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**Comparative Assessment of the Sensitivity of Ozone to Nitrogen Oxides and Volatile  
Organic Compounds in Two Dissimilar Metropolitan Areas of North America:  
Cincinnati, OH (U.S.A.) and Mexico City, DF (Mexico)**

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## ABSTRACT

Because of the complexity of urban ozone ( $O_3$ ) formation, photochemical modeling and ambient data analysis are now needed to provide feedback regarding the effectiveness of  $O_3$  control strategies. However, procedures for diagnosing the  $O_3$  sensitivity to nitrogen oxides ( $NO_x$ ) or volatile organic compounds (VOC) based on ambient data analysis still require further development and testing. This dissertation presents the results of an experimental investigation into the usefulness of using a combination of observational-driven techniques to determine the  $O_3$ - $NO_x$ -VOC sensitivity in urban areas. Afternoon  $O_3/NO_y$ ,  $O_3/NO_z$  and  $NO_y$  (where  $NO_y$  represents the total oxidized nitrogen species and  $NO_z$  represents the reaction products of  $NO_x$ ) photochemical indicators were measured in two dissimilar cities: Cincinnati, Ohio (U.S.A.) and Mexico City, Federal District (Mexico). The evaluation the photochemical indicator analysis diagnosed VOC-sensitive conditions in both cities. The prevalence of these conditions, were confirmed with a complementary statistical analysis of the differences in average weekend/weekday  $O_3$  peak and morning  $NO$  emissions (WE/WD effect). The comparative assessment of the factors determining the  $O_3$ - $NO_x$ -VOC sensitivity in Cincinnati and Mexico City revealed strong differences on population, urban characteristics, topography, precursors emissions densities, and magnitude of VOC emissions, but an apparent similarity in  $NO_x$  emissions between these two urban areas. However, measured maximum  $O_3$  were high in Mexico City as compared to Cincinnati. This phenomenon could be explained as the result of the higher VOC emissions and stronger VOC reactivity in Mexico City than in Cincinnati. A comparative assessment between the photochemical indicators method and the VOC/ $NO_x$  emission ratio to diagnose  $O_3$ - $NO_x$ -VOC sensitivity demonstrated the consistency of the first approach to perform this identification under severe conditions of  $O_3$  formation. The  $O_3$  air quality management implications were, due to the predominating VOC-sensitive conditions, that additional reductions in  $NO_x$  emissions, stipulated by previous model-based  $O_3$  control strategies now in progress in both areas, should be avoided or at least thoroughly reviewed. Otherwise,  $O_3$  levels might increase. This research demonstrates that the combination of the photochemical indicators method with the analyses of the WE/WD effect can be a useful tool in the development and evaluation of  $O_3$  air quality management strategies.



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## GLOSSARY OF ACRONYMS AND SYMBOLS

APC-CAS	Atmospheric Physical Chemistry-Center for Atmospheric Sciences
BUWAL	Bundesamt für Umwelt, Wald und Landschaft (Federal Office for Environment, Forest and Landscape)
CIT	California/Carnegie Institute of Technology
CO	Carbon monoxide
EKMA	Empirical Kinetical Modeling Approach
EPA	Environmental Protection Agency
GC	Greater Cincinnati
$h$	Planck's constant ( $6.6256 \times 10^{-34}$ J·s)
$H_0, H_1$	Terms to describe the statistical null hypothesis and the statistical alternative hypothesis, respectively
H <sub>2</sub> O <sub>2</sub>	Hydrogen peroxyde
HCDOES	Hamilton County Environmental Services
HNO <sub>3</sub>	Nitric acid
HO <sub>2</sub>	Hydroperoxyl radical
HONO	Nitrous acid
HO <sub>x</sub>	Hydrogen-containing free organic radicals
IMP	Instituto Mexicano del Petróleo
ININ	Instituto Nacional de Investigaciones Nucleares
LST	Local Standard Time
M	A third molecule in a chemical reaction, usually oxygen or nitrogen
MCCM	Multiscale Climatic and Chemistry Model
MCMA	Mexico City Metropolitan Area
MoO	Molybdenum oxide
MTBE	Methyl- <i>t</i> -butyl ether
N	Number of samples
NARSTO	North American Research Strategy for Tropospheric Ozone
NBS	National Bureau Standards
NCDC	National Climatic Data Center
NH <sub>3</sub>	Ammonia
NIST	National Institute of Standards and Technology

NMHC	Non-methane Hydrocarbons
NO	Nitrogen oxide
NO*	NO <sub>y</sub> compounds reduced to NO in a thermal catalytic converter
NO <sub>2</sub>	Nitrogen dioxide
NO <sub>2</sub> *	Estimated nitrogen dioxide concentration from empirical numerical approaches
NO <sub>3</sub> <sup>-</sup>	Particulate nitrate
NO <sub>x</sub>	Nitrogen oxides (NO <sub>x</sub> = NO + NO <sub>2</sub> )
NO <sub>x</sub> *	NO <sub>x</sub> measurement performed with a chemiluminescent analyzer
NO <sub>x</sub> -sensitive	A term to define the condition in which lowering NO <sub>x</sub> emissions most effectively reduces peak ozone.
NO <sub>y</sub>	Total reactive nitrogen (NO <sub>y</sub> = NO <sub>x</sub> + NO <sub>z</sub> ), also, total oxidized nitrogen species
NO <sub>y</sub> *	NO <sub>x</sub> analyzer adapted to measure NO <sub>y</sub>
NO <sub>z</sub>	NO <sub>x</sub> reaction products (NO <sub>y</sub> - NO <sub>x</sub> ), also the sum of: PAN + HNO <sub>3</sub> + NO <sub>3</sub> <sup>-</sup> + others
NO <sub>z</sub> *	Estimated NO <sub>z</sub> concentration from empirical numerical approaches
O <sub>3</sub>	Ozone
OBM	Observational-Based Models
OH	Hydroxyl radical
<i>p</i>	Level of significance to express the probability of making a given percent of a Type I error (degree of confidence “alpha”); this error occurs when one reject a hypothesis when it should be accepted.
PAN	Peroxyacetyl nitrate
<i>P<sub>i</sub></i>	Production rate of formation of any “ <i>i</i> ” compound
PIM	Photochemical Indicators Method
ppb	Parts per billion, or parts per 10 <sup>9</sup> , by volume
ppm	Parts per million, or parts per 10 <sup>6</sup> , by volume
ppmC	Parts per million of one organic carbon compound or hydrocarbon compound X number of carbon atoms per molecule
RAMA	Red Automática de Monitoreo Atmosférico (Automated Air Monitoring Network)
RCHO	Carbonyl
RH	Generic hydrocarbon
RO	Alcoxy radical
RO <sub>2</sub>	Alkyl peroxy radical
RO <sub>2</sub> NO <sub>2</sub>	Generic organic peroxy nitrates
RONO <sub>2</sub>	Generic organic nitrates



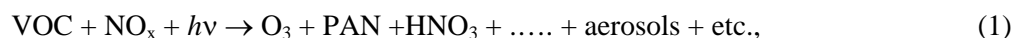
ROOH	Generic higher peroxides
S.D.	Standard Deviation
$S_H$	Summed source of odd hydrogen radicals
SUMS	Santa Ursula Monitoring Station
URG	University Research Glassware (trade mark)
VOC	Volatile organic compounds
VOC/NO <sub>x</sub>	Ratio of emissions of volatile organic compounds to nitrogen oxides
VOC-sensitive	A term to define the condition in which lowering VOC emissions most effectively reduces peak ozone.
WD	Weekday (Monday, Tuesday, Wednesday, Thursday, or Friday)
WE	Weekend day (Sunday or Saturday)
WE/WD effect	Term that refers to the change on peak ozone due to the change on precursors emissions in weekends as compared with the weekdays
3-D model	Three-dimensional photochemical air quality simulation model
$\nu$	Frequency associated with a particular photon (s <sup>-1</sup> )
$\lambda$	Wavelength (nm)
$\mu$	Population mean; also, micro or 1 x 10 <sup>-6</sup>
$\mu\text{g}$	micrograms, 1 x 10 <sup>-6</sup> g

## Chapter 1

# INTRODUCTION

### 1.1 Background and Motivation

The design of effective control strategies for tropospheric ozone ( $O_3$ ), whose harmful effects on human health, forest ecosystems, agricultural crops, and materials are well known, still represents a complex scientific, technological, and social issue, in both developed and developing countries (Kley et al., 1999; Molina and Molina, 2004).  $O_3$  is a secondary pollutant formed in the ambient air as a result of a complex, nonlinear function of: (1) the relative proportion of the emissions of its precursors, primary emissions of nitrogen oxides ( $NO_x = NO + NO_2$ , where NO is nitrogen oxide and  $NO_2$  is nitrogen dioxide) and volatile organic compounds (VOCs) from anthropogenic and natural sources; (2) the reactivity of the VOC; (3) the advective transport of these emissions and their reaction products; (4) the temperature, humidity, and solar radiation; and (5) the complex set of sunlight initiated chemical reactions occurring in the troposphere concurrent with the transport and dispersion of the emissions (US EPA, 1996). These processes lead to the formation of  $O_3$  and other photochemical oxidants, such as peroxyacetyl nitrate (PAN), nitric acid ( $HNO_3$ ), and aerosol nitrates ( $NO_3^-$ ). The overall reaction of this system can be written as:



The complex dependence of tropospheric ozone formation on the availability of its precursors  $NO_x$  and VOC gives origin to the so-called  $O_3$ - $NO_x$ -VOC sensitivity term. Under  $NO_x$ -sensitive conditions (high VOC/ $NO_x$  ratios), ambient  $O_3$  can be effectively reduced by reductions on  $NO_x$  emissions and is, practically, not sensitive to VOC changes. Under VOC-sensitive conditions, reductions on anthropogenic VOC emissions (low VOC/ $NO_x$  ratios) will be effective to reduce  $O_3$  but can also increase peak ozone concentrations if  $NO_x$  emissions are decreased (Finlayson-Pitts and Pitts, Jr., 2000). 3-D photochemical models are useful to predict the features of this complex  $O_3$ - $NO_x$ -VOC chemistry, but the accuracy of the

predictions that determine whether O<sub>3</sub> formation during specific circumstances is associated with NO<sub>x</sub>-sensitive chemistry or VOC-sensitive conditions is still limited and unknown (Blanchard and Stoeckenius, 2001; Chameides et al., 2000). Because similar quantities of O<sub>3</sub> can be predicted in VOC-sensitive and NO<sub>x</sub>-sensitive environments and because of the possibility of compensating errors in photochemical models, one photochemical model can acceptably predict diurnal O<sub>3</sub> but incorrectly predict the sensitivity of O<sub>3</sub> to changes in NO<sub>x</sub> and VOC emissions (Sillman, 1995; Toneesen and Dennis, 2000). Moreover, there is also no direct way to test whether the NO<sub>x</sub>-VOC predictions are accurate—a central concern when designing O<sub>3</sub> control strategies.

Despite the considerable analytical power that photochemical models offer to the policy maker to define an O<sub>3</sub> air quality strategy, a key limitation of these models is that they are driven by highly uncertain emission inventories (Sillman, 1999; Chameides, et al. 2000). Major inaccuracies in simulation results have been associated with discrepancies in current NO<sub>x</sub> and VOC emission inventories (expressed as the ratio VOC/NO<sub>x</sub>).

The failure of models to accurately predict the response of O<sub>3</sub> to reductions in NO<sub>x</sub> and VOC emissions has resulted in considerable skepticism about the use of these models as basis for environmental policy (Chameides et al., 2000; Finlayson-Pitts and Pitts, Jr., 2000). However, the design of control strategies still relies on the predictions of photochemical models. For example, with recent court approval of the new U.S. Environmental Protection Agency (EPA) 8-hr Ozone Standard, state and local agencies must improve emissions inventories, develop or strengthen control strategies, and demonstrate, with photochemical modeling, that the proposed strategies will be sufficient to attain and maintain the standard (Wierman, 2003).

In recognition of the above limitation, the North American Research Strategy for Tropospheric Ozone group (NARSTO) has pointed out the necessity of improving the present emission-based O<sub>3</sub> air quality management process. To achieve this, they have recommended using all available information to determine whether a control strategy is expected to achieve a specific standard or whether a current model-based O<sub>3</sub> control strategy is appropriately functioning (Chameides et al., 2000).

The cases of Cincinnati, OH (U.S.A.) and Mexico City, DF (Mexico) provide a neat illustration of the complexity of this situation. Cincinnati, typical of medium sized U.S. metropolitan areas, is in compliance with the EPA's 1-hr O<sub>3</sub> standard but has the potential to be noncompliance with the new, more restrictive 8-hr O<sub>3</sub> standard. The O<sub>3</sub> control strategies imposed in this region were the result of recommendations from a multi-state photochemical modeling effort undertaken in order to control ozone regional transport over the East and Midwest of the U.S. These strategies have focused mainly on the reduction of NO<sub>x</sub> emissions, as if the whole region (East and Midwest) were under a NO<sub>x</sub>-sensitive environment (Guinnup and Collom, 1997; The Ohio Environmental Council, 2002; US EPA, 2004).

The Mexico City Metropolitan Area (MCMA) is an example of an urban metro-complex in a developing country that still has yet to find the appropriate control strategy to reduce the observed daily O<sub>3</sub> levels. Past and present O<sub>3</sub> control strategies have been the result of recommendations from preliminary photochemical modeling efforts and linear programming analysis combined with political and economical assessments (Streit and Guzman, 1996; Molina and Molina, 2002; McKinley et al., 2003). Although the non-systematic combination of both NO<sub>x</sub> and VOC controls has been common issue in the several control plans enforced in the past 15 years, the general recommendation to policymakers has been to control NO<sub>x</sub>-emissions, as if it was a case of NO<sub>x</sub>-sensitive conditions (Molina and Molina, 2002).

Two central questions arise from these examples. How can the effectiveness of ongoing O<sub>3</sub> control strategies be scientifically demonstrated? If determined to be ineffective, what basis is there for recommending changes to a given control strategy? One possible way of answering these questions is an objective analysis of the O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity based on ambient air-quality measurements.

In recognition of the limitations of photochemical models, scientists have developed a new category of analytical tools. These tools are the so-called "observation-based methods" (OBMs). These methods use ambient air-quality observations to infer important relationships between O<sub>3</sub>, its precursor compounds, and/or the sources of these precursor compounds. These tools differ significantly from emission based methods (i.e., photochemical models) that use emission inventories to simulate photochemical scenarios (Blanchard et al., 1999; Sillman, 1999; Chameides et al., 2000). OBMs offer

several general advantages over photochemical models for assessing the sensitivity of  $O_3$  to VOC and  $NO_x$  in a given air mass. Because OBMs rely on ambient measurements (real-world conditions) and do not depend on the accuracy of emission inventories, they provide a way to evaluate the consistency of typical model predictions and thus reduce the level of uncertainty associated with these predictions. In turn, this improves the characterization of  $O_3$  precursor relationships. Sillman (2002) has suggested that the combination of OBMs with photochemical models can provide an excellent link between photochemical models and ambient measurements that is often missing in typical modeling analysis.

However, OBMs have some operational drawbacks, as they depend on measurements of relevant constituents (Chameides et al., 2000). If the data are inaccurate or incomplete, there is the potential for uncertain or erroneous results. There is also the potential for misinterpretation if the OBM is not tested for uncertain assumptions. Because OBMs are diagnostic approaches, they do not predict the future state of the atmosphere (e.g., the peak  $O_3$  concentration resulting from a given reduction in pollutant emissions).

OBMs include receptor models, regression techniques, ambient ratios, indicator species, and semi-empirical based OBMs. From among the various indicator species techniques, the photochemical indicators method (PIM) has emerged as the most practical for evaluating the  $O_3$ - $NO_x$ -VOC sensitivity. It does so directly from an analysis of correlations and ratios between measured secondary photochemical products such as  $O_3$ , total reactive nitrogen ( $NO_y$ ), and other key secondary species (Sillman et al., 1997; Sillman, 1999; Sillman and He, 2002). This method has also been used to evaluate the consistency of photochemical models currently used to set up  $O_3$  control policies.

However, the acceptance of the PIM is not universal. The NARSTO group has cited the need of further confirmation and development of these complementary tools, from both a policy and a scientific/technical standpoint, before such iterative scheme process could be fully implemented (Chameides et al., 2000). To achieve a wider acceptance, two important disadvantages of OBMs must be tackled. First, their sources of errors must be clarified, and, second, there needs to be more documented experiences to develop appropriate standard procedures.

The present applied research has three motivations. The first is to investigate the expediency of combining the photochemical indicators approach with other data analysis techniques for assessing current O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity, instead of using a single insulated OBM. The second is to demonstrate the usefulness of the photochemical indicators to explain confusing NO<sub>x</sub>-VOC sensitive situations predicted by other modeling and measurement methods. The third is to provide research grade evidence of the convenience of observational-measurement analyses to support O<sub>3</sub> air quality management policies.

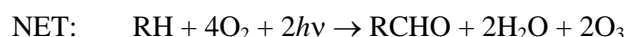
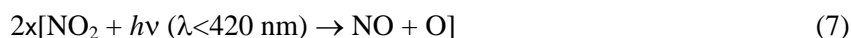
## 1.2 Overview of Ozone-NO<sub>x</sub>-VOC Sensitivity

Ozone production occurs via the free-radical-initiated oxidation of VOC or CO in the presence of NO<sub>x</sub> and sunlight (Chameides et al., 1992). The solar radiation dissociates stable molecules present in the atmosphere (i.e., carbonyls or aldehydes and hydrogen peroxide) to form hydrogen-containing free radicals (HO<sub>x</sub>), which, in the presence of NO<sub>x</sub>, catalyze the oxidation of VOCs to form, ultimately, carbon dioxide and water vapor. Partially oxidized organic species such as aldehydes, ketones, and carbon monoxide are produced as intermediate oxidation products, with O<sub>3</sub> formed as a by-product. The typical chemical mechanism for the oxidation of a generic hydrocarbon (RH) is represented by Reactions (2) through (6):



In this set, R is a hydrocarbon alkyl radical chain (for example, CH<sub>3</sub>CH<sub>2</sub>CH<sub>2</sub>-H will stand for RH = propane), and M is a third molecule (N<sub>2</sub> or O<sub>2</sub>). RO represents an alkoxy radical, and RO<sub>2</sub> an alkyl peroxy radical. RCHO represents a carbonyl. Alkyl peroxy radicals reactions (R4) with NO reduce the concentration of NO and increase the concentration of NO<sub>2</sub>. Reaction (6) regenerates hydroxy (OH)

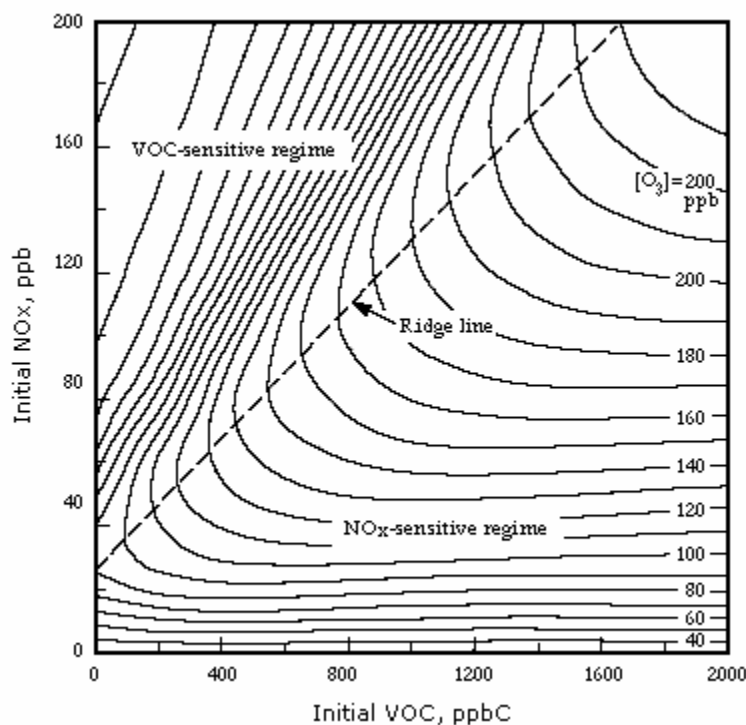
radicals and also produces NO<sub>2</sub>. This mechanism can be interpreted as a catalytic cycle, where radicals (OH, R, RO<sub>2</sub>, RO, and HO<sub>2</sub>) act to propagate a chain of additional radicals. Reactions (4) to (6) play a key role in O<sub>3</sub> formation by oxidizing NO to NO<sub>2</sub>. Subsequently, photodissociation of NO<sub>2</sub> by sunlight generates O<sub>3</sub> by Reactions (7) and (8):



As shown in the net oxidation set of reactions, one molecule of RH (i.e., one cycle) is accompanied by the generation of two molecules of ozone. The carbonyl compound produced in the sequence (RCHO in reaction (5) and net reaction) is, in general, subject to further oxidation that can lead to the production of additional O<sub>3</sub>. In this sequence, VOCs are consumed, while both OH/HO<sub>2</sub> and NO<sub>x</sub> act as catalysts. Termination occurs when the catalysts are removed, often by one of the two paths:



The predisposition for O<sub>3</sub> formation in the air is essentially proportional to the number of free-radical propagated cycles that can occur before radical removal (Jenkin and Clemitshaw, 2000). This is usually referred to as the chain for O<sub>3</sub> formation. Therefore, when considering the influence of changing the concentration of NO<sub>x</sub> or VOC on the production of O<sub>3</sub>, the effect of such a change on the chain-length is very important. This length is determined by the rate of the chain terminating reactions as compared to the competing chain propagating reactions. Oxidation mechanisms for olefins, which react with O<sub>3</sub> as well as OH radicals, aldehydes, which photodissociate into free radicals and react with OH, and aromatics are more complex but have similar chain initiation, carrying, and termination steps. The most important features of an O<sub>3</sub> abatement strategy are illustrated in Figure 1.1.



**Figure 1.1.** Typical peak ozone isopleths diagram. Peak  $O_3$  concentrations in parts per billion (by volume) as a function of various initial VOC and  $NO_x$  concentrations. Ozone isopleth based on simulation of chemistry along air trajectories in Atlanta. (Adapted from Jefferies and Crouse, 1990).

Figure 1.1 shows the typical model-calculated isopleths of peak  $O_3$  concentrations. Such isopleths have been commonly generated using computer models, such as the Empirical Kinetic Modeling Approach (EKMA), where the results of the model are tested against smog chamber data (Finlayson-Pitts and Pitts, Jr., 2000). Although EKMA plots have been replaced by sophisticated 3-D photochemical models, they still give valuable information about the expected peak  $O_3$  levels as a function of initial VOC and  $NO_x$  concentrations or as a function of the amounts of VOC or  $NO_x$  that are added to the system as volumetric emissions. The diagonal line that extends from the lower left to the upper right corner of the graph has a slope corresponding to a VOC/ $NO_x$  ratio of about 8:1. The OH radical chain length, equal to the number of times a newly formed OH radical is regenerated through radical chain propagation before it is destroyed, reaches a maximum at this VOC/ $NO_x$  ratio. Thus, the ridgeline corresponds to the VOC/ $NO_x$  ratio, at which  $O_3$  is most efficiently formed.



At low VOC/NO<sub>x</sub> ratios to the left of the ridgeline, a decrease in the concentration of RH, would decrease the chain length and hence the rate of O<sub>3</sub> formation. The rate of formation of HNO<sub>3</sub> surpasses the rate of formation of H<sub>2</sub>O<sub>2</sub> and O<sub>3</sub> production is *VOC-sensitive* (i.e., lowering VOC emissions most effectively reduces peak O<sub>3</sub>). The OH radicals that propagate VOC oxidation and NO-to-NO<sub>2</sub> conversion are scavenged by the relatively high concentrations of NO<sub>x</sub>. However, a decrease in the concentration of NO<sub>2</sub>, which might result from NO<sub>x</sub> emission reductions, would increase the chain length, as more of the OH pool is available to react with the VOCs (Jenkin and Clemitshaw, 2000). The reduction of NO<sub>x</sub> lowers the rate at which OH and NO<sub>2</sub> are removed by formation of HNO<sub>3</sub> and leads to an increase in maximum O<sub>3</sub>. This situation is frequently named “NO<sub>x</sub>-disbenefit” and occurs only in the VOC-sensitive region.

For high VOC/NO<sub>x</sub> ratios to the right of the ridgeline, lowering NO<sub>x</sub> concentrations with a constant VOC concentration, results in an effective reduction on peak O<sub>3</sub> concentrations. However, decreasing VOCs alone at constant NO<sub>x</sub> results only in a slowly decreasing O<sub>3</sub> peak levels. At these high VOC/NO<sub>x</sub> ratios, O<sub>3</sub> formation is said to be *NO<sub>x</sub>-sensitive* (i.e., lowering NO<sub>x</sub> emissions most effectively reduces peak O<sub>3</sub>). In this region of the isopleths on Figure 1.1, the source of OH radicals is greater than the source of NO<sub>x</sub>. The chain propagating reactions (4) and (6) involve NO<sub>x</sub>, and, consequently, any reduction in NO<sub>x</sub> decreases the O<sub>3</sub> formation chain length. The rate of formation of H<sub>2</sub>O<sub>2</sub> is greater than the rate of formation of HNO<sub>3</sub>. Since RH is not itself directly involved in these competitions, the O<sub>3</sub> formation chain length is insensitive to changes in the concentrations of RH, which might result from VOC control.

It is important to point out that variation in the concentration of some reactive VOCs has an effect on O<sub>3</sub> formation under NO<sub>x</sub>-sensitive conditions, because the degradation of some VOCs can lead to a significant removal of NO<sub>x</sub> as organic nitrates and peroxy nitrates (PANs) as it is shown with general reactions (11) and (12) (Jenkin and Clemitshaw, 2000).

Reaction (12) is important in the photochemical formation process because can act as a temporal reservoir of NO<sub>2</sub> and as a sink of RO<sub>2</sub> radicals.



Finally, between the  $\text{NO}_x$ - and VOC-sensitive regions (the ridgeline zone) lies a transition region in which  $\text{O}_3$  is equally sensitive to VOC and  $\text{NO}_x$ , but, compared to its sensitivity to VOC on the left side of the ridgeline and its sensitivity to  $\text{NO}_x$  on the right, it is relatively insensitive to both.

### 1.3 Observation-Based Methods for Evaluating $\text{O}_3$ - $\text{NO}_x$ -VOC Sensitivity

As previously discussed, OBM refers to a number of techniques that can be used to translate field observations into information on  $\text{O}_3$  production. These techniques qualitatively assess the effectiveness of reductions of emissions of VOCs versus  $\text{NO}_x$ . OBMs offer several advantages as an analytical tool for evaluating the consistency of  $\text{O}_3$  regulatory policies (Sillman, 2002). Examples of OBMs include: smog production (extent-of-reaction) algorithms (Chang and Suzio, 1995, Chang et al., 1997; Blanchard et al., 1999; Blanchard 2000); observation-based models using  $\text{NO}_x$  and VOC concentrations, meteorological data, and the concept of incremental reactivity (Cardelino and Chameides, 1995); constrained steady state (Kleinman 2000; Kleinman et al., 2001; Tonnesen and Dennis, 2000); and the photochemical indicators method (Sillman 1995; Sillman et al., 1997; Sillman and He, 2002).

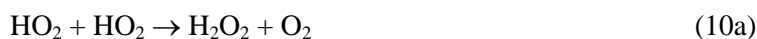
The typical measurements required by most OBMs are:  $\text{O}_3$ , total reactive nitrogen ( $\text{NO}_y$ ),  $\text{NO}$ ,  $\text{CO}$ , and speciated VOC. Some of these methods actually combine two different techniques, such as modeling with one observational analytical technique. A detailed analysis of the theoretical bases, advantages, and disadvantages of these OBMs is given in Kleinman (2000) and Sillman (2002).

OBMs are classified into two broad categories: “present time-frame methods,” based on measurements of present ambient species such as VOC,  $\text{NO}_x$ , and  $\text{CO}$ , and “past time-frame methods,” based on measurements of secondary photochemical products (Kleinman, 2000). The photochemical indicators method (PIM) makes use of measurements of secondary reaction products. These measurements are absolute values of selected species or species ratios that have been found as consistent indicators of the

NO<sub>x</sub>-VOC sensitivity by analyzing 3-D photochemical simulations results under different conditions of NO<sub>x</sub> and VOC hypothetical emissions reductions. An example of one photochemical indicator is the ratio of the peak O<sub>3</sub> concentration (representing the O<sub>3</sub> that has built up since the precursor emissions occurred) to that of another secondary species (such as NO<sub>z</sub>, the amount of NO<sub>x</sub> that has been oxidized to higher nitrogen oxides or the sum of HNO<sub>3</sub>, PAN, NO<sub>3</sub><sup>-</sup>, and others). The species ratios can then provide clues as to how an air mass upwind of a receptor location has evolved to its present state.

The species ratios O<sub>3</sub>/NO<sub>y</sub>, (where NO<sub>y</sub> = NO<sub>x</sub> + NO<sub>z</sub>), O<sub>3</sub>/NO<sub>z</sub>, O<sub>3</sub>/HNO<sub>3</sub>, H<sub>2</sub>O<sub>2</sub>/HNO<sub>3</sub>, and other equivalent species correlations or individual species (i.e., NO<sub>y</sub> as a whole) show different values depending on whether O<sub>3</sub> was built up under a NO<sub>x</sub>- or VOC-sensitive environment. This approach is based on 3-D modeling results obtained by Milford et al. (1994), Sillman (1995, 1999), and Sillman et al. (1997, 1998). These authors found that, in general, high values of these ratios (and low values of NO<sub>y</sub>) were associated with NO<sub>x</sub>-sensitive chemistry and low values of the same ratios (and high values of NO<sub>y</sub>) were associated with VOC-sensitive chemistry.

The theoretical justification for the main indicator ratios is based on the assumption that the split between NO<sub>x</sub>-sensitive and VOC-sensitive regimes is driven by the chemistry of odd hydrogen radicals (OH, HO<sub>2</sub>, and RO<sub>2</sub>) (Sillman et al., 1990; Sillman, 1995). This split has been attributed to the relative strengths of odd hydrogen radical sources ( $S_H$ ) and the production rates of its major sinks (peroxides, nitric acid and PAN):  $S_H = 2P_{perox} + P_{HNO_3} + P_{PANs}$ . VOC-sensitive chemistry would occur whenever  $P_{HNO_3}$  (and in a lower extent  $P_{PANs}$ ) exceeds  $2P_{perox}$ . Under this condition, the production rate of O<sub>3</sub> is assumed to be proportional to the rate of the VOC + OH reactions (reaction 2), and its rate increases with increasing VOC and decreases with increasing NO<sub>x</sub>. By contrast, NO<sub>x</sub>-sensitive conditions would occur when the peroxide-forming Reactions (10a) and (13) are the dominant radical sinks:



Under this condition, the production rate of  $O_3$  is assumed to be approximately equal to the rate of Reactions (4) and (6) ( $RO_2 + NO$ ,  $HO_2 + NO$ ). This rate increases with increasing  $NO_x$  but has little direct dependence on VOC.

In this way, the ratio  $O_3/NO_z$  and the correlation between  $O_3$  and  $NO_z$  could be interpreted as the rate of production of  $O_3$  to the rate of production of production of  $NO_x$  reaction products. This ratio has been interpreted as the “ozone production efficiency per  $NO_x$ ” factor (OPE) in the analysis of rural events (Trainer et al., 1993). The ratio  $O_3/NO_y$  and the correlation between  $O_3$  and  $NO_y$  have several interpretations. They reflect two separate processes, both of which are related to  $O_3$ - $NO_x$ -VOC sensitivity. They reflect the process of photochemical production of  $O_3$  (represented by the ratio  $O_3/NO_z$ ) and the process of NO titration from fresh  $NO_x$  emissions (represented by the  $NO_x$  fraction in the total  $NO_y$ ). Afternoon high  $O_3$  and low  $NO_y$  would reflect relatively large radical sources and smaller radical sinks such as the  $NO_2 + OH$  reaction. This ratio is assumed also to indirectly account for photochemical aging. The ratio  $O_3/NO_y$  can be viewed as the ratio  $O_3/NO_z$  multiplied by an aging term ( $NO_z/NO_y$ ) (Sillman and He, 2002). In photochemically young air masses the ratio  $NO_z/NO_y$  is close to 0. In photochemically old air masses a large fraction of  $NO_x$  has been processed to  $NO_z$  and under these conditions the  $NO_z/NO_y$  ratio is close to 1. On the other hand, Milford et al. (1994) and Sillman (1995) have shown that  $NO_y$  is a good empirical indicator of high precursor concentrations because acts as a measure of  $NO_x$  emissions through the period of photochemical activity and also reflects the impact of stagnant meteorology on  $NO_x$ -VOC sensitivity and on the photochemical production of  $O_3$ .

A number of reference values have been suggested for delineating the transition from VOC to  $NO_x$  sensitivity. These values are obtained from 3-D photochemical models in which the input  $NO_x$  and VOC emissions were varied systematically to predict an  $O_3$  reduction in response to reduced  $NO_x$  and anthropogenic VOC. Table 1.1 shows some of the suggested transition values for different  $O_3$  polluted conditions in the U.S.

The PIM allows one to identify  $NO_x$ - and VOC-sensitive conditions between different locations or for different smog events (Sillman, 2002). Comparisons of 3-D model predictions and measured values

have demonstrated that the photochemical model can accurately infer the appropriate VOC- or NO<sub>x</sub>-sensitive chemistry. Alternatively, when 3-D model predicted correlations between O<sub>3</sub> and NO<sub>y</sub> and similar indicator species are compared with measured values, the resulting scatterplot patterns can be interpreted as evidence of bias in the photochemical model to interpret the right VOC-sensitive or NO<sub>x</sub>-sensitive chemistry.

**Table 1.1.** Values of selected photochemical indicators and their NO<sub>x</sub>-VOC sensitive transitional range.

Indicator	Transition range	Reference
NO <sub>y</sub>	10 to 25 ppb <sup>a</sup>	Milford et al., 1994; Sillman, 1995
Ozone/NO <sub>y</sub>	11 to 15 (when [O <sub>3</sub> ] < 80 ppb) <sup>b</sup> 6 to 8 (when 80 ppb < [O <sub>3</sub> ] < 200 ppb) <sup>b</sup>	Sillman and He, 2002
Ozone/NO <sub>z</sub>	15 to 20 (when [O <sub>3</sub> ] < 80 ppb) <sup>b</sup> 8 to 10 (when 80 ppb < [O <sub>3</sub> ] < 200 ppb) <sup>b</sup>	Sillman and He, 2002
Ozone/Nitric Acid	20 to 25 (when [O <sub>3</sub> ] < 80 ppb) <sup>b,c</sup> 12 to 15 (when 80 ppb < [O <sub>3</sub> ] < 200 ppb) <sup>b,c</sup>	Sillman and He, 2002
Hydrogen Peroxide/Nitric Acid	0.2 to 0.3 <sup>b</sup>	Sillman and He, 2002
Hydrogen Peroxide/NO <sub>y</sub>	0.12 to 0.17 <sup>b</sup>	Sillman and He, 2002
Hydrogen Peroxide/NO <sub>z</sub>	0.20 to 0.25 <sup>b</sup>	Sillman and He, 2002

<sup>a</sup> Ratios lower than the transition range generally correspond to NO<sub>x</sub>-sensitive conditions.

<sup>b</sup> Ratios lower than the transition range generally correspond to VOC-sensitive conditions.

<sup>c</sup> Interpolated from Figure 3 in Sillman and He (2002).

In order to use the method in field applications, several requirement must be fulfilled: (1) photochemical indicators can only be collected, modeled, and/or used for the afternoon period (between 13:00 and 17:00 hr) during relatively sunny days and without rain; (2) measurements of indicator species must be performed in the peak or near-peak O<sub>3</sub> regions of a given airshed; and (3) NO<sub>y</sub> measurements must have uncertainties of less than 20% (preferably less than 10%) and must include HNO<sub>3</sub> species (Sillman, 2002). If measurements of photochemical indicators are available for a given airshed, the method itself can work as a “rule of thumb” to identify the NO<sub>x</sub>- and VOC-sensitive regions. However, a

thorough application of the method includes testing with a 3-D photochemical model, as well as the evaluation of measured species correlations patterns (i.e.,  $O_3$  versus  $NO_y$  or  $O_3$  versus  $NO_z$ ) in comparison with model results.

Locations of successful applications of the method include: the northeast corridor, Lake Michigan and Atlanta (Sillman, 1995); New York and Los Angeles (Sillman et al., 1997); Nashville (Sillman, et al., 1998); Switzerland (Staffelbach et al., 1997, Dommen et al., 1999); Milan, Italy (Martilli et al., 2002; Hammer et al., 2002); and Paris, France (Sillman et al., 2003). Some of these applications have included both measurements and model-measurement comparisons for the relevant species. Although several different photochemical mechanisms were used in these 3-D model simulations, the results were consistent with photochemical indicator theory.

However, evidence running contrary to this method has been reported. Model-based investigations of the San Joaquin Valley (Lu and Chang, 1998) and Los Angeles (Chock et al., 1999) both gave results different from those reported by Sillman. Blanchard and Stoeckenius (2001) also reported opposing results in a review of OBM applications. Results from Paris, France (Sillman et al., 2003) revealed situations where measurements resulted in poor accuracy. Sosa et al. (2000) obtained contradictory results in their study of Mexico City. However, there is evidence that most of the above disagreements were associated either with problems in the input emissions inventories and in the treatment of deposition mechanisms in the models or with erroneous measurements (Sillman, 2002). Consequently, more research is required to confirm the universality of the method, especially in situations where the PIM has failed to produce consistent results.

## **1.4 Scope of the Research**

### **1.4.1 Approach and Methodology**

The overall strategy envisioned in this applied research consisted of sequential stages including: field experiments; statistical, graphical, and conceptual evaluations; and assessments of the consistency of the results and of the PIM application with modeling experiments and published results. The idea behind

this strategy was to combine the assessment of the particular conditions linked to the  $O_3$ - $NO_x$ -VOC sensitivity of two dissimilar environments with the investigation of factors that may affect, improve and demonstrate the consistency of the PIM.

The first stage focused on obtaining data on photochemical processes related species for each of the two study sites. These data included measurements of  $O_3$ ,  $NO$ , and  $NO_y$ , in Cincinnati and Mexico City during the smog season of each region. Additional measurements of  $HNO_3$  and  $H_2O_2$  were made in Cincinnati only. Several methods recommended by the EPA were used for ozone and the nitrogen oxidized species. An experimental procedure published in peer-reviewed literature was used for the measurement of  $H_2O_2$ . Quality assurance and quality analytical procedures recommended by the EPA were followed during the monitoring and sampling campaigns.

The second stage consisted in the diagnosis of the  $O_3$ - $NO_x$ -VOC sensitivity for Cincinnati and Mexico City using the PIM. Several statistical, graphical, and conceptual evaluations were performed on the observed data. In both cases, empirical correlations from peer-reviewed literature or from observed data were used to obtain either an indirect corroboration of a given assumption or to obtain a surrogate of species concentrations not measured during the field campaigns. A combination of graphical and numerical comparison procedures recommended by Sillman (2002) was used to evaluate the consistency of the results.

The central objective of the third stage was the corroboration of the diagnosis previously generated with the results of a statistical analysis of weekend/weekday differences in observed peak  $O_3$  and morning  $NO$  concentrations. A corroboration of the diagnosis with the results of an evaluation of the duration of morning-to-afternoon  $O_3$  accumulation and the rate of  $O_3$  accumulation was also undertaken. The confirmation of the diagnosis for Mexico City was extended with the inclusion of a comparison of the measured photochemical indicators with the results of an  $O_3$ - $NO_x$ -VOC sensitivity analysis previously performed for this area. An intercomparison between the results of the PIM diagnosis in Mexico City with two photochemical models previously used to study the local  $O_3$  control strategies was also performed.

The last stage of the research consisted in a comprehensive assessment of the factors associated with the specific O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity in each city making use of the photochemical indicators results and several empirical O<sub>3</sub> formation concepts. This assessment included an evaluation of the photochemical indicators results with regard to the VOC/NO<sub>x</sub> ratio in the identification of the NO<sub>x</sub>-VOC sensitivity.

### **1.4.2 Research Goals**

Since ozone control strategies require predictions that are difficult to test with present photochemical models, and because observation-based methods represent a practical option for evaluating predicted O<sub>3</sub> control management scenarios, this research investigated the consistency and usefulness of a combination of two OBMs for evaluating O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity in two dissimilar urban areas. In doing so, the goal was to make a scientific/technical contribution to the improvement of O<sub>3</sub> air quality management practices within the context of the NARSTO recommendations for North America.

The specific objectives of the research were as follows: (1) to explore the expediency of combining the PIM with the weekend-weekday effect analysis to improve the accuracy of O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity diagnosis; (2) to characterize the O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity of Cincinnati and Mexico City and evaluate the consistency of current O<sub>3</sub> control strategies in both urban areas; (3) to investigate the usefulness of the photochemical indicator approach to identify inconsistencies of the VOC/NO<sub>x</sub> ratio to diagnose NO<sub>x</sub>-VOC sensitivities; and (4) to generate data for the enrichment of the PIM database towards the development of standard procedures useful in a regulatory context.

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## Chapter 2

# EXPERIMENTAL METHODS

### 2.1 Measurement of Indicator Species

The program for the measurement of ambient air species in Cincinnati and Mexico City was designed to provide the aerometric databases needed to apply the photochemical indicators method (PIM) and to corroborate the consistency of the application of this method. Continuous measurements of  $O_3$ ,  $NO$ , and total  $NO_y$  were made in the near-peak  $O_3$  region of each of the two cities, while 1-hr  $H_2O_2$  and  $HNO_3$  measurements were made only in Cincinnati during the afternoon hours (13:00-17:00 hr). As a result of a review of previous studies on the dynamics of the photochemical air pollution in both cities (see chapter 3), it was assumed that the measurements done at only one site within the near-peak  $O_3$  region of each of these urban areas would be representative of the  $O_3$ - $NO_x$ -VOC chemistry of their airsheds.

The measurements were performed at the times of year in which ozone reached its highest levels (summer for Cincinnati and winter for Mexico City). The study and analysis of days with high  $O_3$  concentrations provides elements appropriate to infer the interrelationship between precursor emissions and meteorology associated with the occurrence of these “event” days. In addition, the design of  $O_3$  control strategies has been traditionally based on the selection of days with higher ozone concentrations (US EPA, 1996).

The selection of the methods for the continuous measurement of  $O_3$  and  $NO_y$  species was done following the recommendations of Parrish and Fehsenfeld (2000) and of the US EPA (US EPA, 2001). Methods for the sampling and analysis of  $H_2O_2$  and  $HNO_3$  were selected after a review of the specialized literature (Sakugawa et al., 1990; US EPA, 1999; Parrish and Fehsenfeld, 2000). Table 2.1 shows a summary of the methods used for monitoring and sampling of the several indicator species.

The methods used by commercial analyzers for measuring  $O_3$  and  $NO$  are well accepted to obtain ambient concentrations of these species (Parrish and Fehsenfeld, 2000). Therefore, the instruments used in the experiments did not require special considerations.

**Table 2.1.** Summary of monitoring and sampling methods for the measurement of indicator species.

Compound	Type of measurement	Method	References	Reported detection limit
NO	Continuous	Chemiluminescence	US EPA, 2001	15 ppt
NO <sub>y</sub>	Continuous	Thermal conversion with converter located near the sample inlet/chemiluminescence	US EPA, 2001	15 ppt
O <sub>3</sub>	Continuous	Ultraviolet absorption photometry	Parrish and Fehsenfeld, 2000	2-5 ppb
HNO <sub>3</sub>	1-hr integrated sample	Na <sub>2</sub> CO <sub>3</sub> coated annular denuder/analysis by ion chromatography	US EPA, 1999	0.78 ppb for 1-hr at a 10 L/min sampling
H <sub>2</sub> O <sub>2</sub>	1-hr integrated sample	Ti(IV)-H <sub>2</sub> SO <sub>4</sub> coated annular denuders/analysis by spectrophotometry with a ferrous thiocyanate reagent at 475 nm	Possanzini et al., 1988	0.037 ppb for 1-hr at a 10 L/min sampling

However, measurement of NO<sub>y</sub> species is not straightforward and requires of a number of special adaptations. As previously mentioned, NO<sub>y</sub> species consist primarily of NO, NO<sub>2</sub>, PAN, HNO<sub>3</sub>, particulate nitrate (NO<sub>3</sub><sup>-</sup>), and nitrous acid (HONO), roughly in that order of importance (Finlayson-Pitts and Pitts, Jr., 2000). NO<sub>y</sub> is commonly measured by passing the airstream containing NO and the other oxides of nitrogen over a catalyst to convert the oxides into NO. Commercial heated molybdenum converters have been shown to readily reduce these species to NO, although the reduction efficiency for particulate nitrate may vary with the converter design (US EPA, 2001). The conversion is accomplished with the converter operating at 325 to 375°C. Above 375 degrees, the converter begins to convert other gases (e.g., ammonia), although the conversion has been reported at less than a few percent under typical operating conditions (Parrish and Fehsenfeld, 2000).

Plumbing changes or adaptations to standard NO<sub>x</sub> sampling configurations in commercial NO<sub>x</sub> analyzers allows one to maximize NO<sub>y</sub> sample collection (Fitz, 2002). These alterations include external placement of the heated converter, shorter non-reactive sampling inlet lines to the converter, moving the location of the particle filter at the exit of the converter, and adding a larger capacity pump. Under this re-

arrangement of the measuring system, the  $\text{HNO}_3$  and aerosol nitrate are not lost prior to the converter and they as well as other oxidized nitrogen species (i.e., PANs) are effectively reduced in the external converter to  $\text{NO}^*$  (where  $\text{NO}^*$  represents all of the  $\text{NO}_y$  compounds reduced to NO). Because the  $\text{NO}_x$  analyzer works on the principle that chemiluminescence emitted by  $\text{NO}_2$  is directly proportional to its concentration, the measured  $\text{NO}^*$  is taken as equivalent to the  $\text{NO}_y$  concentration. Results of an intercomparison of several  $\text{NO}_y$  measurement techniques conducted at a suburban site of Nashville, Tennessee (Williams et al., 1998), indicated that these instruments can reliably measure  $\text{NO}_y$  in typical urban and suburban environments (Parrish and Fehsenfeld, 2000). However, several concerns have been mentioned on the possibility of having measurement errors depending on the history of the converter's exposure (Fitz et al., 2003).

For the purposes of this research,  $\text{NO}_y$  converters were placed outside, leaving the instrument detector within the air-conditioned monitoring shelter. This was done to insure that the sampling lines did not absorb nitric acid (Fitz et al., 2003). In addition, the  $\text{NO}_y$  sampling was conducted using the shortest possible PFA Teflon line to the converter in order to prevent the adsorption (and loss) of nitric acid. The filters to protect the instruments of  $\text{MoO}$  dust were located at the exit of the converters insuring in this way the conversion of  $\text{NO}_3^-$  aerosols. In the Cincinnati and Mexico City monitoring campaigns, a commercial  $\text{NO}_x$  analyzer was adapted to the above specifications in order to measure the  $\text{NO}_y$  species.

The ideal method of measuring  $\text{NO}_y$  would be to measure each of the contributing species to the total expected  $\text{NO}_y$ . However, this effort requires specialized instruments currently only available in a few research centers. In several prior field studies, the sum of the concentrations of the individual  $\text{NO}_y$  species has been found to be less than the measured total  $\text{NO}_y$  using conversion to NO (Finlayson-Pitts and Pitts, Jr., 2000). A number of possible sources of error have been proposed to explain this negative difference; the most likely being the cumulative uncertainty resulting from taking successive small measurements.

$\text{NO}_z$  species ( $\text{NO}_z = \text{NO}_y - \text{NO}_x$ ) were not determined directly from the continuous measurement of the chemiluminescence  $\text{NO}_x$  analyzers because of the recognized difficulty this equipment has in accurately measuring  $\text{NO}_2$  (US EPA, 2001). These kind of analyzers report  $\text{NO}_x$  as the sum of NO and

NO<sub>2</sub> detected by the instrument. However, the chemiluminescent NO<sub>x</sub> analyzers not only measure NO<sub>2</sub> but also simultaneously respond to PAN and other organic nitrates (Winer et al., 1974; Grosjean and Harrison, 1985; Gerboles et al., 2003). As a result, routine NO<sub>x</sub> measurements overestimate the actual NO<sub>x</sub> and are not suitable for the calculation of indicators (US EPA, 2001).

Because of this operative deficiency of the NO<sub>x</sub> analyzers, several empirical approaches were investigated to derive a surrogate concentration for the NO<sub>z</sub> species. These included: use of direct NO<sub>x</sub> measurements; a gross “corrected” estimation of NO<sub>x</sub> from the sum of measured NO and an assumed measured-surrogate of the sum NO<sub>2</sub> + PAN; the estimation of NO<sub>2</sub> from the difference between the sum of estimated and measured individual contributions to the total measured NO<sub>y</sub>; the empirical calculation of NO<sub>x</sub> from the difference between the measured NO<sub>y</sub> and an estimated fraction of the cumulative NO<sub>y</sub> mass loss by deposition; the estimation of NO<sub>2</sub> from an empirical equation derived from selected ambient measurements of NO, NO<sub>2</sub>, and NO<sub>x</sub>; and the graphical interpolation of an uncertainty correction factor to NO<sub>2</sub> readings from measured NO and NO<sub>x</sub>. Table 2.2 shows a summary of the general characteristics of the investigated approaches. All approaches shown in Table 2.2 were evaluated for consistency by means of several conceptual and graphical comparisons between the measured data and other information available in the literature. The BUWAL empirical equation ( $\text{NO}_2 \cong 0.055[\text{NO}_x] + 55[1 - e^{-\text{NO}_x \cdot 0.01173}]$ ) was selected to calculate the surrogate NO<sub>z</sub>. The criterion for its selection was that this relationship has been used by the European Environmental Agency to investigate the air quality trends of more than 200 cities in 15 countries of the European Union (de Leeuw et al., 2001).

Although the annular denuder methods have been tested in field measurements and have been suggested as appropriate for the determination of H<sub>2</sub>O<sub>2</sub> and HNO<sub>3</sub> levels typical of urban areas (Possanzini et al., 1988; US EPA, 1999), there are several factors that may affect measured concentrations.

All the measured data were statistically analyzed in periods of 1-hr for their mean, standard deviation and range. An interpretation of the resulting average diurnal profiles was performed for each of the monitoring campaigns as an introduction for the subsequent analyses.



**Table 2.2.** Summary of empirical approaches investigated as alternatives to approximate concentration values of NO<sub>x</sub> or NO<sub>2</sub> species useful for the estimation of NO<sub>z</sub> concentrations.

Estimating approach	Equation/Basic procedure	Comments
NO <sub>x</sub> as an upper limit of regular NO <sub>x</sub> measurements <sup>a</sup>	$\text{NO}_x \cong \text{NO} + \text{NO}_2$ <p>where: NO<sub>x</sub> is the concentration measured by the chemiluminescent analyzer</p>	The assumption that contributions of PAN and HNO <sub>3</sub> are minor and that they do not interfere may lead to important uncertainties in the calculation of NO <sub>z</sub> .
NO <sub>x</sub> as the sum of NO, and a surrogate of the sum of NO <sub>2</sub> and PAN	$\text{NO}_x \cong \text{NO} + \text{NO}_2^*$ <p>where: NO is the measured concentration and NO<sub>2</sub>* is the NO<sub>2</sub> reading from the instrument, which is then interpreted as a surrogate of the sum of the concentrations of NO<sub>2</sub> and PAN.</p>	This approximation is based on the fact that PAN is quantitatively detected by the NO <sub>x</sub> analyzer in the NO <sub>2</sub> mode and on the assumption that all of the HNO <sub>3</sub> is lost in the sample tubing during the transport of the air sample to the analyzer and that all the NO <sub>3</sub> <sup>-</sup> is retained on the filter of the air inlet sample to the analyzer.
NO <sub>2</sub> from the difference between measured and/or empirically derived NO <sub>y</sub> species <sup>b</sup>	$\text{NO}_2 \cong \text{NO}_y - \text{NO} - \text{HNO}_3 - \text{PAN}^* - \text{NO}_3^-$ <p>where: NO<sub>y</sub>, NO and HNO<sub>3</sub> are the hourly average measured species; PAN is an estimated derived from a power regression<sup>c</sup> performed to reported concurrent measured PAN and O<sub>3</sub> data from urban and rural sites<sup>d</sup>; and NO<sub>3</sub><sup>-</sup> is an estimated<sup>e</sup> from the reported ratio HNO<sub>3</sub>/NO<sub>3</sub><sup>-</sup> observed in a suburban site of Tennessee<sup>f</sup>.</p>	The assumptions that total NO <sub>y</sub> is given by the sum of only the main nitrogen oxidized species (NO <sub>y</sub> = NO + NO <sub>2</sub> + PAN + HNO <sub>3</sub> + NO <sub>3</sub> <sup>-</sup> ); that PAN is related to the O <sub>3</sub> levels as reported for other regions; and that the ratio HNO <sub>3</sub> /NO <sub>3</sub> <sup>-</sup> is universally constant, have not been tested for the Cincinnati and Mexico City regions.
NO <sub>x</sub> from the hypothetical mass balance of the NO <sub>x</sub> input and the NO <sub>y</sub> deposition losses in the tropospheric regional airshed system from time zero (6:00 hr) to time <i>t</i> <sup>g</sup>	$\text{NO}_x(i) \cong \text{NO}_y(t) + \text{DNO}_y(t)$ <p>where: NO<sub>x</sub>(<i>i</i>) is the hypothetical concentration at time <i>t</i> corresponding to the mass of NO<sub>x</sub> input into the system from time zero to time <i>t</i>. NO<sub>y</sub>(<i>t</i>) is the measured concentration at time <i>t</i>. DNO<sub>y</sub> is the cumulative deposition of NO<sub>y</sub> at time <i>t</i>, and where the deposition term is a function of a time-varying deposition velocity and of the depth of the mixed layer in that region.</p>	This empirical estimation is dependent on the accuracy of the measurement of the mixing layer, and on the assumption that the deposition coefficients required to obtain the deposition velocity are universal and do not change with time of the day.
NO <sub>2</sub> from NO <sub>x</sub> measurements <sup>h</sup>	$\text{NO}_2 \cong 0.055[\text{NO}_x] + 55[1 - e^{-\text{NO}_x \cdot 0.01173}]$ <p>where: NO<sub>x</sub> are the readings reported by the chemiluminescent analyzer.</p>	Empirical equation based on selected monitoring data in Germany, Austria and Switzerland. Used to estimate the air quality trends of a number of cities in the European Union. <sup>i</sup>
NO <sub>2</sub> uncertainty factor from measurements of NO, NO <sub>2</sub> and NO <sub>x</sub> <sup>j</sup>	Graphical interpolation of the NO <sub>2</sub> relative expanded uncertainty using a standard graph of NO <sub>x</sub> versus NO.	The interpolation is difficult to perform. It gives very approximated values particularly at low NO <sub>x</sub> and NO concentrations.

<sup>a</sup> Chameides et al., 1992.

<sup>b</sup> For the Cincinnati case.

<sup>c</sup> PAN = 0.0004[O<sub>3</sub>]<sup>1.9524</sup>; R<sup>2</sup> = 0.97.

<sup>d</sup> Roberts et al., 1995.

<sup>e</sup> NO<sub>3</sub><sup>-</sup> = 0.2489[HNO<sub>3</sub>]; R<sup>2</sup> = 0.878

<sup>f</sup> Williams et al., 1998.

<sup>g</sup> Blanchard et al., 1999.

<sup>h</sup> BUWAL, 1997.

<sup>i</sup> de Leeuw et al., 2001.

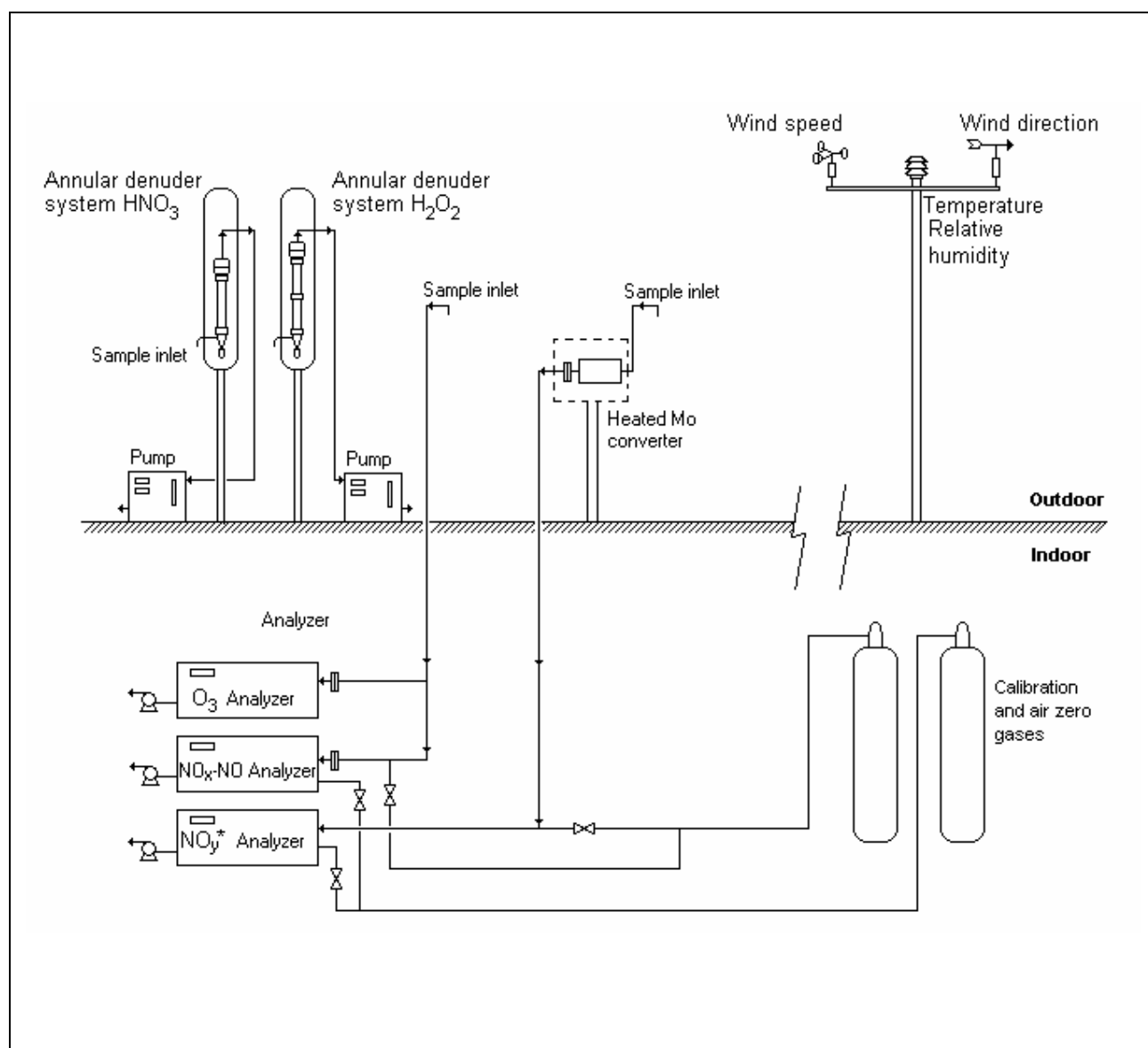
<sup>j</sup> Gerboles et al., 2003.

These factors include: the possibility of uptake and conversion of other nitrogen containing compounds in the  $\text{HNO}_3$  sampling and artifact  $\text{H}_2\text{O}_2$  formation from  $\text{O}_3$  on denuder walls (Possanzini et al., 1988; Sakugawa et al., 1990; Finlayson-Pitts and Pitts, Jr., 2000; Saver et al., 2003).

The monitoring and sampling programs in Cincinnati and Mexico City were supported with tracking down procedures and data validation assessment recommendations taken from the US EPA (US EPA, 1994, 1998). Because in both cases the measurements of  $\text{O}_3$ ,  $\text{NO}$ , and  $\text{NO}_y$  were conducted at official monitoring sites, the calibration and operation of the equipment was done under protocols of the US EPA by the personnel in charge of the monitoring stations (HCDOES, 2003; RAMA, 2004). These protocols include the accuracy evaluation between the field measurements and a quantitative transfer standard.

The adapted chemiluminescent  $\text{NO}_x$  analyzers (or  $\text{NO}_y^*$  analyzers) were calibrated with the same protocol as used for typical  $\text{NO}_x$  analyzers. One standard gas was supplied at the inlet of the converters. The common calibrated audit systems consisted of zero air, NIST-traceable  $\text{NO}$  gas in a cylinder, and an ozone generator. The response of the molybdenum converter to convert  $\text{NO}_y$  species was verified through a response comparison of the  $\text{NO}_y^*$  instrument with the concurrent measurements of  $\text{NO}_x$  from the analyzer that was measuring  $\text{NO}$ . Figure 2.1 shows a general scheme of the setup distribution of the different instruments in the monitoring sites.

An additional validation of the performance of the  $\text{NO}_y^*$  analyzer in Mexico City was done through the evaluation of the statistical correlation between the  $\text{NO}_y$  concentrations and their concurrent  $\text{CO}$  levels. Because both  $\text{CO}$  and  $\text{NO}_y$  have the same origin, a good correlation should be observed between them under indirect evaluation (Sillman et al., 1997). The sampling and analysis of  $\text{HNO}_3$  and  $\text{H}_2\text{O}_2$  had their own quality assurance program. The procedures used were adapted from the original methods themselves and from general recommendations made by the US EPA (1998, 1999). Special emphasis was placed on sampling travel blanks and analytical calibration curves. The correlation between  $\text{O}_3$  and  $(2\text{H}_2\text{O}_2 + \text{NO}_z)$  was used to indirectly evaluate the consistency of the  $\text{NO}_y$  measurements in Cincinnati (Sillman et al., 1997).



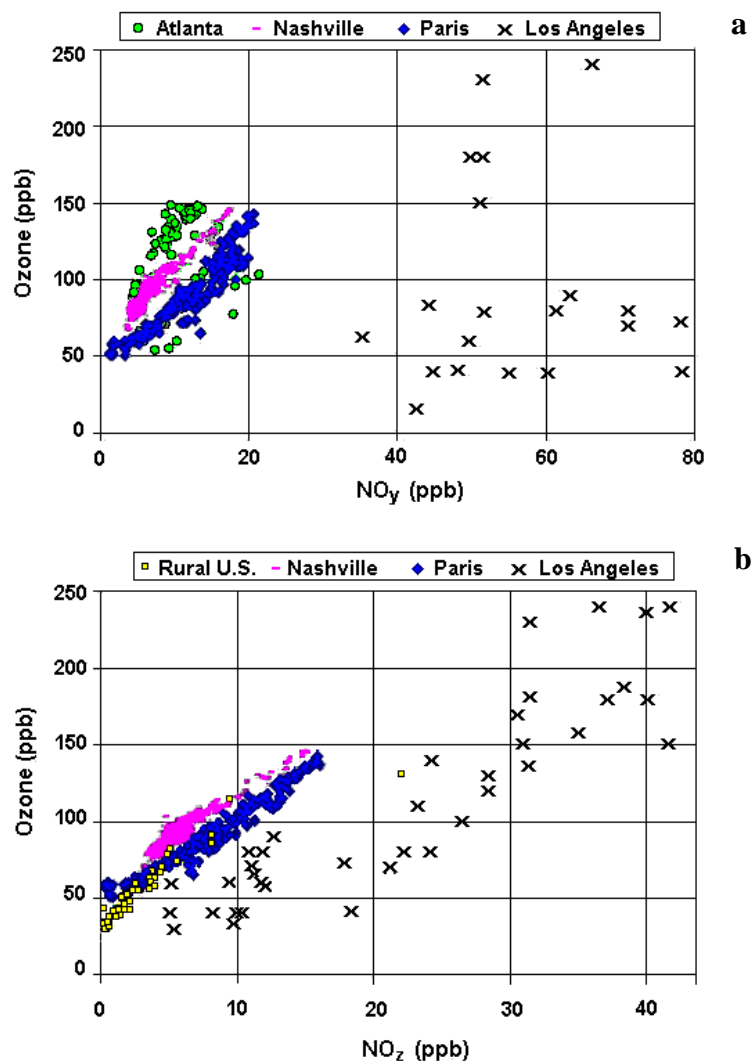
**Figure 2.1.** Schematic diagram of the general setting up of the instruments in the monitoring sites.

## 2.2 Photochemical Indicators Evaluation

Photochemical indicators consist of correlations and calculated ratios of secondary photochemical products. Because of the lack of a standard procedure for using PIM, however, the PIM, the consistency of the measurements must be validated.  $\text{NO}_y$  measurements were evaluated for two purposes: (1) to insure that  $\text{HNO}_3$  was included in the measurement (section 2.1), and (2) to insure that ambient conditions were appropriate. Because the connection between  $\text{NO}_y$  and  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity is due primarily to  $\text{HNO}_3$ , rather than to other components of  $\text{NO}_y$ , it is not acceptable to substitute  $\text{NO}_x$  or the sum of  $\text{NO}_x$  and organic nitrates for  $\text{NO}_y$  in the PIM determinations (Sillman 1995; 2002). In addition, measurements affected by rainfall events (at the site or upwind) cannot be used as  $\text{NO}_y$ -VOC indicators because  $\text{HNO}_3$  has been removed. Other suspicious conditions, such as the presence of high  $\text{NH}_3$  concentrations in the location of the measurement, should be carefully evaluated, because the chemistry of aerosol formation acts as an important removal path for  $\text{HNO}_3$ .

The validity of the measurements performed during the monitoring campaigns was evaluated following the recommendations of Sillman (2002). These evaluations were done by examining scatterplot patterns of  $\text{O}_3$  versus  $\text{NO}_y$ , as well as versus  $\text{NO}_z$  and  $\text{HNO}_3$ , and comparing them to previously measured and/or modeled correlations. According to Sillman, these scatterplots should be imposed on the pattern of research-grade measured and/or model correlations for well identified  $\text{NO}_y$ -sensitive and VOC-sensitive conditions. To be valid, measured values should fall within the broad range of model values. However, if a significant portion of the measurements fall outside of model values, then the measurements for the day in question should not be interpreted as  $\text{NO}_y$ -VOC indicators. In all cases, the evaluations must be done to measurements performed during afternoon hours (13:00 to 17:00 hr).

Figure 2.2 shows the scatterplots of measured correlations between  $\text{O}_3$  and  $\text{NO}_y$  and  $\text{NO}_z$  obtained in several research experiments performed in regions previously identified with either VOC- or  $\text{NO}_x$  sensitive conditions. The overlapping of the measured values on these patterns should be interpreted as a preliminary identification of either  $\text{NO}_x$ - or VOC-sensitive conditions associated to each region.

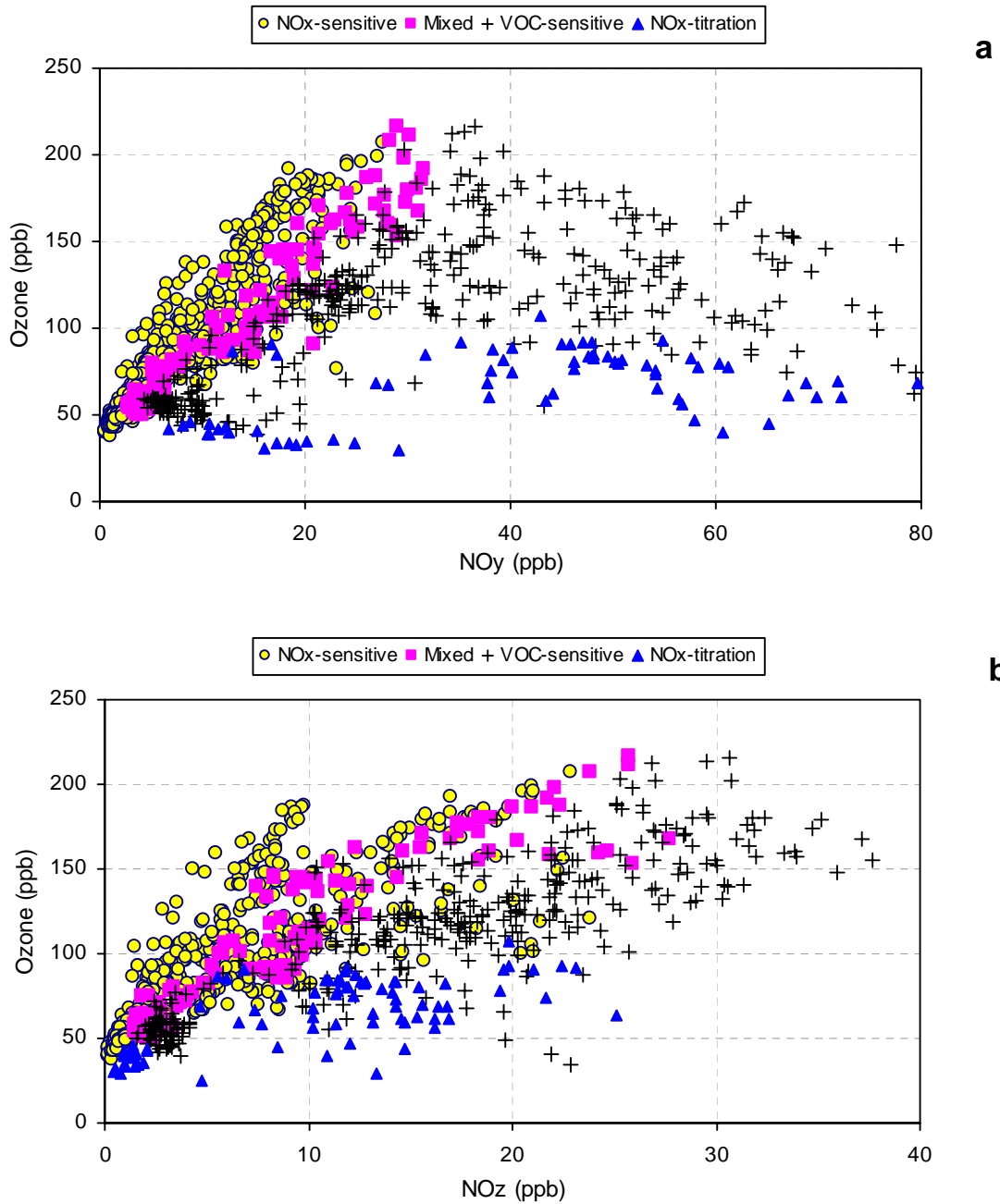


**Figure 2.2.** Measured scatterplot comparison between  $\text{O}_3$  and  $\text{NO}_y$  in (a) Atlanta (circles), Nashville (dashes); Paris (diamonds) and Los Angeles (crosses); (b) measured correlation between  $\text{O}_3$  and  $\text{NO}_z$  in Nashville (dashes), Paris (diamonds), Los Angeles (crosses) and from rural sites in the Eastern U.S. (squares), from measurements reported by Sillman et al. (1997, 1998, 2003) and Trainer et al. (1993). The identified conditions are:  $\text{NO}_x$ -sensitive in Atlanta, GA., mixed sensitivity in Nashville, TN. and Paris, France, and strongly VOC-sensitive in Los Angeles, CA. Figures adapted from Sillman 2002 and 2003.

Figure 2.3 shows the scatterplots of modeled  $O_3$ ,  $NO_y$ , and  $NO_z$  correlations for VOC- and  $NO_x$ -sensitive conditions obtained from 3-D photochemical simulations of several representative regions in the U.S., performed by Sillman 2002; Sillman and He, 2002. These plots show four different pattern conditions:  $NO_x$ -sensitive, VOC-sensitive, mixed, and  $NO_x$ -titration. The interpretation of the first three conditions directly arises from the definitions previously mentioned in this dissertation. The  $NO_x$ -titration condition was included by Sillman, and Sillman and He, to consider that situation in which the locations under study are near large sources of  $NO_x$ . These locations typically have relatively low  $O_3$ . At the same time,  $O_3$  has been affected primarily by the titration reaction ( $O_3 + NO \rightarrow NO_2 + O_2$ ) in the presence of directly emitted NO, rather than by  $O_3$  production mechanisms. This condition should be carefully evaluated. Appendix A presents the modeling-base definition of each of the conditions mentioned above.

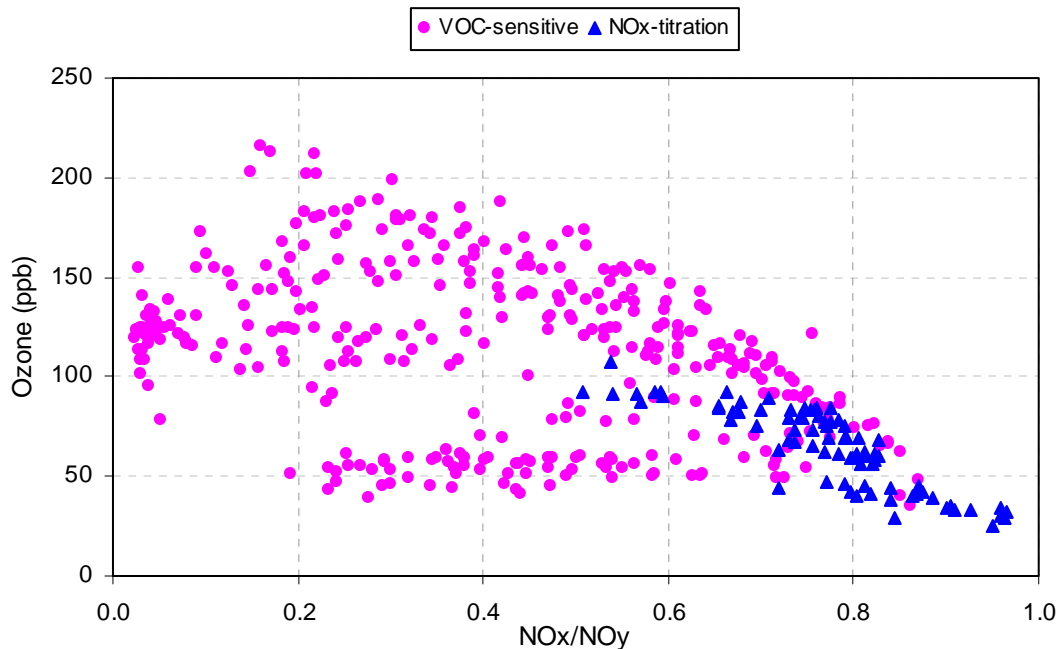
Comparison with model correlations is considered important because the criteria for  $NO_x$ -sensitive and VOC-sensitive indicator ratios were based on models with specific scatterplot patterns. If measurements are not consistent with these model scatterplot correlations, the indicator ratios may not be valid. Sillman (2003) and Sillman and He (2002) have provided data files and plots of modeled values for the previously mentioned correlations that can readily be used for evaluation purposes.

It should be noted that several predicted  $NO_x$ -titration data points overlap other data points for other conditions. Therefore, at one extreme of Figure 2.3(b), there is the possibility that measured data points may fall within this zone of scatterplots patterns and create confusion in the interpretation of the data. To solve this problem, simple procedures that would identify when a value was associated with  $NO_x$ -titration conditions were investigated. The best approach resulted from the combination of photochemical age concept (Trainer et al., 1993) with the scatterplot patterns for each of the VOC-sensitive and  $NO_x$ -titration conditions (Sillman, 2002). As previously mentioned, the  $NO_x$ -titration is related to situations in which a given location is near the influence of relatively large sources of  $NO_x$ . Trainer et al. (1993) have suggested that to discriminate the influence of fresh emissions in the measurement of  $NO_y$  species in rural areas, a figure  $\leq 0.6$  for the ratio  $NO_x/NO_y$  could be indicative of the arriving of chemically aged air masses.



**Figure 2.3.** Predicted scatterplot patterns for (a) O<sub>3</sub> vs. NO<sub>y</sub>, and (b) O<sub>3</sub> vs. NO<sub>z</sub> from 3-D simulations listed in Table 1 in Sillman and He (2002). Each location is classified as NO<sub>x</sub>-sensitive (circles), VOC-sensitive (crosses), mixed or with near-zero sensitivity (squares), and dominated by NO<sub>x</sub>-titration. Scatterplots constructed from data file available from Sillman (2003).

However, in urban areas, this criterion could not be appropriate, because there is always the possibility that local traffic will cause fresh emissions. Sillman (2002) has suggested a limit value of 0.7 for this ratio. The correlation between  $O_3$  and the corresponding ratios  $NO_x/NO_y$  for the VOC-sensitive and  $NO_x$ -titration conditions is shown in Figure 2.4.



**Figure 2.4.** Scatterplot patterns for predicted  $O_3$  vs. the ratio  $NO_x/NO_y$  from 3-D simulations listed in Table 1 in Sillman and He (2002). From data file available from Sillman (2003).

The idea behind this gross evaluation was simple. If the superimposing of the scatterplot correlation of measured  $O_3$  vs. the estimated ratio  $NO_x/NO_y$  over the patterns clustered in Figure 2.4 matched the data points of the predicted  $O_3$  vs.  $NO_x/NO_y$  correlation for the  $NO_x$ -titration condition, or if the measured  $NO_x/NO_y$  ratio was greater than 0.7, then this observation might not be appropriate for the indicator method application.

Two very important assumptions were made in the application of PIM. The first was that, due to the lack of photochemical modeling information for Cincinnati and of accurate simulated results for



Mexico City, the predicted correlation values of Sillman and He (2002) and Sillman (2003) would be valid for inferring the consistency of the indicator evaluations for the study cities.

The second assumption involved the fact that indicators were measured at only one representative site in both cities. Sillman (2002) suggested that measurements of indicators be made over a network broad enough to include most of the locations with elevated  $O_3$  in the study region. If the latter were not possible, however, the selected site should be representative of the ozone air pollution situation in the region. The monitoring sites in Cincinnati (Taft monitoring station) and in Mexico City (Santa Ursula monitoring station) were selected for this very reason. It was also assumed that the single measurement site could be compensated for with the collection of data during a representative monitoring period. In this situation, the sampling would include air masses coming from the dominant sectors of the region and not from just one particular event. There was one attempt to measure  $O_3$ ,  $NO_y$ , and  $HNO_3$  species at a rural site of Mexico City, but several inconsistencies were observed in the measurements. Additional details regarding the experimental work are presented in chapter 4.

Once the consistency of the results was evaluated, the data were interpreted numerically and graphically. The afternoon species ratios of  $O_3/NO_y$  and  $O_3/\text{surrogate } NO_z$  (or  $NO_z^*$ ) and afternoon  $NO_y$  concentrations for both cities were directly compared with the original indicator transitional values shown in Table 1.1. Subsequently, the measured correlations of  $O_3$  versus  $NO_y$  (and of measured  $O_3$  vs.  $NO_z^*$ ) were superimposed over the respective simulated correlations for mixed sensitivity conditions given by Sillman (2003) and Sillman and He (2002). Measured data points close to the mixed values were identified as having mixed sensitivity, while  $O_3/NO_y$  (or  $O_3/NO_z^*$ ) ratios significantly above or below this correlation were interpreted as representing  $NO_x$ -sensitive or VOC-sensitive conditions, respectively.

## **2.3 Complementary Data Analysis Tests**

Several additional approaches were used to validate and/or reinforce the findings of the  $O_3$ - $NO_x$ -VOC sensitivity analysis in Cincinnati and Mexico City. The first approach was the investigation of the weekend/weekday effect on the peak  $O_3$  concentrations. The basic assumption supporting this semi-

empirical evaluation was that, in  $\text{NO}_x$ -sensitive locations (i.e., locations with abundant VOC relative to  $\text{NO}_x$ ), weekend  $\text{O}_3$  peak should decrease significantly compared with the average weekday  $\text{O}_3$  peak, as a consequence of the weekend reductions in  $\text{NO}_x$  emissions (i.e., reduction in diesel trucks and gasoline vehicles traffic) (right lower corner of EKMA diagram in Figure 1.1). In areas where ozone formation is VOC-limited (left upper corner of EKMA diagram in Figure 1.1), weekend  $\text{O}_3$  peak should exhibit a significantly higher value relative to the average weekday  $\text{O}_3$  peak due to the weekend reductions in  $\text{NO}_x$  emissions (Altshuller et al., 1995; Fujita et al., 2002; Heuss et al., 2003; Pun et al., 2003).

A two-tailed  $t$ -test was used to test the statistical significance of the mean differences. The  $t$ -test has been used to interpret weekend/weekday differences in ambient air (US EPA, 1979) and is appropriate for the analysis of small populations (Spiegel, 1961). The means tested were the weekend and weekday peak  $\text{O}_3$  and mean weekend and weekday 6:00 to 9:00 hr (6-9 am) NO concentrations. The mean 6-9 am NO was assumed to be a surrogate of the  $\text{NO}_x$  morning emissions because the difficulty to know the exact uncertainty of the  $\text{NO}_x$  measurements. The test of hypotheses and significance were formulated based on the definition of  $\text{NO}_x$ - and VOC-sensitive conditions, as explained in chapter 1. Table 2.3 shows the main characteristics of the statistical tests. This evaluation was extended to the max 8-hr  $\text{O}_3$  average.

The comparison of weekend/weekday differences in ozone and precursor species with OBM has been explored previously. Blanchard and Fairley (2001) investigated the use of the “extent of reaction” concept, which is based on the smog production ( $sp$ ) algorithm, to strengthen the delineation of the  $\text{NO}_x$ - or VOC-limitation in urban areas. However, the strong dependence of the  $sp$  algorithm calculation on a relatively high number of empirical and practical assumptions seems to be associated to the uncertainty of the results.

Sillman and He (2002) have investigated the  $sp$  algorithm as a tool to delineate the  $\text{NO}_x$ -VOC sensitivity transition, in an approach similar to the PIM graphical concept using 3-D models, and found that, although there was a certain correlation with the  $\text{NO}_x$ -VOC sensitivity predicted with models, its behavior showed a systematic variation with  $\text{O}_3$ , making the identification of the transition range difficult.

**Table 2.3.** Statistical tests used to differentiate the weekend/weekday effect on maximum afternoon O<sub>3</sub>.

Parameter	Hypothesis <sup>a</sup>	Implication
Maximum 1-hr and 8-hr O <sub>3</sub> concentration	$H_0: \mu_{O_3WE} = \mu_{O_3WD}$	The change in WE maximum O <sub>3</sub> is insignificant, likely to be caused by a VOC-sensitive condition provided that the reduction in morning WE-NO <sub>x</sub> emissions is significant
	$H_I: \mu_{O_3WE} < \mu_{O_3WD}$	The decrease in maximum WE maximum O <sub>3</sub> is significant, likely due to a NO <sub>x</sub> -sensitive condition on condition that morning WE-NO <sub>x</sub> emissions is significant
Morning 6-9 a.m. NO average concentration	$H_0: \mu_{6-9am\ NO\ WE} = \mu_{6-9am\ NO\ WD}$	The reduction in morning WE-NO <sub>x</sub> emissions is insignificant probably due to mere chance
	$H_I: \mu_{6-9am\ NO\ WE} < \mu_{6-9am\ NO\ WD}$	The reduction in morning WE-NO <sub>x</sub> is significant likely due to the reduction in WE activities

<sup>a</sup> Two-tailed *t*-test assuming equal variances.

Several empirical evaluations based on the chemistry of tropospheric formation of O<sub>3</sub> were used for assessing the characteristics of the O<sub>3</sub>-formation scenarios in Cincinnati and Mexico City. These evaluations consisted basically on the analysis of the average duration of O<sub>3</sub> accumulation and the average rate of O<sub>3</sub> accumulation following a procedure suggested by Fujita et al. (2002). These evaluations were complemented with the analysis of the average rate of NO depletion. The duration of O<sub>3</sub> accumulation was estimated by calculating the difference between the occurrence of maximum O<sub>3</sub> ( $t_{maxO_3}$ ) and the time in the morning when NO and O<sub>3</sub> crosses over ( $t_{NO=O_3}$ ). The NO-O<sub>3</sub> crossover time represents the end of the morning inhibition period in which fresh NO emissions effectively titrate O<sub>3</sub> and restrain radical formation. The rate of O<sub>3</sub> accumulation (in ppb/hr) is the ratio of the increase in O<sub>3</sub> from  $t_{NO=O_3}$  to  $t_{maxO_3}$  and the duration of O<sub>3</sub> accumulation is the time from the end of the inhibition to the maximum O<sub>3</sub> concentration (Rate =  $\{[O_3]_{max} - [O_3(t_{NO=O_3})]\} / \{t_{O_3max} - t_{NO=O_3}\}$ ). The rate of NO depletion was assumed to be equivalent to the ratio of the difference in NO concentrations from the time of its morning peak to the time when the O<sub>3</sub> peak occurs (Rate =  $\{[NO]_{max} - [NO](t_{NO=O_3\ max})\} / \{t_{NO\ max} - t_{NO=O_3\ max}\}$ ).

An empirical interpretation of the severity of the event suggested by Milford et al. (1994) and Sillman (1999) was used for the evaluation of the consistency of the  $\text{NO}_x$ -VOC sensitivity suggested by photochemical indicators results and by theoretical sensitivity determined by the ratio of initial VOC/ $\text{NO}_x$  emissions. These authors have suggested that environmental conditions prone to have higher ozone precursors can be associated with either higher emission densities (i.e., larger urban areas) or with more restricted meteorological conditions (i.e., light winds and low daytime vertical mixing) or with a combination of both. Sillman has proposed an explanation for the severity of event effect by doing an analogy with the theoretical concept of Kleinman (1991, 1994) that indicates that the split between the  $\text{NO}_x$ -sensitive and  *$\text{NO}_x$ -saturated* (or VOC-sensitive) regimes is related to the relative supply of  $\text{NO}_x$  (from emissions) in comparison with the supply of radicals generated by sunlight. Freshly emitted plumes of polluted air have an initial  $\text{NO}_x$  supply that greatly exceeds the supply of radicals. As the plume ages the total amount of radicals created during the process of photochemical evolution becomes equal with and eventually surpasses the initial  $\text{NO}_x$ -source, causing a switch from  *$\text{NO}_x$ -saturated* to  $\text{NO}_x$ -sensitive conditions. Under the analogy of Sillman, if the source of  $\text{NO}_x$  source per unit volume in an airbasin is high (after considering the effect of dilution through daytime vertical mixing), it will take more time for the accumulated source of radicals (from the VOCs) to become equal with and surpass the  $\text{NO}_x$  source. As a result, a  *$\text{NO}_x$ -saturated* (VOC-sensitive) condition will persist for a longer time. But, if the  $\text{NO}_x$  source is low, the accumulated radical source will exceed the  $\text{NO}_x$  source after a short period of photochemical aging and the system will shift to  $\text{NO}_x$ -sensitive conditions.

## 2.4 Intercomparison Tests

Sillman and He (2002) have suggested that their transition values for indicator ratios may not be valid for every region in the world. On the one hand, they are appropriate for the Cincinnati airshed region, because they were obtained from photochemical simulations in regions of the U.S. On the other, those transition values might not be adequate for airshed regions outside the geographic region of North America such as Mexico City.

One attempt was made to evaluate the consistency of the photochemical indicator method (PIM) to the Mexico City observations using a state-of-the-art photochemical model (the Multiscale Climatic and Chemistry Model (MCCM)). However, the lack of sufficient input data to the model for the measurement period limited its application.

As an alternative, a semi-quantitative evaluation of the consistency of the  $\text{NO}_x$ -VOC sensitivity diagnosed with the photochemical indicator approach in Mexico City was done through the intercomparison of the measured indicator results with the findings of a preliminary effort carried out by Sosa et al. (2000) to identify the local  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity, and with the results of a simple box model results by Gaffney et al., (1999) to simulate some secondary species indicators. The measurements used by Sosa et al. and Gaffney et al. were obtained from a research-grade campaign carried out in Mexico City during the winter of 1997 (Doran et al., 1998; Edgerton et al., 1999). In addition, the measured species indicators and indicators ratios from the present research monitoring campaign were intercompared with the results of a modeling experiment with the MCCM for a historical ozone event. A direct comparison between all the available information could not be performed because the data had different averaging times, different simulating days, and/or different approaches of estimation.

Sosa et al. (2000) attempted to apply the PIM of Sillman based on a compilation of measurements taken at three different locations in Mexico City over an 18-day period. The monitoring locations were: La Merced (close to the downtown), Azcapotzalco (~ 13 km northeast of La Merced), and IMP (~ 8 km north of La Merced). They studied the correlations  $\text{O}_3/\text{NO}_y^*$ ,  $\text{O}_3/\text{NO}_z^*$ , and  $\text{O}_3/\text{HNO}_3^*$  (where the asterisks indicates surrogates), but they could not arrive at conclusive results.

In addition to the above evaluation, Sosa et al. (2000) carried out a number of modeling experiments with the CIT 3-D photochemical model to study  $\text{O}_3$  formation in MCMA and to diagnose the  $\text{NO}_x$ -VOC sensitivity. Although they could produce in a relatively good agreement the  $\text{O}_3$  daily pattern with the observed  $\text{O}_3$  measurements of one selected day, and found that all of the urbanized sites were roughly on the crest between  $\text{NO}_x$ -sensitive and VOC-sensitive regions, they could not confirm the consistency of these results. They arrived at their final conclusions after performing an artificial increment

(a factor of three) to the VOC emissions in the base emission inventory. They also reported that the CIT model had problems reproducing the vertical meteorology of the region. The available simulation data from the CIT study was limited to the Merced site (Sosa et al., 2000).

The modeling experiment with the MCCM was a simulation of the spatial distribution of indicator species and indicator ratios for one historical event (March 3, 1997 from 14:00-15:00) using a previously integrated input database. It covered the Mexico City's entire region, including parts of two neighboring states. The emissions for the simulation were taken from an ozone precursors official inventory for the year 1995 (GDF, 1995). The results of the modeling experiment were obtained after a number of preliminary evaluations and discussions of the most representative modeling scenario. A graphical interpolation software package (Surfer 7) was used to generate the approximate species concentrations for the different locations of interest. García (2002) implemented the MCCM to study air quality and meteorological impact for a project to restore an artificial lake in northeast Mexico City and to study the reduction of  $\text{NO}_x$  emissions from a power plant in the region. He found that the MCCM could better reproduce ozone concentrations with a  $\sim 40\%$  reduction in the  $\text{NO}_x$  emissions of the official base inventory. Jazcilevich et al. (2003) used the same input database and MCCM to investigate the influence of wind patterns on the 3-D behavior of air pollutants in central Mexico. They concluded that MCCM was able to reproduce more accurately the meteorology of the region than other models, and, as a consequence, to better reproduce the dispersion of the low reactive air pollutants. These two studies did not investigate further the  $\text{NO}_x$ -VOC sensitivity of the region. Appendix B presents a summary of the main components of the MCCM.

Gaffney et al. (1999) used a simple box model approximation to predict the "average" concentrations of PAN,  $\text{NO}_y$  (as  $\text{NO}_y\text{-NO}$ ), and  $\text{O}_3$  for Mexico City on a typical smog-polluted day based on a set of data collected in March 1997. The results of the simple box modeling were assumed to apply to the entire Mexico City air basin. Because no attempt was done by Gaffney et al. to define a  $\text{NO}_x$ -VOC sensitivity for Mexico City, the data were worked to obtain several different indicator species useful required for the intercomparison.

Additional details on the preparation of the CIT and MCCM 3-D models, are given elsewhere (Grell et al., 1994; García, 2002; Sosa et al., 2002; Jazcilevich et al., 2003). Gaffney et al. (1999) did not provide more details on the particular characteristics of their box model. Table 2.4 shows a summary of the type of data available to perform the intercomparison.

**Table 2.4.** Measured and simulated data available and characteristics of the information for the intercomparison of measured indicator results in Mexico City.

Approach	Species	Averaging time and site of measurement/simulation
Measurement approach <sup>a</sup> (March 2-19, 1997)	O <sub>3</sub>	1-hr average at La Merced and Azcapotzalco sites <sup>b</sup>
	NO <sub>x</sub>	1-hr average at La Merced and Azcapotzalco sites <sup>b</sup>
	O <sub>3</sub>	6-hr average at La Merced and Azcapotzalco sites
	NO <sub>x</sub>	6-hr average at La Merced and Azcapotzalco sites
	HNO <sub>3</sub> + NO <sub>3</sub> <sup>-</sup>	6-hr average at La Merced during IMADA-AVER <sup>c</sup>
	PANs	30-min and 6-hr average at IMP during IMADA-AVER <sup>d</sup>
	O <sub>3</sub> /NO <sub>y</sub> ratio	6-hr average at La Merced and Azcapotzalco sites
CIT simulation <sup>a</sup> (Sunday, March 2, 1997)	O <sub>3</sub>	Surface hourly modeled at La Merced position
	NO, NO <sub>2</sub>	Surface hourly modeled at La Merced position
	PAN	Surface hourly modeled at IMP position
	HNO <sub>3</sub>	Surface 6-hr average modeled at La Merced position
MCCM simulation <sup>e</sup> (Monday, March 3, 1997)	O <sub>3</sub>	Surface 1-hr (14:00-15:00 hr) covering the MCMA region (~ 80 x 75 km in 3600 grid data points)
	NO, NO <sub>2</sub>	
	PAN	
	HNO <sub>3</sub>	
	NO <sub>3</sub> <sup>-</sup>	
	H <sub>2</sub> O <sub>2</sub>	
	HCHO	
Simple Box Model <sup>d</sup> (Winter, 1997)	O <sub>3</sub>	Average for afternoon hours for the whole Mexico City region
	NO <sub>y</sub> (as NO <sub>y</sub> – NO)	
	PAN	

<sup>a</sup> Sosa et al., (2000).

<sup>b</sup> SMA historic data files. (Red Automatica de Monitoreo Atmosferico, SIMA, <http://sma.df.gob.mx/simat/>)

<sup>c</sup> Edgerton et al., (1999) and Chow et al., (2003).

<sup>d</sup> Gaffney et al., (1999).

<sup>e</sup> Obtained from García and Jazcilevich (2004).

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## **Chapter 3**

### **CHARACTERIZATION OF THE METROPOLITAN AREAS**

#### **3.1 Greater Cincinnati**

The Greater Cincinnati area is located in the American Midwest (39.8° North, 84.3° West) at the southwestern tip of Ohio. It is comprised of the urban and suburban settlements that have extended around the city of Cincinnati and cover parts of 12 counties of Ohio, Kentucky, and Indiana. The City of Cincinnati is a part of Hamilton County, OH. The city is located on the bank of the Ohio River and extends over two ranges of hills. The hill elevation is about 120 m above the river valley level. Although the valley is not very steep, it influences meteorological conditions and pollutant dispersion in the area (Martuzevicius et al, 2004). The climate of the region shows a wide range of temperatures through the year. Summers are warm and humid, while winters are cold and dry. The region is subject to frequent changes due to the passage of cyclonic storms in winter and spring and thunderstorms during summer. The average temperature is ~12°C and prevalent wind direction is from south-southwest (NCDC, 2003). As of the 2000 census, Cincinnati had a total population of around 332,000. However, the estimated population in the Greater Cincinnati area is nearly 2 million residents (US Census Bureau, 2004).

Greater Cincinnati is crisscrossed by a network of highways that leads to important traffic flow throughout the day, as well as during the peak rush hours. A recent study found that during the year 2000, an estimated 124,000 vehicles passed through a 17-mile stretch of I-75 (Martuzevicius et al, 2004). In addition to very intensive traffic, Greater Cincinnati hosts almost 1000 major and small industrial facilities that emit a number of air pollutants into its airshed. According to the US EPA (2004), the estimated anthropogenic emissions of VOC and NO<sub>x</sub> for the years 2000-1 were 121,921 and 239,277 tons per year, respectively. These values include the emissions of five counties in Ohio, six counties in Kentucky, and 2 counties in Indiana. Table 3.1 shows a composite of the emission inventory for Greater Cincinnati.

**Table 3.1.** Cincinnati-Hamilton, OH-KY-IN Emissions Inventory (year 2000).

Source Category	Annual (tons per year)				
	VOC	NO <sub>x</sub>	CO	SO <sub>2</sub>	PM <sub>10</sub>
Fuel Combustion Electricity Utilities	438	119,115	3,386	260,990	3,696
Fuel Combustion Industrial	308	17,832	3,353	55,447	1,530
Fuel Combustion Other sources	2,657	6,089	14,339	5,486	1,941
Chemical & Allied products	2,192	188	0	2,051	306
Metals processing	1,151	1,737	22,073	3,176	2,740
Petroleum & Related Products	490	64	105	170	252
Other Industrial Processes	1,850	1,469	1,319	372	458
Solvent Utilization	41,769	0	0	5	150
Storage & Transport (Petroleum products)	8,487	0	84	0	411
Waste Disposal & Recycling	3,677	820	14,658	234	2,501
Highway Vehicles	41,510	63,861	408,338	2,482	2,405
Off-Highway	17,327	28,089	185,820	4,051	2,322
Miscellaneous	65	11	408	0	73,605
Natural Emissions (Biogenic)	39,151 <sup>a</sup>	0	0	0	0
Total	161,072	239,277	653,943	334,459	92,173
Percent of total	10.87	16.16	44.16	22.58	6.23

<sup>a</sup> Estimated as proportional to the total annual emissions of biogenic hydrocarbon emission inventory and total land area for the US EPA Region V (Lamb et al., 1993), for the 13 counties of Ohio, Kentucky and Indiana considered by US EPA for the emission inventory (US EPA, 2004).

Solvent utilization and highway vehicles categories account for 52% of the total VOC, while biogenic hydrocarbon emissions account for 24%. Power plants and highway vehicles account for ~ 77% of the total NO<sub>x</sub> emissions. The estimated VOC/NO<sub>x</sub> ratio for the Greater Cincinnati area is 0.67, expressed as total annual emissions. The tons VOC/tons NO<sub>x</sub> ratio must be converted to ppmC VOC/ppm NO<sub>x</sub> in order to compare the present approximation to previous VOC/NO<sub>x</sub> ratio estimates for the area. Using the method suggested by Baughes (1991) and the average composition of ambient hydrocarbons in 39 U.S. cities (Seinfeld, 1989), the estimated VOC/NO<sub>x</sub> ratio is ~ 1.9 ppmC/ppm.

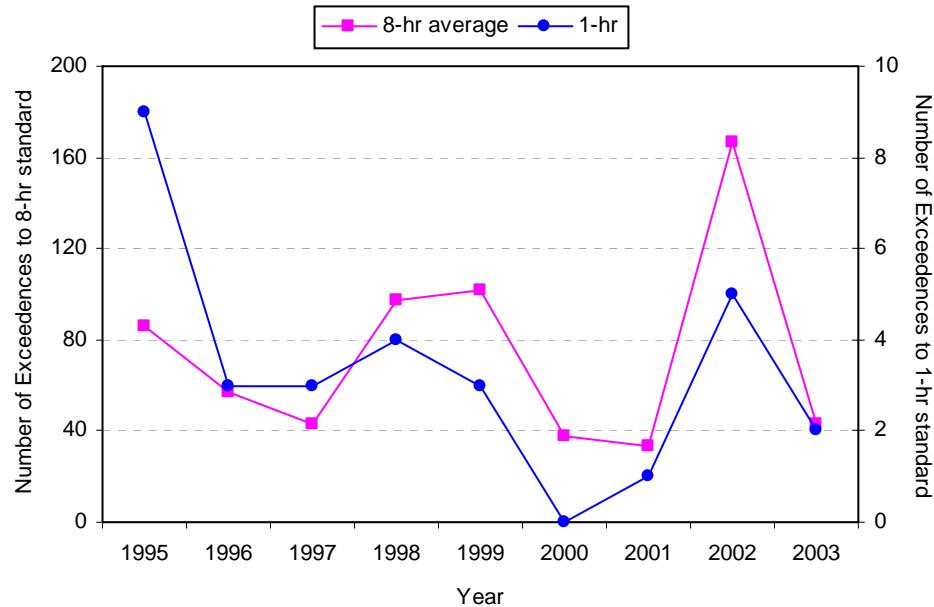
This value contrasts with a previously reported VOC/NO<sub>x</sub> ratios for Cincinnati from 6:00-9:00 hr Non Methane Hydrocarbons (NMHC) and NO<sub>x</sub> monitoring readings of 3.8 in 1976 (OKI, 1979) and of 9.1 in 1984 (Baugues, 1986), and of 6.4 from 1988 emissions inventories (Carter, 1994). The measured values should be taken as indicative because measurements of NO<sub>x</sub> were carried out with chemiluminescence analyzers and the NMHC with flame ionization analyzers. The NMHC analyzers were not explicitly for detecting aromatic compounds. However, the downward trend is in agreement with the trend observed in these ratios in several Northeast U.S. cities from 1986 to 1991 (Wolff, 1993). This trend can be attributed to the more reductions in VOC relative to NO<sub>x</sub> that have occurred in the past 15 years. Theoretically, the estimated VOC/NO<sub>x</sub> ratio of 1.9 for Greater Cincinnati implies that VOC controls should be more effective than NO<sub>x</sub> reductions. However, the Ohio Environmental Council (2000) has stated that, based on photochemical modeling analysis, reduction of NO<sub>x</sub> emissions from power plants in the Ohio Valley would prevent excessive formation of O<sub>3</sub> in the airshed.

O<sub>3</sub> formation in Greater Cincinnati is well correlated with ambient temperature and humidity (Jung, 2000). As a result, the smog season in this airshed is typically from April through October. However, there is evidence that the differences in anthropogenic weekday activities are correlated with the local O<sub>3</sub> formation. A statistical analysis of the frequency of days exceeding an assumed O<sub>3</sub> 1-hr threshold value of 125 ppb during 1992-1999 showed that the lowest frequencies occur early in the week, with a gradual increase through the week towards the highest frequency on Sunday (Jung, 2000). A recent study of the photochemical system in this airshed showed that O<sub>3</sub> concentrations at all sites within the Greater Cincinnati domain are strongly linked dynamically (Kanda, 1997). This means that high O<sub>3</sub> concentrations measured at one site of the region lead to high O<sub>3</sub> concentrations at other sites and vice versa, indicating that O<sub>3</sub> is formed in a well-mixed large-scale air mass. Kanda has suggested that control of some VOCs at one point in the domain could reduce O<sub>3</sub> concentrations over the entire airshed.

Although Greater Cincinnati has maintained the 1-hour O<sub>3</sub> standard since 1996 and had obtained the designation of an attainment area, in June of 2003, the US EPA demoted the area to “nonattainment” status. This was a result of the Ohio EPA’s failure to adopt the federal Reasonably



Available Control Technology (RACT) rules into the Ohio State Implementation Plan (SIP) (HCDOES, 2004a). In addition, preliminary analysis of O<sub>3</sub> monitoring data has shown that southwestern Ohio will be in nonattainment with the new 8-hour O<sub>3</sub> standard. Figure 3.1 shows the trends in 1-hr and 8-hr average O<sub>3</sub> exceedences to their respective US EPA ozone air quality standard in the period 1983-2003.



**Figure 3.1.** Trends in number of exceedences to the 1-hr and 8-hr average US EPA ozone air quality standards in Greater Cincinnati metropolitan area from 1995 to 2003. (US EPA 1-hr O<sub>3</sub> standard: 0.12 ppm max no to be exceeded more than three times over the three most recent years; 8-hr average O<sub>3</sub> standard: 0.08 ppmv max based on highest three-year average of the fourth highest 8-hour concentration). (Source: HCDOESa, 2004)

In addition, because Ohio is a part of the 22 states in the US within the US EPA's NO<sub>x</sub> State Implementation Plan Call, it has the obligation to reduce ~ 33% of NO<sub>x</sub> emissions to support the control of transported ozone in the region (68 FR No. 150 August 5, 2003). The most important O<sub>3</sub> control strategy has been the reduction of NO<sub>x</sub> in electricity utilities. Other strategies that have been applied in Greater Cincinnati include reductions in VOC emissions from stationary sources and combined NO<sub>x</sub> and VOC reductions from transportation activities. According to the Hamilton County Department of Environmental Services in Cincinnati (HCDOES, 2004b), from 1996 to 2002 there was a reduction of ~ 13% in VOC

emissions, as a result of stricter inspection and maintenance programs, vapor recovery systems on fuel pumps, and cleaner fuels.

According to a 1998 investigation into health effects, residents in the Greater Cincinnati area were exposed to higher levels of smog more often than residents in Boston and New York and had higher rates of smog-related respiratory hospital admissions, including asthma attacks and emphysema (The Ohio Environmental Council, 2000).

### **3.2 Mexico City Metropolitan Area**

The Mexico City Metropolitan Area (MCMA) lies at a high altitude (2240 m) on an inland flat basin ( $\sim 5,000 \text{ km}^2$ ) at a subtropical latitude (Lat.  $19.5^\circ\text{N}$ ). Mountains to the east, south, and west, reaching to more than 2500 m above the basin floor, inhibit dispersion of air pollution and contribute to frequent wintertime thermal inversions that further trap pollutants near the surface (Bravo and Torres, 2002). Prevalent winds are from the NE and NW sector. Low-level regional winds during the cool season are usually not strong enough to provide ventilation of the air pollution in the basin. At other times, weak regional winds give way to cool, thermally driven down-slope winds that during the night reinforce the centripetal circulation, thus contributing to the lateral confinement of air pollutants (Jauregui, 2002).

The climate of the basin is wet-dry tropical in character, tempered by altitude, with a mean annual temperature of  $16^\circ\text{C}$ . During the dry season (November to April), clear skies and anti-cyclonic weather prevail with little rain. From May to October, the dry westerlies give way to moist trade winds that bring convective rain showers (Jauregui, 2002). The high elevation and latitude combine to provide plentiful sunshine, especially during the late winter, which in comparison to more northern latitudes, is enhanced by the UV radiation that drives atmospheric photochemistry to produce secondary pollutants such as  $\text{O}_3$ .

The MCMA consists of the 16 delegations of the Federal District (DF) and 17 municipalities of the State of Mexico, which border the DF. The population in the MCMA has grown rapidly, from 3 million in 1950 to nearly 20 million in 2000. The urbanized area of the region has reached an extension of

almost 1500 km<sup>2</sup>. More than 40 million L of fuel are consumed per day, producing thousands of tons of pollutants. Table 3.2 shows the emission inventory for the MCMA that has been developed for the year 2000.

**Table 3.2.** Mexico City Metropolitan Area Emissions Inventory (year 2000).<sup>a,b</sup>

Source Category	Annual (tons per year)				
	VOC	NO <sub>x</sub>	CO	SO <sub>2</sub>	PM <sub>10</sub>
Fuel Combustion Electricity Utilities	146	11,626	1,816	16	202
Fuel Combustion Industrial	7,876	6,155	3,594	5,130	1,071
Fuel Combustion Other sources	1,688	8,079	3,385	23	215
Chemical & Allied products	1,688	2,311	3,380	2,332	394
Metals processing	7,453	2,311	3,380	2,332	394
Petroleum & Related Products	5,159	2,563	2,371	1,588	887
Other Industrial Processes	1,515	4,586	964	1,237	447
Solvent Utilization	129,565	0	0	0	0
Storage & Transport (Petroleum products)	33,241	0	0	0	0
Waste Disposal & Recycling	0	0	0	0	0
Highway Vehicles & Off-Highway	194,517	157,239	2,018,788	4,348	5,287
Miscellaneous	33,170	33	1,127	7	1,838
Natural Emissions (Biogenic)	15,425	859	0	0	0
Total	429,755	193,451	2,035,425	14,681	10,341
Percent of total	16.05	7.2	75.8	0.55	0.4

<sup>a</sup> Includes 16 Delegations from the Federal District and 17 municipalities from the State of Mexico.

<sup>b</sup> Source: SMA (2003)

Fuel combustion in motor vehicles accounts for 81% of the total NO<sub>x</sub>. Power plants contribute only with a 6% of the total. Most of the electricity consumed in MCMA is provided by a national interconnected network distribution from a number of power plants located across the country. Vehicles account for 45% of the total VOC, while solvent utilization accounts for 30% of VOC emissions. Leaks

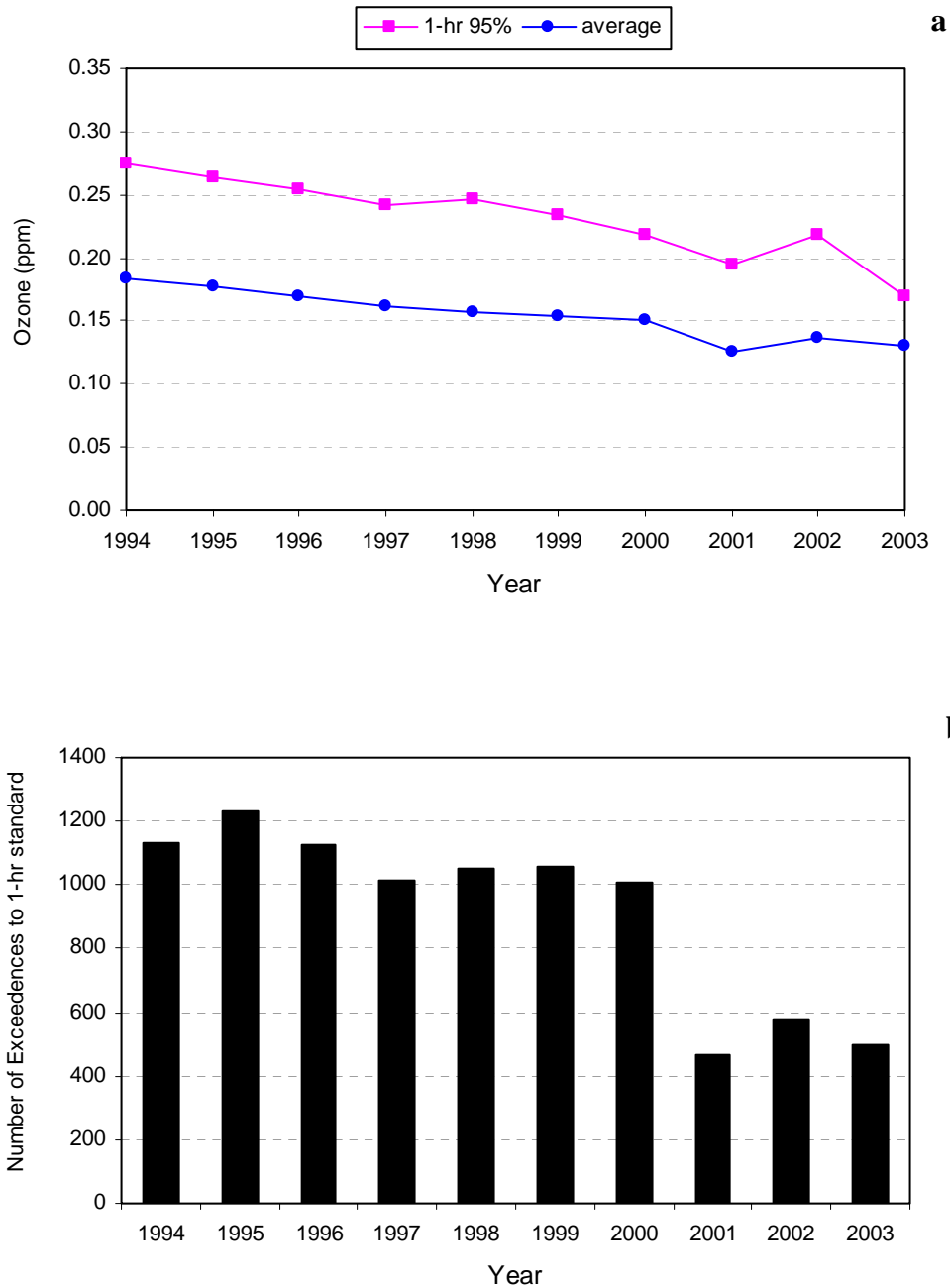
from storage and distribution of Liquefied Petroleum Gas (LPG) account for 7.5% of the total VOC. The estimated VOC/NO<sub>x</sub> ratio from the emission inventory is ~ 2.2. However, comparisons between measured VOC/NO<sub>x</sub> ratio (as ppbC/ppm) and the equivalent ratio derived from emission inventories indicate a strong incongruence. Molina and Molina (2004) reported that, while the VOC/NO<sub>x</sub> ratio derived from past inventories was ~ 3:1 (ppbC/ppm), the same estimation from measurements in representative locations were 15:1 or higher for the same period (1992-1997). These very high VOC/NO<sub>x</sub> ratios have been extremely influential in past studies, which have concluded that ozone is likely to be NO<sub>x</sub>-sensitive (MARI, 1994; Sosa et al., 2002).

The MCMA's large population, 35 000 industries, 3.5 million vehicles, complex topography, and meteorology cause high air pollution levels. Peak concentrations of O<sub>3</sub> are typically reached during the afternoon in the downwind southwest sector of the basin. One important effect of the regional meteorology of the Mexico Basin is the trapping of very high levels of aged O<sub>3</sub> aloft in the nocturnal mixing layer (Bravo and Torres, 2002). This aged O<sub>3</sub> may contribute effectively to the formation of new diurnal O<sub>3</sub>.

The smog season in MCMA is typically between February and March, when the incoming UV radiation is larger and the sky is cleaner. During the wet summer months (June to September), clouds inhibit photochemistry and high O<sub>3</sub> episodes are less frequent.

Although significant improvement in ambient air quality in the MCMA has been reached in the past 12 years, levels remain high. In 2003, O<sub>3</sub> levels exceeded the local air quality standard (0.11 ppm 1-hr max. any day in the year) on 75% of days in the year. Figure 3.2 shows the trends in O<sub>3</sub> concentrations of the 1-hr max 95<sup>th</sup> percentile and 1-hr max annual average from data collected at five representative sites in the MCMA.

MCMA health studies indicate that most of the inhabitants in this area have been affected to some degree by the high levels of O<sub>3</sub> and high dosage of exposure (Calderón-Garcidueñas et al., 1992). Borja-Aburto et al. (1997) have demonstrated a statistically significant association between O<sub>3</sub> levels and cardiovascular mortality in inhabitants of the MCMA.



**Figure 3.2.** Trends in O<sub>3</sub> concentrations for (a) 1-hr max 95<sup>th</sup> percentile and 1-hr max annual average, and (b) number of exceedences to the 1-hr Mexican Ozone Air Quality Standard in MCMA from 1994 to 2003. (Mexican O<sub>3</sub> 1-hr standard: 0.11 ppm max no to be exceeded any time in a year). (Source: SINAICA, 2004).

A new O<sub>3</sub> 8-hr average standard (0.08 ppb) was approved in 2002 by the Mexican authorities. A preliminary analysis of the data indicates that since 1994, every year has not been below or not even close to this standard (SINAICA, 2004).

Historically, air quality management in the MCMA has focused on the whole local air pollution problem and not on one specific air pollutant. Most strategies from 1990-1995 were based on technological modernization and fuel improvement that had been demonstrated as useful in other countries (Molina et al., 2004). The most significant reductions in air pollution in that period, including ambient O<sub>3</sub> levels, can be attributed to the introduction of catalytic converters and the improvement in fuel quality.

In 1996, a five-year control program was initiated based on linear programming involving political and economical factors, as well as results from photochemical modeling efforts (Streit and Guzman, 1996). It is not clear if the concept of O<sub>3</sub>-NO<sub>x</sub>-VOC sensitivity was taken into account in this air quality management program. However, the authorities enforced a strong reduction in VOC emissions. For the period 2001-2010, a new control program was elaborated considering recommendations of a research group coordinated by Dr. Mario Molina at the Massachusetts Institute of Technology and modeling works done by the Fraunhofer Institute (CAM, 2002). The aim of this new plan was to reduce ambient O<sub>3</sub>, PM<sub>10</sub>, and PM<sub>2.5</sub> levels. Control strategies were based on NO<sub>x</sub> reduction, although VOC reductions were also included. The MIT group concluded that, even though there is evidence from modeling and chamber experiments that O<sub>3</sub> formation in MCMA is likely NO<sub>x</sub> sensitive, more research is needed to clarify this situation.

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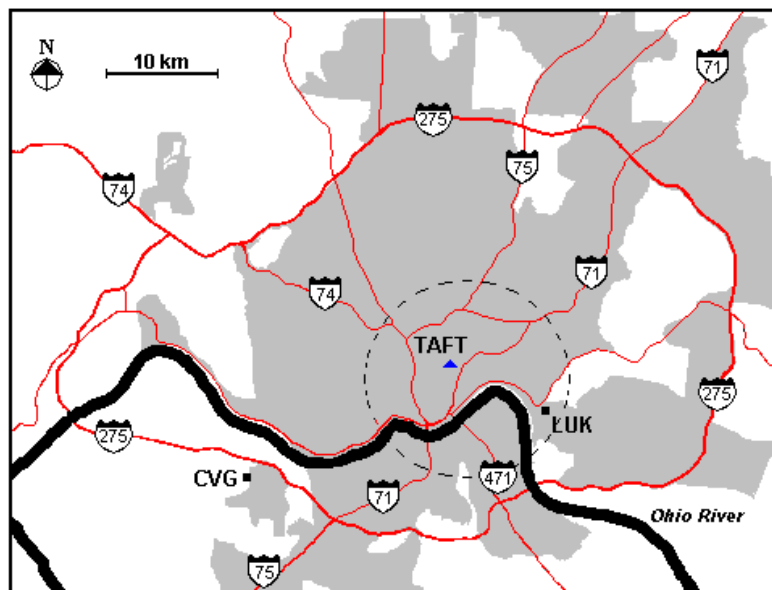
## Chapter 4

### RESULTS AND ANALYSIS

#### 4.1 Greater Cincinnati

##### 4.1.1 Indicator Species Measurements

Continuous measurements of  $O_3$ ,  $NO_x$ , and  $NO_y$  were made from September 3-30, 2003 at the Taft monitoring station of the Hamilton County Department of Environmental Services (HCDOES) in the city of Cincinnati, Ohio ( $39.06^\circ N$ ,  $84.3^\circ W$ , 230 m.a.s.l.), ~ 5 km north of downtown (see Figure 4.1).



**Figure 4.1.** Map showing the Greater Cincinnati area and the location of the Taft Monitoring Station. The dashed circle encloses the approximate extent of the city while the remainder of the gray area outside of the circle shows the extent of the suburban area. LUK represents the Lunken Municipal Airport and CVG is the Cincinnati/Northern Kentucky International Airport.

The hourly average concentrations of  $O_3$ ,  $NO_x$ , and  $NO_y$  were processed by HCDOES protocols. All species were sampled at ~ 6 m above ground level and all instruments were kept in a temperature-controlled room.  $O_3$  measurements were made with a Dasibi Ozone Monitor model 1008-RS equipped with an internal  $O_3$  generation system.  $NO_x$  ( $NO$  and  $NO_2$ ) measurements were made with a Thermo

Environmental Instrument model 42C.  $\text{NO}_y$  measurements were taken with a chemiluminescent method, using a partially modified Columbia Scientific Instruments (CSI) model 1600  $\text{NO}_x$  analyzer with an external heated converter mounted to the gas sample inlet. The external heated converter consisted of a molybdenum (Mo) converter assembly removed from a Monitor Labs 8840  $\text{NO}_x$  analyzer. Energy to heat in the Mo converter was supplied by a cartridge heater. A type K thermocouple embedded in the assembly kept the temperature of the converter at approximately  $365 \pm 5^\circ\text{C}$ . In order to insure that the measured total reactive nitrogen ( $\text{NO}_y$ ) included gas-phase  $\text{HNO}_3$  and aerosol nitrate, the inlet Teflon line to the external converter was cut as short as possible ( $\sim 30$  cm) to minimize the loss of nitric acid on the inlet. In addition,  $\text{NO}_y$  sampling was conducted without a particulate filter.  $\text{NO}_x$  and  $\text{NO}_y$  instrument calibrations were accomplished by the addition of a NBS standard reference material gas for  $\text{NO}_2$  via a Sabio unit calibration system. Conversion efficiency of the Mo converter was calculated to be  $> 98\%$ . Additional meteorological information was obtained from the Cincinnati Municipal Airport Lunke (NCDC, 2003).

Integrated 1-hour average samples (15:00–16:00 hr) of gas-phase  $\text{H}_2\text{O}_2$  and  $\text{HNO}_3$  were collected simultaneously on several weekdays between September 8-17 using two annular denuder systems (ADSs). The ADSs were located on the roof of the HCDOES laboratory; about  $\sim 20$  m away from the inlet sampling of the monitoring equipment. After sampling, the denuders were sealed and kept in a dark enclosure immersed with blue-ice packs. The extractions were performed no later than 30 min after the sampling. The analysis of the samples was performed in the Civil and Engineering Department of the University of Cincinnati.

$\text{H}_2\text{O}_2$  was collected in two Ti(IV)- $\text{H}_2\text{SO}_4$  coated URG annular denuders (3 channels x 30 mm O.D. x 150 mm length glass tubes) in series at a flow rate of 10 L/min (Possanzini et al., 1988). The leachates were then analyzed with a Hewlett-Packard UV-VIS Spectrophotometer Model 8453 using a ferrous thiocyanate reagent at 475 nm. The Ti(IV)- $\text{H}_2\text{SO}_4$  denuder based technique has a reported detection limit for  $\text{H}_2\text{O}_2$  of 0.037 ppb ( $0.05 \mu\text{g}/\text{m}^3$ ) for one hour of sampling. This technique does take into account a correction for possible  $\text{O}_3$  interferences. Hydrogen peroxide content was calculated from a

calibration curve obtained with dilute solutions of H<sub>2</sub>O<sub>2</sub>, previously standardized with KMnO<sub>4</sub> solutions titrated against a sodium oxalate oximetric standard.

HNO<sub>3</sub> was collected in one Na<sub>2</sub>CO<sub>3</sub> coated URG annular denuder (3 channel x 30 mm O.D. x 242 mm length glass tube) at a flow rate of 10 L/min. The aqueous extracts were analyzed by ion chromatography (US EPA, 1999). A Dionex DX-120 Ion Chromatograph equipped with an AS14A IonPack column was used for the analysis. The eluent consisted of a mix of sodium carbonate and sodium bicarbonate. The standard curve had a range of 0-100 mg/L and was constructed using several dilutions of sodium nitrate. The reported detection limit for HNO<sub>3</sub> with this method is 0.78 ppb (2.0 µg/m<sup>3</sup>) for a 1-hour sampling at a flow rate of 10 L/min (US EPA, 1999).

As previously mentioned in chapter 2, an empirical model employed by the European Environmental Agency was used to determine the afternoon NO<sub>2</sub> concentrations from measured NO<sub>x</sub> concentrations and to estimate “real” NO<sub>2</sub><sup>\*</sup> concentrations (de Leeuw et al., 2001). The empirical model (BUWAL, 1997) has the form:  $NO_2 = 0.055 * NO_x + 55 * (1 - e^{-NO_x * 0.01173})$ .

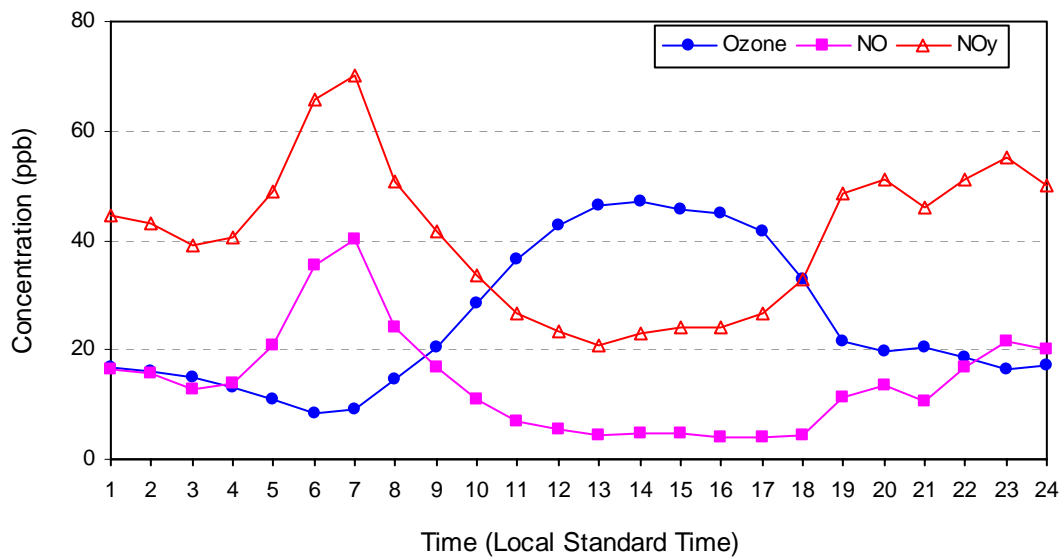
In general, the calculated NO<sub>2</sub><sup>\*</sup> value was ~ 65% of the measured non-corrected NO<sub>2</sub> value. This percentage was within to the operational uncertainty of chemiluminescent NO<sub>x</sub> analyzers suggested by Gerboles et al. (2003) for this study’s range of NO<sub>x</sub> and NO concentrations. A surrogate NO<sub>z</sub><sup>\*</sup> value was estimated from the measured NO<sub>y</sub> and NO and the estimated NO<sub>2</sub><sup>\*</sup> concentrations (ppb):

$$NO_z^* = NO_y - NO - NO_2^* \quad (13)$$

The estimations of NO<sub>2</sub><sup>\*</sup> and, as a result, NO<sub>z</sub><sup>\*</sup> were assumed valid only for the afternoon hours. It was also assumed that conversion efficiency of NO<sub>y</sub> and NO was the unity.

#### 4.1.2 General Findings

Hourly average concentrations of NO, NO<sub>y</sub>, and O<sub>3</sub> from September 3-30, 2003 are shown in Figure 4.2. The basic statistical analysis of NO, NO<sub>x</sub><sup>\*</sup>, NO<sub>y</sub>, and O<sub>3</sub> data is presented in Table 4.1.



**Figure 4.2.** Composite diurnal profiles of hourly average NO, NO<sub>y</sub> and O<sub>3</sub> concentrations measured at the Taft monitoring station in Cincinnati, September 3-30, 2003.

**Table 4.1.** Statistical summary of hourly NO, NO<sub>x</sub>\*<sup>(a)</sup>, NO<sub>y</sub> and O<sub>3</sub> and meteorological parameters<sup>(b)</sup> measured at the Taft monitoring station Cincinnati, OH (September 3-30, 2003).

Parameter	Unit	Mean	S.D.	Range	No. of Samples
NO	ppb	14.3	22.4	0.05-184	625
NO <sub>x</sub> *	ppb	36.1	28.1	5-213	625
NO <sub>y</sub>	ppb	41.2	29.1	10-213	625
O <sub>3</sub>	ppb	24.8	17.4	1-73	621
Temperature	°C	18.1	5.3	3.9-28.9	626
Wind Speed	m/sec	1.8	1.7	0-8.75	626
Wind Direction	degrees	Dominant: 225 (10.4%)		0-360	593

<sup>(a)</sup> NO<sub>x</sub> values as measured by the NO<sub>x</sub> analyzer.

<sup>(b)</sup> Measured at the Cincinnati Municipal Airport Lunke.

Hourly average values of O<sub>3</sub> ranged from 1 ppb to 73 ppb—relatively low values. The dependence of “low” O<sub>3</sub> concentrations on “low” ambient temperature was clearly evident during the study period. According to historic mean temperature data from the Cincinnati-Northern Kentucky

Airport, the mean temperature during September 2003 had a 1.44°C difference less than the normal mean for this month (NCDC, 2003). As a consequence, the O<sub>3</sub> levels were relatively low. They were even below the O<sub>3</sub> 8-hr standard of 0.08 ppm.

The hourly average O<sub>3</sub>, NO, and NO<sub>y</sub> diurnal profile showed the typical urban diurnal variation with a plateau of maximum O<sub>3</sub> concentrations in the afternoon between 13:00 and 15:00 hr and a minimum concentration early morning (~ 6:00 hr). The increase in O<sub>3</sub> during morning hours had a nearly constant positive slope that coincided with the decrease in NO concentrations. The net production (or accumulation) of O<sub>3</sub> started before 9:00 hr—once the equivalent concentration of NO titration was surpassed. The average rate of O<sub>3</sub> accumulation was ~ 5.3 ppb/hr, and the average duration of the ozone build-up was 5.2 hr. The “ridge” of average O<sub>3</sub> concentrations above the generally accepted O<sub>3</sub> background of 40 ppb extended from around 11:30 hr to shortly after 17:00 hr, with the average peak at 14:00 hr. The decrease in O<sub>3</sub> started at 16:00 hr as a result of the decreasing solar radiation and photochemical activity. After the evening, “fresh” NO emissions combined with other deposition mechanisms caused O<sub>3</sub> levels to fall.

The peak in the average hourly concentrations of NO and NO<sub>y</sub> clearly coincided with the morning 5:00 to 8:00 hr traffic. These peaks were, in general, short—with their maximum levels (40 and 70 ppb respectively) around 7 hr. After a 6-hour period, both NO and NO<sub>y</sub> reached a minimum plateau at around 13:00 hr. NO depletion rate after the start of the O<sub>3</sub> accumulation was ~ 3.3 ppb/hr. While the NO levels kept at a nearly constant level, a minimum of ~ 3 ppb from 13 to 18 p.m., NO<sub>y</sub> concentrations showed a minimum (or background) average concentration of ~ 21 ppb by the early afternoon. After this, minimum NO<sub>y</sub> levels showed a slight increasing trend, moving in the opposite direction of the trend for the O<sub>3</sub> concentrations. The difference in this pattern could be the result of regional contributions of HNO<sub>3</sub> and NO<sub>3</sub><sup>-</sup> to the NO<sub>y</sub> balance, because NO did not show a significant increase along this period. The constant levels of afternoon NO indicate that the Greater Cincinnati airshed was under the influence of constant NO<sub>x</sub> emissions. Two additional NO and NO<sub>y</sub> peaks were registered from late afternoon to

midnight, the first one associated with the evening rush hour and the second one to the resetting of the shallow nocturnal boundary layer. The reduction in  $\text{NO}_y$  levels between midnight and 3:00 hr followed the same reduction pattern as  $\text{NO}$  concentrations. Both moved in concurrence with a continuous nocturnal titration with  $\text{O}_3$ .

The set of data collected during the monitoring and sampling campaigns was statistically analyzed for the afternoon hours, as required by the PIM. Table 4.2 shows the mean, standard deviation of the mean, and range of afternoon  $\text{NO}$ ,  $\text{NO}_y$ , and calculated values of  $\text{NO}_x^*$  and  $\text{NO}_z^*$  concurrent with  $\text{O}_3 \geq 50$  ppb (WE and WD, September 3-30, 2003). The limit of 50 ppb on  $\text{O}_3$  concentrations was chosen arbitrarily to exclude days not representative of intense photochemical activity. Table 4.2 also gives the concurrent 15:00-16:00 hr measurements of  $\text{O}_3$ ,  $\text{NO}_y$ ,  $\text{HNO}_3$ ,  $\text{H}_2\text{O}_2$ , and associated  $\text{NO}_x^*$  and  $\text{NO}_z^*$  (WD only, September 8-17).

**Table 4.2.** Statistical summary of average concentrations of measured  $\text{NO}$ ,  $\text{NO}_y$  and estimated  $\text{NO}_x^*$  and  $\text{NO}_z^*$  levels concurrent with measured  $\text{O}_3 \geq 50$  ppb between 13:00-17:00 hr (September 3-30, 2003), and hourly average of concurrent measured  $\text{O}_3$ ,  $\text{NO}_y$ ,  $\text{HNO}_3$ ,  $\text{H}_2\text{O}_2$  and estimated  $\text{NO}_x^*$  and  $\text{NO}_z^*$  between 15:00-16:00 hr (September 8-17) at the Taft monitoring station.

Period	Parameter	Mean (ppb)	S.D. (ppb)	Range (ppb)	N
13:00-17:00 hr (September 3-30) <sup>a</sup>	$\text{NO}$	2.9	1.3	1-6	44
	$\text{NO}_y$	21.6	4.6	13-30	44
	$\text{NO}_x^*$	13.6	3.7	7-21	44
	$\text{NO}_z^*$	7.9	1.8	5-12	44
	$\text{O}_3$	59.2	6.0	50-73	44
15:00-16:00 hr (September 8-17) <sup>b</sup>	$\text{NO}_y$	27.8	8.9	18-40	8
	$\text{HNO}_3$	6.8	0.59	6.24-7.08	8
	$\text{H}_2\text{O}_2$	1.4	0.76	0.30-2.46	7
	$\text{O}_3$	53.7	8.9	42-66	8
	$\text{NO}_x^*$	18.9	4.9	14.5-27.9	8
	$\text{NO}_z^*$	9.0	3.7	3.1-14.6	8

<sup>a</sup> Weekend and weekdays.

<sup>b</sup> Weekdays only.

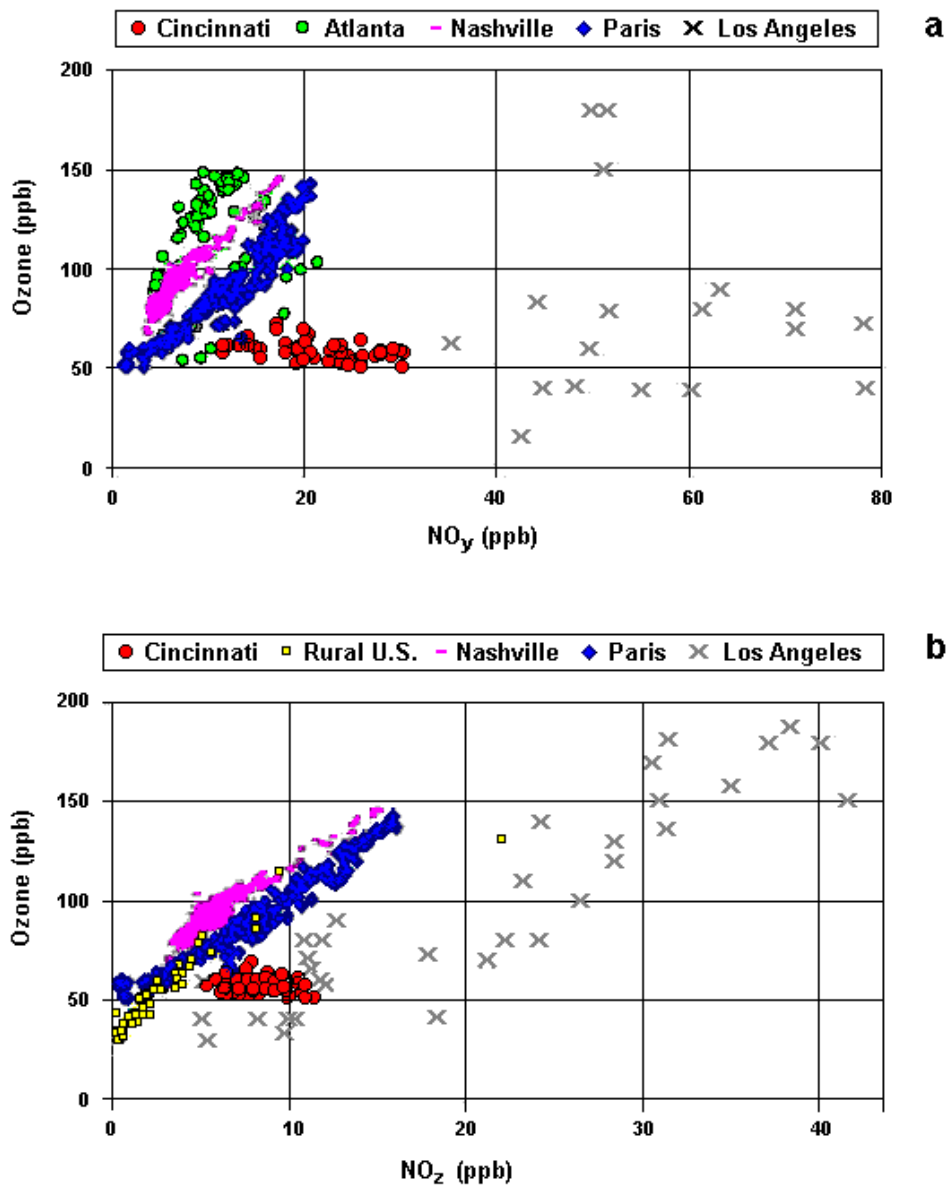
In general, the concentrations of all the compounds measured during the afternoon hours did not have strong variation. The range of the  $\text{NO}_y$  average levels (13-30 ppb) was somewhat similar to reported afternoon  $\text{NO}_y$  values for a suburban site in Tennessee (13-27 ppb) (Williams, 1998). For comparison, the range of  $\text{NO}_y$  values reported for several sites of the Los Angeles Basin in 1987 were between 25 to 140 ppb (Sillman et al., 1997). The range of measured  $\text{HNO}_3$  levels in Cincinnati (6.2-7.0 ppb) was slightly higher than that reported for a suburban site in Tennessee in the afternoon period (2.5-5.9 ppb) (Williams, 1998). However, it was approximately three times lower than the average of ranges reported for Claremont California from noon to 16:00 hr (11-18 ppb) (Hering, 1988). The range of measured  $\text{H}_2\text{O}_2$  concentrations (0.30-2.4 ppb) was within the range of measurements reported for other locations in the Ohio region. For example, Daum (1990) reported a range of  $\text{H}_2\text{O}_2$  concentrations of 0.2-6 ppb in Central Ohio for June 1987 and June 1988. Kok et al. (1978) reported a range of  $\text{H}_2\text{O}_2$  concentrations from 10-30 ppb in the Los Angeles Basin during July 1977. The range of estimated  $\text{NO}_z^*$  values of 5-12 ppb was also in agreement with the range reported for the suburban site in Tennessee (6-16 ppb) (Williams, 1998). The range of  $\text{NO}_z$  values reported for the Los Angeles Basin in 1987 was 3-52 ppb (Sillman et al., 1997).

#### 4.1.3 Photochemical Indicators Evaluation

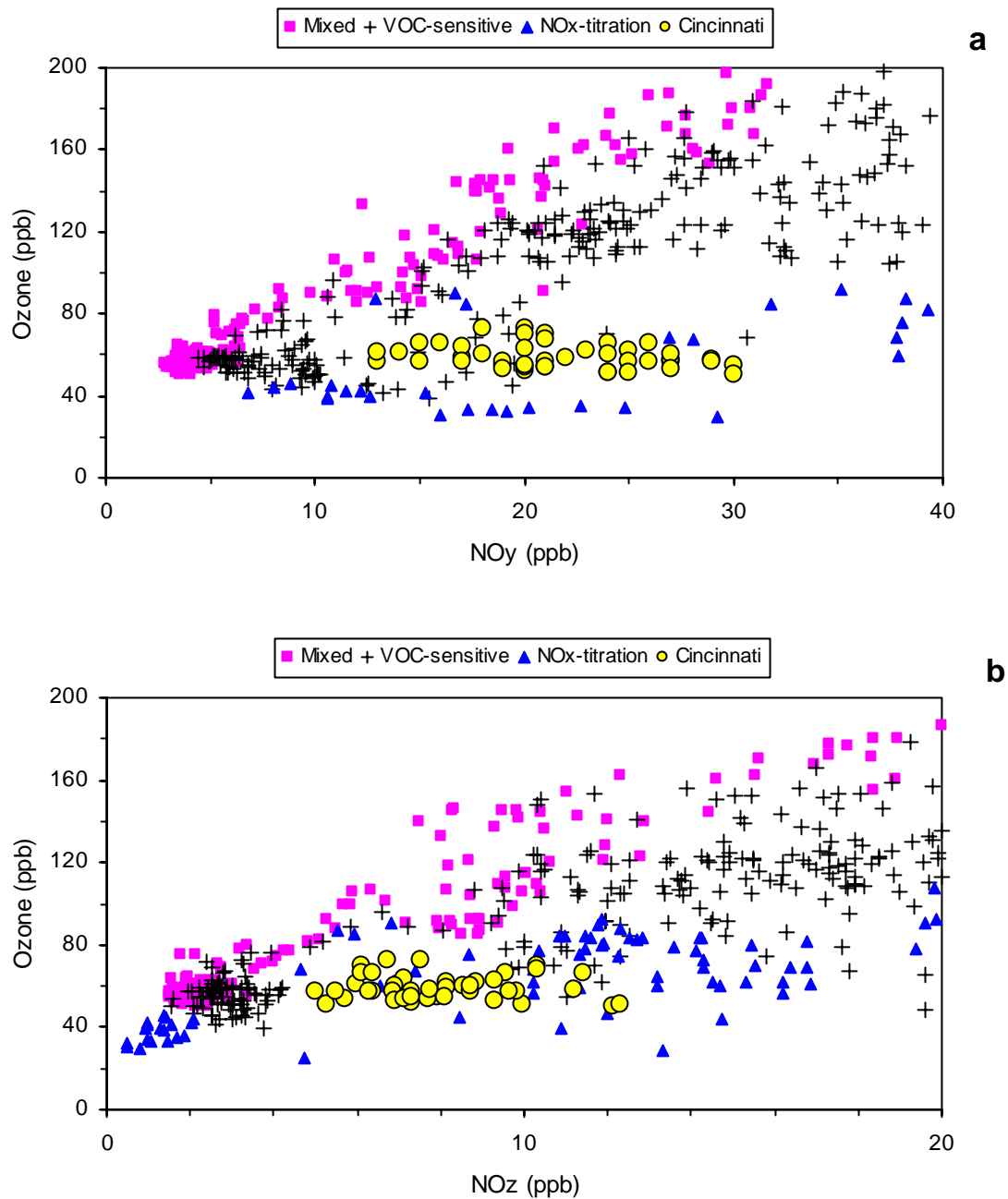
Figure 4.3 shows the comparison between hourly measured  $\text{O}_3$  ( $\geq 50$  ppb) and both  $\text{NO}_y$  and estimated  $\text{NO}_z^*$  afternoon concentrations in Cincinnati. Also shown are the reported  $\text{O}_3$  and  $\text{NO}_y$  concentrations and the  $\text{O}_3$  and  $\text{NO}_z$  concentrations measured in several representative sites of the U.S. and Paris (Sillman, 2002). This figure shows that the correlation patterns for Cincinnati tended to be more towards the VOC-sensitive conditions of the Los Angeles correlation pattern.

Figure 4.4(a) and 4.4(b) show, respectively, the comparison between the scatterplot correlation of measured  $\text{O}_3 \geq 50$  ppb,  $\text{NO}_y$ , and estimated  $\text{NO}_z^*$  from 13:00-17:00 hr and a composite of scatterplot patterns of predicted  $\text{O}_3$  versus  $\text{NO}_y$  and  $\text{NO}_z$  afternoon values for mixed-, VOC-, and  $\text{NO}_x$ -titration conditions for several regions in the U.S. (Sillman and He, 2002).





**Figure 4.3.** Scatterplot patterns of: (a) measured  $\text{O}_3$  and  $\text{NO}_y$ , and, (b) measured  $\text{O}_3$  and estimated  $\text{NO}_z^*$  in Cincinnati compared with measured correlations of  $\text{O}_3$  and  $\text{NO}_y$  and  $\text{NO}_z$  patterns for different locations identified as:  $\text{NO}_x$ -sensitive (Atlanta and Rural U.S.), mixed sensitive (Nashville and Paris), and VOC-sensitive (Los Angeles). Scatterplots adapted from Sillman (2002 and 2003).



**Figure 4.4.** Scatterplot patterns of: (a) measured  $O_3$  and  $NO_y$ , and (b) measured  $O_3$  and estimated  $NO_z^*$  at Cincinnati compared with 3-D model predicted  $O_3$  and  $NO_y$  and  $NO_z$  patterns, respectively, for locations with mixed or with near-zero sensitive conditions (squares), VOC-sensitive conditions (crosses), and dominated by  $NO_x$ -sensitive titration (triangles). Simulation data obtained from data file available from Sillman (2003).

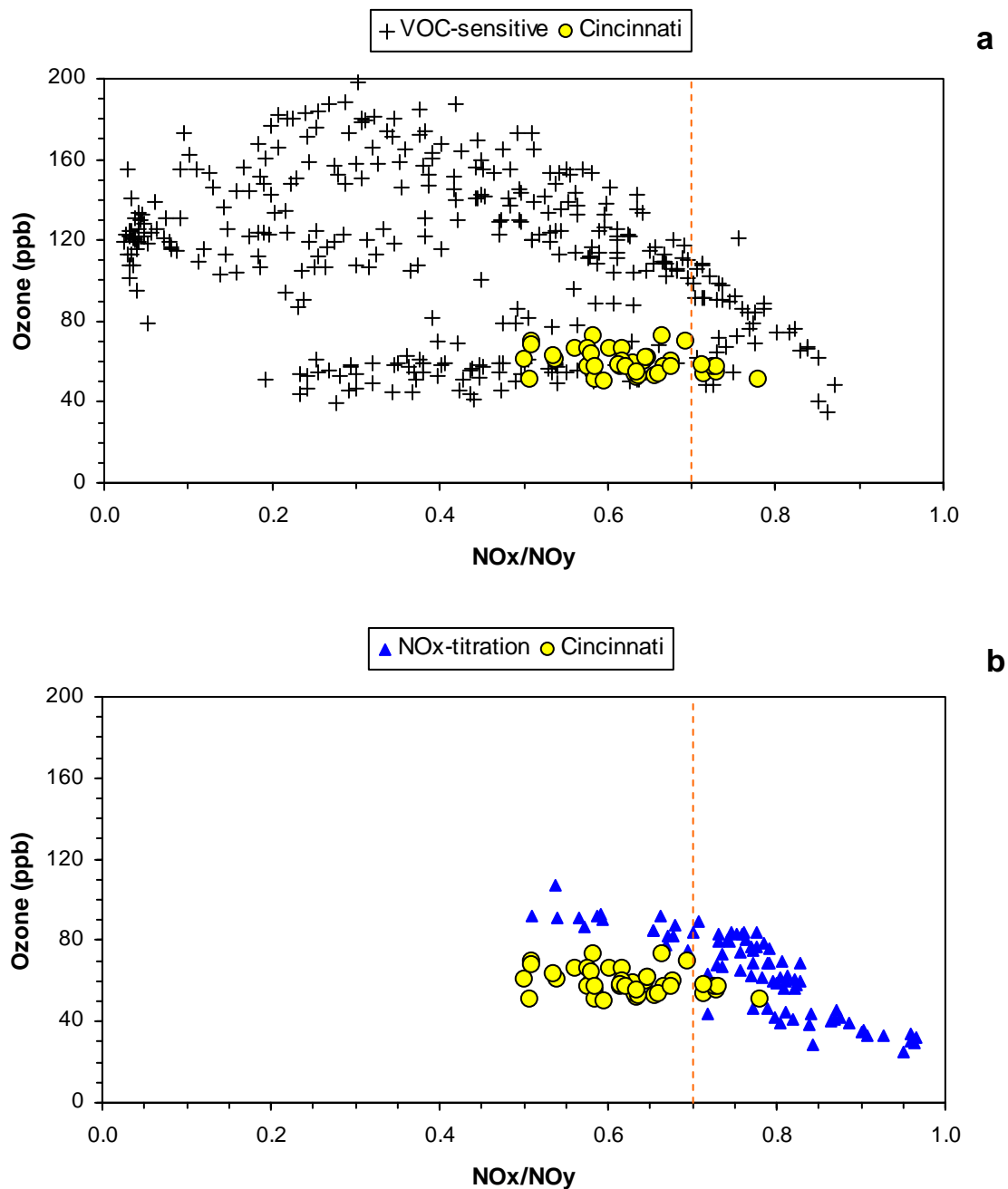
According to Figure 4.4(a), the  $O_3$  vs.  $NO_y$  data points for Cincinnati matched with the predicted VOC-sensitive condition and a great portion of the  $NO_y$  data points were above 20 ppb (VOC-sensitive condition according Sillman, 1995). However, the outline of the  $O_3$  vs.  $NO_z^*$  scatterplot pattern, shown in Figure 4.4(b), suggests the possibility that the air masses sampled could be under conditions dominated by  $NO_x$ -titration.

Figures 4.5(a) and 4.5(b) show the results of the empirical procedure outlined in chapter 2 to identify VOC-sensitive conditions from those dominated by  $NO_x$ -titration environments, which involves the associated  $NO_x/NO_y$  ratios. Figure 4.5(a) shows that most of the data points of the  $NO_x^*/NO_y$  ratios for Cincinnati matched perfectly with a portion of the scatterplot pattern for locations with VOC-sensitive conditions. Figure 4.5(b) demonstrates that just a few of these data points overlapped the  $NO_x$ -titration pattern and had a  $NO_x^*/NO_y$  ratio of less than 0.7.

Although the  $NO_x^*$  values represented an approximation of the real  $NO_x$  concentrations, the results displayed in Figure 4.5 one could argue that the air masses that arrived at the monitoring site were not significantly impacted by local sources of NO and that, independent of origin and age, they were relatively homogeneous in their maximum  $O_3$  concentrations.

On the other hand, one attempt was made to evaluate the consistency of the measurements of  $HNO_3$  and  $H_2O_2$  by a comparison of the correlations of measured  $O_3$  versus  $2H_2O_2 + NO_z^*$  with simulated values (as explained in chapter 2). However, the small number of observations (8 pairs) and the very low values of concurrent  $O_3$  resulted in a small cluster located at the lower edge of the scatterplot. As an alternative, the  $O_3/(2H_2O_2 + NO_z^*)$  ratio was calculated, resulting in a value of  $\sim 4.7$ . This value was very close to the lower limit of the range (5-8) given by Sillman (2002).

Although, at the Taft monitoring site, there were no measurements of CO to verify the consistency of the registered  $NO_y$  levels, the ensemble of qualitative evaluations allowed one to conclude that the measurements of  $NO_y$ , the estimated  $NO_z^*$ , and the  $HNO_3$  and  $H_2O_2$  indicator species were consistent with the indicator theory and that they could be used to evaluate the  $O_3$ - $NO_x$ -VOC sensitivity of Cincinnati with the PIM.



**Figure 4.5.** Scatterplot patterns for afternoon  $\text{O}_3$  versus  $\text{NO}_x^*/\text{NO}_y$  data for Cincinnati compared with the calculated  $\text{NO}_x/\text{NO}_y$  ratio from data results of 3-D model simulations for locations with: (a) VOC-sensitive conditions, and (b) dominated by  $\text{NO}_x$ -titration. The vertical dashed line represents the hypothetical limit for differentiating photochemically aged air parcels. Simulation data obtained from data file available from Sillman (2003).

In addition, the comparisons of the indicator data for Cincinnati with measured and predicted correlations for other locations were consistent with the identification of VOC-sensitive conditions.

Table 4.3 shows the comparison of the calculated photochemical indicator ratios from measurements performed in Cincinnati with the mean transition values proposed by Sillman (1995) and Sillman and He (2002). Figure 4.6 displays the scatterplots of measured  $O_3$  versus  $NO_y$  and  $O_3$  versus surrogate  $NO_z^*$  in Cincinnati between 13:00 and 17:00 hr for late September 2003. It also compares them with the suggested mean transition lines for  $NO_x$ - and VOC-sensitive conditions for “clean” air conditions with  $O_3$  maximum levels < 80 ppb, as is the case for Cincinnati. The combined results from Table 4.3 and Figure 4.6 provide a confirmation of the dominance of VOC-sensitive conditions in Cincinnati according to the PIM.

**Table 4.3.** Calculated average indicator ratios for data observed at the Taft Monitoring Station (Cincinnati) and transition ranges according to Sillman (1995) and Sillman and He (2002). Values of ratios lower than their transition ranges indicate VOC-sensitivity. Values of  $NO_y$  higher than the transition value indicate VOC-sensitive conditions.

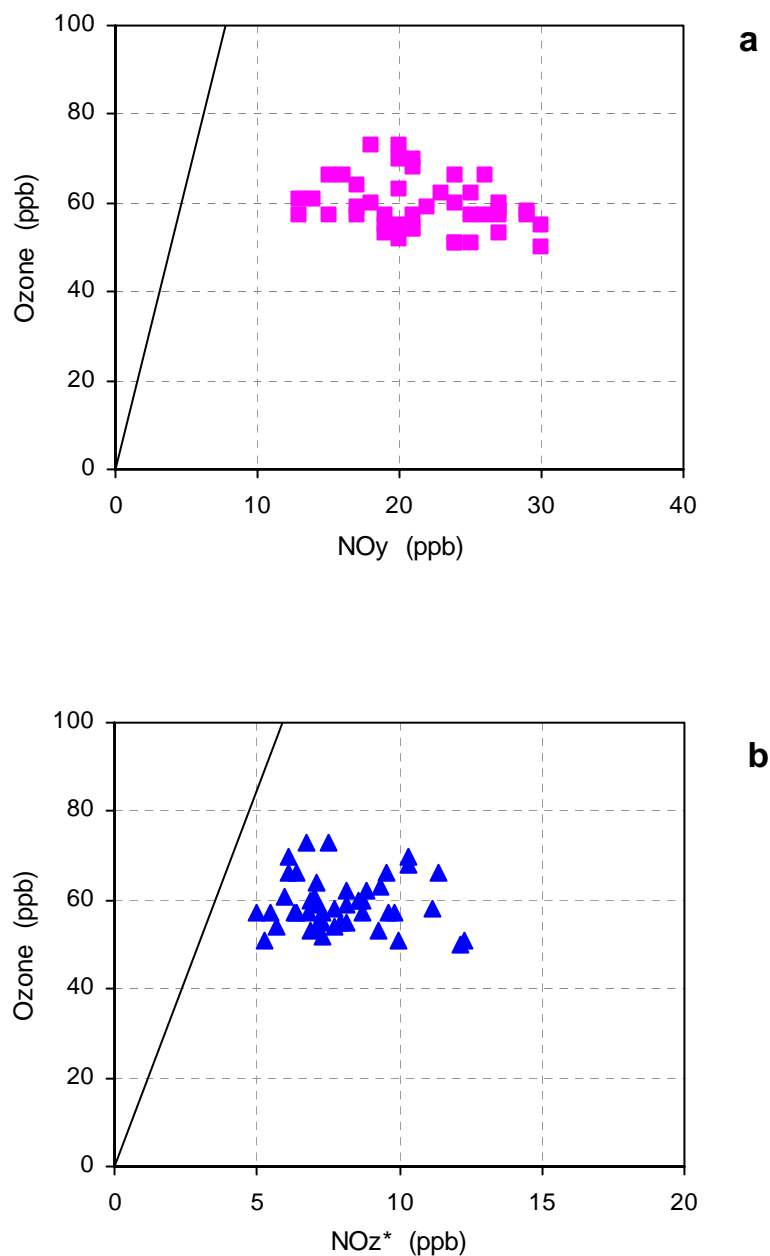
Indicator	Taft average ratios	Transition range
$O_3/NO_y$	2.7 <sup>a</sup> -1.93 <sup>b</sup>	11–15 <sup>c</sup>
$O_3/NO_z$	5.96 <sup>a</sup> -5.92 <sup>b</sup>	15–20 <sup>c</sup>
$O_3/HNO_3$	7.89 <sup>b</sup>	20–25 <sup>c,d</sup>
$H_2O_2/NO_y$	0.05 <sup>b</sup>	0.12-0.17
$H_2O_2/NO_z$	0.15 <sup>b</sup>	0.20-0.25
$H_2O_2/HNO_3$	0.22 <sup>b</sup>	0.25-0.35
$NO_y$	21.6 <sup>a</sup> -27.8 <sup>b</sup>	20

<sup>a</sup> From the whole monitoring period

<sup>b</sup> From the weekdays sampling period.

<sup>c</sup> Transition range for low ozone (< 80 ppb) polluted conditions.

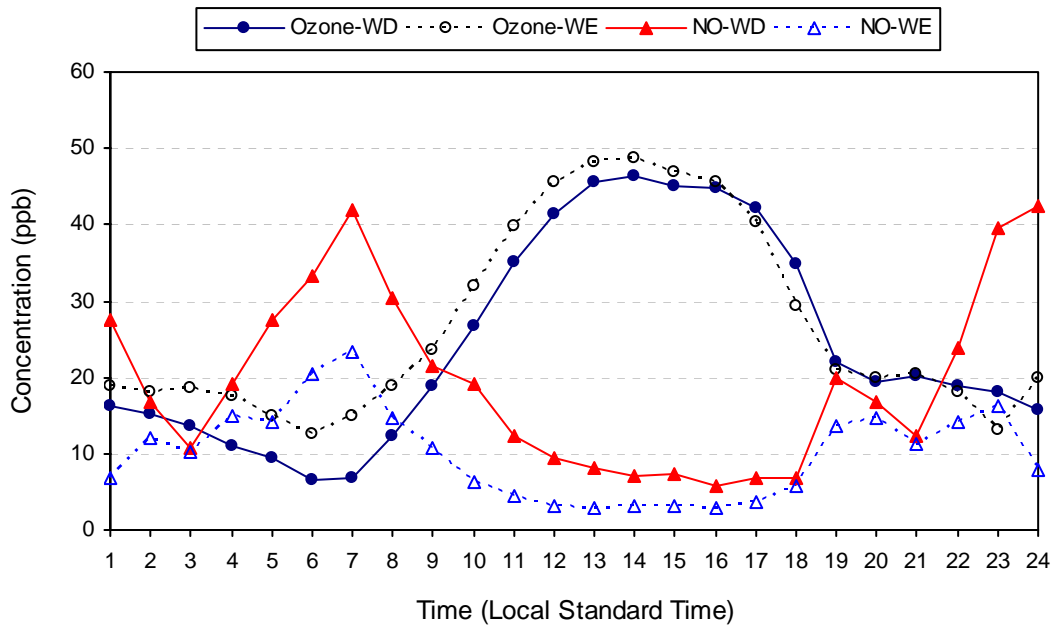
<sup>d</sup> Interpolated from Figure 3 in Sillman and He (2002).



**Figure 4.6.** Scatterplot patterns of: **(a)** measured O<sub>3</sub> and concurrent NO<sub>y</sub> and **(b)** measured O<sub>3</sub> and concurrent estimated NO<sub>z</sub>\* from 13:00 to 17:00 hr, September 3-30, 2003 at Cincinnati. The lines represent the transition between NO<sub>x</sub>- and VOC-sensitive chemistry for O<sub>3</sub>/NO<sub>y</sub> and O<sub>3</sub>/NO<sub>z</sub> indicator ratios for low O<sub>3</sub> conditions (≥ 50 ppb) according to Sillman and He (2002).

#### 4.1.4 Weekend/Weekday Effect Analysis

Figure 4.7 presents the average hourly diurnal weekend/weekday (WE/WD) variations of  $O_3$  and NO from September 3-30, 2003. Table 4.4 presents the results of the tests for statistical significance in WE and WD for 1-hr  $O_3$  maximum, 8-hr  $O_3$  maximum average, and 6:00-9:00 hr NO average for the same sampling period.



**Figure 4.7.** Average weekend and weekday hourly diurnal variations of  $O_3$  and NO concentrations (September 3-30, 2003) measured at the Taft monitoring station, Cincinnati.

Average  $O_3$  concentrations were lower during WD morning rush hours than in WE. This likely caused by the  $O_3$  titration resulting from NO emitted by the heavy traffic during WDs. However, WE  $O_3$  morning concentrations were around 7 ppb higher than those on WDs. Reduced diesel truck and gasoline vehicles traffic in WE could be associated with the reduction in  $NO_x$  emissions leading to a short supply of NO for morning titration. Accordingly, the average NO- $O_3$  crossover time in WE occurred about 1-hr earlier than in WD, and the duration of the build-up and accumulation of  $O_3$  was also higher in WE (~ 6.3 hr) than in WD (~ 4.7 hr). The average WE/WD rate of accumulation of  $O_3$  was also somewhat different.

It was a little faster in WE (5.8 ppb/hr) as compared with WD (5.1 ppb/hr). This change in efficiency and rate of WE /WD O<sub>3</sub> formation was probably due to reduced WE NO<sub>x</sub> emissions and the associated higher WE VOC/NO<sub>x</sub> ratio. This same phenomenon has been identified in California, where O<sub>3</sub> formation is dominated by VOC-sensitive conditions (Fujita et al., 2002).

**Table 4.4.** Results of the tests for statistical significance for WE/WD differences of maxima 1-hr and 8-hr average O<sub>3</sub> concentrations and 6:00-9:00 hr NO average levels for data registered between September 3-30, 2003, at the Taft monitoring station, Cincinnati.

Parameter	Mean (S.D.) ppb	No. of samples	Results of tests of significance	Interpretation
WE max 1-hr O <sub>3</sub>	48.6 (14.4)	8	$p = 0.721$	$H_0^b$ : $\mu_{1\text{-hr O}_3\text{WE}} = \mu_{1\text{-hr O}_3\text{WD}}$ Accepted
WD max 1-hr O <sub>3</sub>	46.4 (14.8)	17		$H_I^c$ : $\mu_{1\text{-hr O}_3\text{WE}} < \mu_{1\text{-hr O}_3\text{WD}}$ Rejected
WE max 8-hr O <sub>3</sub>	43.3 (25.4)	8	$p = 0.739$	$H_0^b$ : $\mu_{8\text{-hr O}_3\text{WE}} = \mu_{8\text{-hr O}_3\text{WD}}$ Accepted
WD max 8-hr O <sub>3</sub>	41.6 (11.6)	17		$H_I^c$ : $\mu_{8\text{-hr O}_3\text{WE}} < \mu_{8\text{-hr O}_3\text{WD}}$ Rejected
WE 6-9 a.m. NO	16.1 (25.4)	54	$p = 0.021$	$H_0^d$ : $\mu_{6\text{-}9\text{am NO WE}} = \mu_{6\text{-}9\text{am NO WD}}$ Rejected
WD 6-9 a.m. NO	31.7 (27.5)	24		$H_I^e$ : $\mu_{6\text{-}9\text{am NO WE}} > \mu_{6\text{-}9\text{am NO WD}}$ Accepted

<sup>a</sup> If  $p \leq 0.05$ , the difference is significant; if  $0.05 \leq p \leq 0.10$  the difference between the means is marginally significant, if  $p > 0.10$ , the difference is insignificant.

<sup>b</sup>  $H_0$ : The change in WE maximum O<sub>3</sub> is insignificant, likely to be caused by a VOC-sensitive condition provided that the reduction in morning WE-NO<sub>x</sub> emissions is significant.

<sup>c</sup>  $H_I$ : The decrease in maximum WE maximum O<sub>3</sub> is significant, likely due to a NO<sub>x</sub>-sensitive condition on condition that morning WE-NO<sub>x</sub> emissions is significant.

<sup>d</sup>  $H_0$ : The reduction in morning WE-NO<sub>x</sub> emissions is insignificant probably due to mere chance.

<sup>e</sup>  $H_I$ : The reduction in morning WE-NO<sub>x</sub> is significant likely due to the reduction in WE activities.

The higher WE O<sub>3</sub> morning levels converted into higher afternoon levels relative to the WD O<sub>3</sub> peak levels. The difference between the average 13:00-15:00 hr WE/WD O<sub>3</sub> values was on the order of 2 ppb. As shown in Table 4.4 however, the difference in the 1-hr maximum WE/WD O<sub>3</sub> was not statistically



significant at a 95% confidence level. Shortly after 16:00 hr, average WE and WD O<sub>3</sub> concentration profiles crossed over and had a strong decreasing that suddenly ceased around 19:00 hr when both WE and WD O<sub>3</sub> were almost equal. From midnight to predawn, WE O<sub>3</sub> had a higher concentration than in WD.

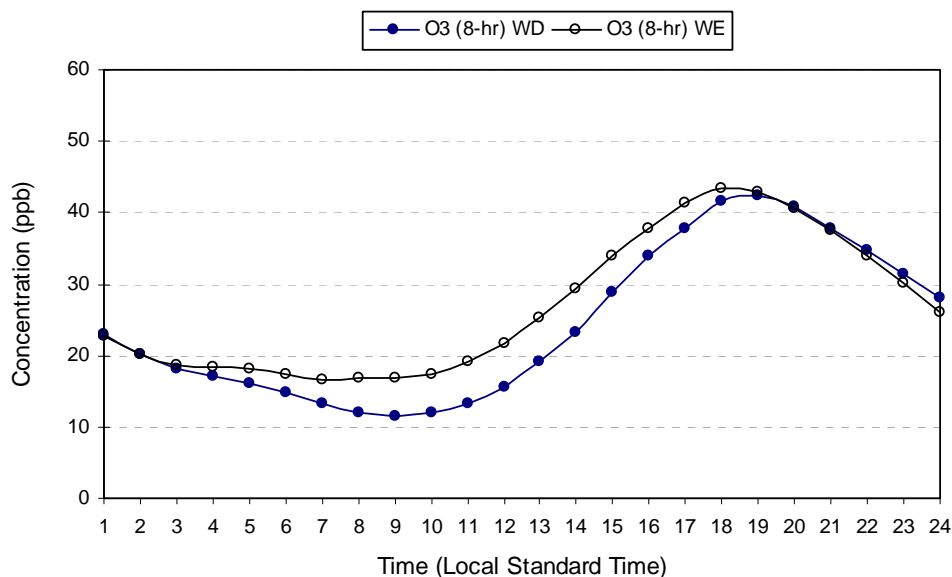
The most obvious variation, shown in Figure 4.7, was in the NO WE/WD hourly average profile. Although the morning WE and WD NO peaks occurred at the same hour (~ 7 a.m.), the WD NO peak was almost 1.8 times higher than the WE NO peak. The previously suggested association between the reduction in WE NO<sub>x</sub> emissions and the reduction in traffic seems to be confirmed by this difference.

Moreover, the importance of morning WE NO<sub>x</sub> emissions reduction for afternoon WE O<sub>3</sub> concentrations is confirmed by the results shown in Table 4.4. The reduction in 6:00-9:00 hr average WE NO levels compared to the respective morning WD NO average concentration was statistically significant at a 95% confidence level. After the occurrence of the NO morning peaks, the diurnal NO levels were dominated by the WE levels. This difference converged around 19:00 hr. There was a second short NO peak in the evening, due likely to evening traffic. This peak was higher for WE compared to WD. A third nocturnal NO peak was observed in average WD NO concentrations equivalent in level to the WD morning peak but longer in duration. A careful examination of the hourly data indicated that, although the occurrence of very high NO concentrations around midnight did not follow a constant pattern, the measured values affected the calculated hourly averages. It is possible that the combination of a strong reconstruction of the shallow boundary layer with the typical nocturnal NO<sub>x</sub> emissions could lead to such accumulation in the NO levels.

Because in the statistical analysis results showed in Table 4.4 one sample (WD) was almost 2.5 times as large as the other (WE), two additional evaluations were performed to corroborate these results. One was the use of a stricter two-tail *t-test* with unequal variances and another was the testing of the differences in the WE and WD means from a matrix of 2 days from the weekdays data set in order to have the closest population numbers (equal variances). In all cases the results of the tests were the same as those shown in Table 4.4. Appendix C presents a summary of the set of two WE/WD combinations.

The O<sub>3</sub> 8-hr average diurnal WE/WD patterns (see Figure 4.8) followed the same outline as those for the O<sub>3</sub> 1-hr patterns. They evidenced higher levels on WE than on WD during the daytime. The highest difference of ~ 6 ppb occurred at 11:00 hr. The WE O<sub>3</sub> 8-hr average peak occurred at 17:00 hr with a difference of less than 2 ppb with respect to the respective WD average. Testing the results of the WE/WD O<sub>3</sub> 8-hr average peaks showed that there was not a statistically significant difference between the peaks.

The statistical analyses of the WE/WD differences suggest that, even though the reduction in WE NO<sub>x</sub> emissions (represented by the morning average NO concentrations) was significant, both average maximum 1-hr and 8-hr WE O<sub>3</sub> average concentrations were not significantly lower than the respective WD averages, as would be expected in dominant NO<sub>x</sub>-sensitive conditions. On the contrary, the 1-hr and 8-hr WE O<sub>3</sub> peaks were somewhat higher than the respective WD averages, although not statistically significant resembling a situation somewhat similar of VOC-sensitive conditions.

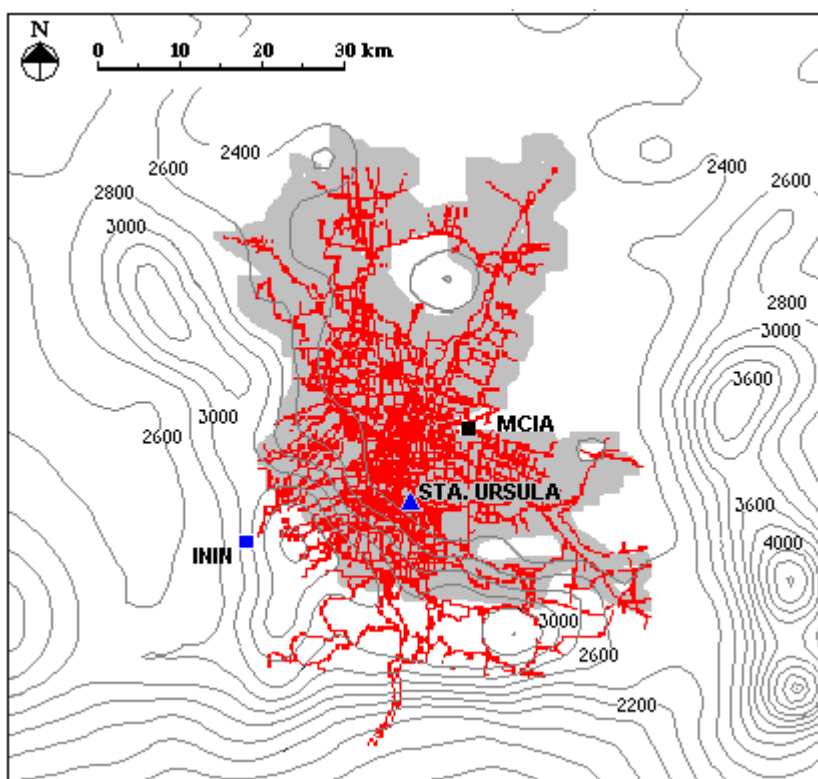


**Figure 4.8** Average weekend and weekday 8-hr diurnal variations of O<sub>3</sub>, September 3-30, 2003, for O<sub>3</sub> data registered at the Taft monitoring station, Cincinnati.

## 4.2 Mexico City Metropolitan Area

### 4.2.1 Indicator Species Measurements

Two monitoring campaigns were carried out from mid-February to mid-April 2004. The longest one was conducted at the urban smog-receptor site of Santa Ursula (19.3° N; 99.14° W, 2300 m.a.s.l.) in the south of Mexico City. The shortest was carried out at the downwind rural reference site of the Instituto Nacional de Investigaciones Nucleares (ININ) (Lat. 19.28° N; 99.38° W; 3119 m.a.s.l.) in the southwest mountains of the MCMA region (Figure 4.9).



**Figure 4.9.** Map showing the Mexico City Metropolitan Area (MCMA) and the location of the Santa Ursula and the ININ monitoring sites. The gray area shows the approximate extent of the MCMA. MCIA represents the Mexico City International Airport. The dark-gray lines show the approximate elevation curves in the region. (Adapted from García, 2002).

#### 4.2.1.1 Santa Ursula Monitoring Site

The Santa Ursula monitoring site (SUMS) is a typical urban-representative monitoring station of the official Automated Air Monitoring Network of the Government of Mexico City (RAMA). The SUMS is located within a typical low-medium income urban neighborhood of Mexico City. The site is not affected by any important fixed emission source and receives only the influence of the local traffic. The station is equipped with the following instruments: a NO/NO<sub>2</sub>/NO<sub>x</sub> API Model 200 Chemiluminescent Analyzer, an O<sub>3</sub> API Model 400 Photometric Analyzer, a CO API Model 300 Infrared Analyzer, a SO<sub>2</sub> API Model 300 Fluorescence UV Analyzer, and a PM<sub>10</sub> Rupprecht and Patashnick Model 1400 TEOM Analyzer. This station also measures other meteorological parameters. NO<sub>y</sub> measurements were made with a chemiluminescence method using the NO<sub>2</sub>/NO<sub>x</sub> NO<sub>x</sub> API Model 200 analyzer (or NO<sub>y</sub>\* analyzer) with an external heated converter mounted at the gas inlet. Personnel of the RAMA routinely calibrated the instruments following US EPA standard procedures.

The heated converter used in the SUMS consisted of a catalytic stainless steel (SS) converter assembly removed from a Columbia Scientific Industries Chemiluminizer Model NA510-2 NO<sub>x</sub> analyzer. The SS tube was packed with ~ 1 g of molybdenum screen (12 x 12 wires cm<sup>-1</sup>) (ThermoShield, Los Altos, CA). The converter was temperature controlled with a thermocouple embedded in the assembly and connected to a separate lab-made temperature controller. The temperature controller kept the temperature of the converter at approximately 370-385°C. Both the converter assembly and the temperature controller were located inside a small shelter that was mounted on the SUMS roof.

A Teflon tube (0.6 cm OD x 25 cm from the tip to heated converter) was used as the inlet line to the converter, while a ~ 3.5 m long Teflon tube connected the outlet of the converter to the NO<sub>y</sub>\* analyzer located inside the air-conditioned monitoring shelter. The NO<sub>y</sub>\* inlet tubing was not heated. A 47 mm Teflon filter pack located at the sampling inlet of the NO<sub>y</sub>\* analyzer protected the instrument from MnO dust. With the aim of keeping a relatively constant flow through the converter, a bypass pump was located ~ 0.8 m from the sample inlet of the instrument. The flow through the converter was ~ 1.5 L/ min. A number of tests (with- and without the Mo mesh packing) were carried out in the field to investigate the

converter configuration and operating conditions. The  $\text{NO}_y^*$  analyzer was operated without any plumbing modification and with the original cycling sampling modes ( $\text{NO}/\text{NO}_x$ ). The  $\text{NO}_x$  signal from the  $\text{NO}_y^*$  analyzer was taken as the  $\text{NO}_y$  final reading. The conversion efficiency of the external converter was checked indirectly by comparing the  $\text{NO}_2$  output reading of the  $\text{NO}_y^*$  analyzer with the concurrent  $\text{NO}_2$  reading of the other  $\text{NO}_x$  API 400 instrument dedicated to measuring  $\text{NO}_x$  species. Both the  $\text{NO}_y^*$  and  $\text{NO}_x$  analyzers were periodically calibrated according to the US EPA  $\text{NO}_x$  calibration method. The  $\text{NO}_x$  analyzer sample inlet was connected to the common manifold of the monitoring shelter. The converter efficiency of the packed SS converter was calculated to be  $> 95\%$ .

In order to estimate the  $\text{NO}_z^*$  concentrations, the empirical model (BUWAL, 1997) was applied to the SUMS  $\text{NO}_x$  data. Due to technical difficulties, it was not possible to operate denuder samplers to collect  $\text{H}_2\text{O}_2$  and  $\text{HNO}_3$  at this site. Additional meteorological information was obtained from the Preparatoria 5 Meteorological Station located  $\sim 5$  km to the east of the SUMS.

#### **4.2.1.2 ININ Site**

The ININ site is a research institute facility, about 35 miles southwest of downtown Mexico City, located in a semi-forested area partially insulated by a series of mountains. Because of its location downwind of the MCMA, the site is ideally suited for monitoring of the transport of air masses from the MCMA. The closest source of emissions to this site is a highway located  $\sim 3$  miles to the north.

An equipped monitoring trailer belonging to the Atmospheric Physical-Chemical Group of the Center for Atmospheric Sciences (APC-CAS) was located in an insulated part of the ININ facility. The monitoring trailer was equipped with the following instruments: a  $\text{NO}/\text{NO}_2/\text{NO}_x$  TEI Model 42C Chemiluminescent analyzer, an  $\text{O}_3$  API Model 400 Photometric Analyzer, a CO TEI Model 48C Infrared Analyzer, and a Non-Methane HC TEI Model 55C, as well as a URG Annular Denuders System for  $\text{HNO}_3$  integrated sampling.  $\text{NO}_y$  measurements were carried out with the  $\text{NO}/\text{NO}_2/\text{NO}_x$  TEI 42C analyzer with an external heated catalytic converter (molybdenum). Calibration of all the instruments was carried out at

the site using standard gases and US EPA protocols. Meteorological parameters were measured with a Davis Weather Monitor III station.

The external converter consisted of an experimental prototype constructed in the electronics workshop of the Center for Atmospheric Sciences. It was made up of a straight stainless steel (SS) tube (0.6 cm x 25 cm) loosely packed with approximately 1 g of molybdenum screen (12 x 12 wires cm<sup>-1</sup>) (Thermoshield, Los Altos, CA). The tube was placed in the center of two-channeled aluminum heating plates (12 cm x 10 cm x 2.5 cm each). This “oven” was heated by four concentrically separated tube-type heating cartridge elements (~ 10 cm long) inserted in each of plates. A thermocouple was inserted parallel to the SS tube in one of the two plates. The oven was inserted into a box made of ~ 5 cm thick solid insulating material. A Met-One data logger connected to a Campbell CR10X-console controlled the temperature of the converter. The operating temperature of the SS converter was set at 375°C.

The converter assembly was located inside a small metallic housing on the roof of the monitoring trailer, a configuration similar to the Santa Ursula site. The temperature controller system was placed inside a separated polypropylene-shelter box that was itself placed underneath the converter assembly. The inlet line to the converter consisted of a ~ 20 cm Teflon tube (0.6 cm OD). Another ~ 3 m Teflon tube (~ 0.6 cm OD) was extended from the outlet of the converter assembly to one of the inlets of an air sample diverter (a 3-way solenoid) connected directly to the NO<sub>x</sub> analyzer. The other inlet of the 3-way solenoid valve was connected to a Teflon tube (~ 0.6 cm OD x 50 cm) that was inserted into the common manifold of the monitoring trailer. The purpose of this diverter was to allow the same NO<sub>x</sub> instrument the measuring of both NO<sub>y</sub> (from the external converter) and NO/NO<sub>2</sub>/NO<sub>x</sub> (from the common manifold) concentrations in alternate equal short time cycles. A timer that switched the pass of the air sample to the analyzer every five minutes controlled the diverter. The out sample of the diverter was connected to a 47 mm diameter filter pack located at the sampling inlet of the analyzer to protect the instrument from MnO dust. In order to keep a constant flow of air through the converter during the “common manifold” diverter mode, an additional pump pulled the NO<sub>y</sub> sample outside of the trailer at the same flow rate as the instrument.

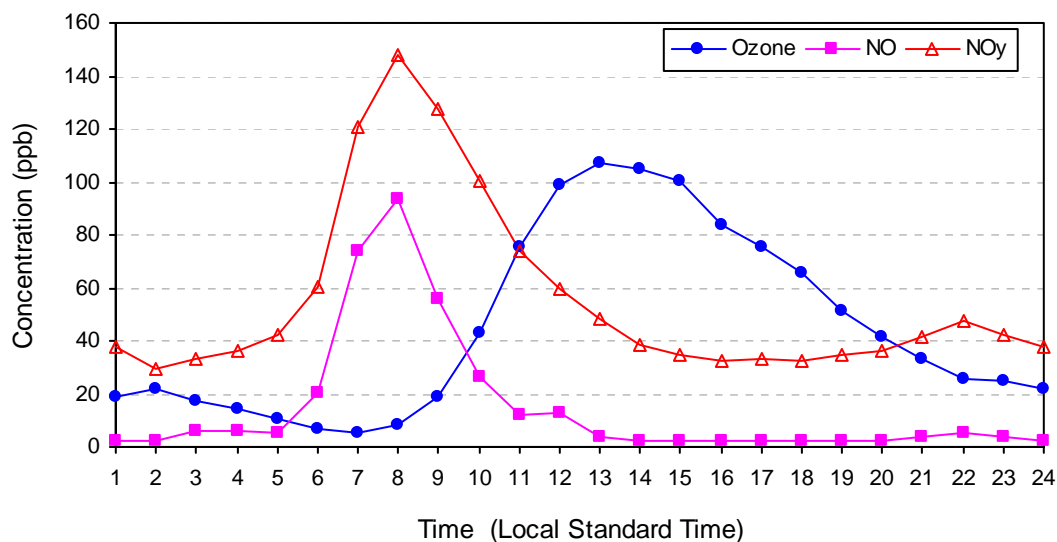
The conversion efficiency of the external Mo converter was checked indirectly by comparing the NO<sub>2</sub> readings from the two inlet air sampling modes (i.e., external converter and common manifold flows). If the converter was operating correctly, under the “external converter” sampling mode, the NO<sub>2</sub> readings in the instrument should be “zero”, because all the NO<sub>2</sub> that had passed previously across the external Mo converter had been reduced to NO. The NO<sub>2</sub>\* concentrations for the estimation of NO<sub>z</sub>\* were derived using the BUWAL equation and the NO<sub>x</sub> readings taken from the “common manifold” air sampling mode.

Several attempts were made for sampling HNO<sub>3</sub> at this site. Instead of the typical 1-hr sampling period, the annular denuder system was operated for two hours in order to enhance the capture of gaseous HNO<sub>3</sub>. Nitric acid was collected in one Na<sub>2</sub>CO<sub>3</sub> coated URG annular denuder (3 channel x 30 mm O.D. x 242 mm length glass tube) at a sampling flow rate of 10 L/min. The extractions were done in the following 24 hours. The aqueous extracts were analyzed with a Perkin Elmer Liquid Chromatograph (Isocratic LC Pump 250) conductivity detector equipped with a Hamilton PRP-X100 IC anion column. The extractions and the analysis were performed at the APC-CAS laboratories. As well as in the SUMS gaseous hydrogen peroxide could not be sampled in this site.

## **4.2.2 General Findings**

### **4.2.2.1 Santa Ursula Monitoring Site**

The monitoring campaign at SUMS consisted of two stages: the first from February 18 to March 11, 2004, with the objective of testing the converter configurations; and the second from April 14 to 25, 2004 to measure, O<sub>3</sub>, NO<sub>y</sub>, NO and NO<sub>x</sub> concentrations. Only the results of the measurement campaign are presented in this dissertation. Figure 4.10 shows the hourly average diurnal profiles of O<sub>3</sub>, NO<sub>y</sub>, and NO concentrations measured at SUMS from April 14-25. Table 4.5 presents the general statistical summary of the hourly measurements for the same monitoring period including the statistical summary of 30-min measurements of some meteorological parameters.



**Figure 4.10.** Composite diurnal profiles of hourly average NO, NO<sub>y</sub> and O<sub>3</sub> concentrations measured at the Santa Ursula monitoring station in Mexico City, April 14 to 25, 2004.

**Table 4.5.** Statistical summary of hourly average NO, NO<sub>x</sub>\*<sup>(a)</sup>, NO<sub>y</sub> and O<sub>3</sub> concentrations and meteorological parameters<sup>(b)</sup> measured at the Santa Ursula monitoring station, Mexico City (April 14-27, 2004).

Parameter	Unit	Mean	S.D.	Range	No. of Samples
NO	ppb	14.8	29.8	0.05-183	271
NO <sub>x</sub> * <sup>(a)</sup>	ppb	45.3	38.2	11-226	271
NO <sub>y</sub>	ppb	54.7	39.0	16-228	243
O <sub>3</sub>	ppb	44.7	37.1	4-172	286
Temperature	°C	19.7	4.86	9.1-28.2	491 <sup>c</sup>
Wind Speed	m/sec	1.57	1.64	0-7.2	480 <sup>c</sup>
Wind Direction	degrees	Dominant: 360 (22.12%)		0-360	480 <sup>c</sup>

<sup>a</sup> NO<sub>x</sub> values as measured by the NO<sub>x</sub> dedicated instrument.

<sup>b</sup> Measured at the "Preparatoria 5" High School.

<sup>c</sup> 30-min average

Diurnal average ozone concentrations followed the typical urban pattern, with the peak between 12:00 to 15:00 hr and the lowest levels during the morning rush hours (6:00-8:00 hr). Average O<sub>3</sub> photochemical production started around 7:00 hr, shortly after the occurrence of the NO peak (~ 92 ppb),



and reached its maximum (~ 107 ppb) when NO levels dropped to their minimum (~4 ppb). The net build-up of O<sub>3</sub> started before 10:00 hr and the duration of effective ozone accumulation lasted around 3.3 hrs. The average rate of accumulation was ~ 21.4 ppb/hr. The ridge of levels above 80 ppb extended for approximately 5 hr. The period with O<sub>3</sub> peak was relatively short and, after 14:00 hr, ozone levels started to decrease, following a smooth slope. The mean of O<sub>3</sub> concentrations was 44.7 ppb and hourly average values ranged from 4 to 172 ppb. A detailed review of the data showed that the Mexican Ozone Air Quality Standard (0.12 ppm max 1-hr) was exceeded 17 times during this 12-day monitoring period.

On the other hand, the diurnal average NO concentrations pattern showed one peak between 6:00-9:30 hr. The NO raised and reached a maximum around 8:00 hr (~ 92 ppb) that coincided with the maximum emissions of NO during the day, in this case, morning peak vehicle traffic. Subsequently, NO concentrations dropped at an average rate of ~ 17 ppb/hr, reaching a low level plateau of less than 10 ppb for the rest of the day. Average NO<sub>y</sub> concentrations had a pattern similar to the NO levels but with a second small peak around 22:00 hr. There was a longer “valley” of NO<sub>y</sub> average concentrations (between 38 and 32 ppb) from 14:00-18:00 hr, before the occurrence of the second evening peak (~ 45 ppb). The meteorological conditions during this campaign were typical for this time of the year, with predominantly light winds coming from the north and, with less frequency, from the northeast. The afternoon hourly concentrations of NO, NO<sub>y</sub>, and O<sub>3</sub>, as well as estimated NO<sub>x</sub>\* and NO<sub>z</sub>\* values were statistically analyzed with the same method as the Cincinnati data (Table 4.6).

**Table 4.6.** Statistical summary of hourly average concentrations of measured O<sub>3</sub>, NO, NO<sub>y</sub> and estimated NO<sub>x</sub>\* and NO<sub>z</sub>\* between 13:00-17:00 hr (April 14-25, 2004) at the Santa Ursula monitoring station, Mexico City<sup>a</sup>.

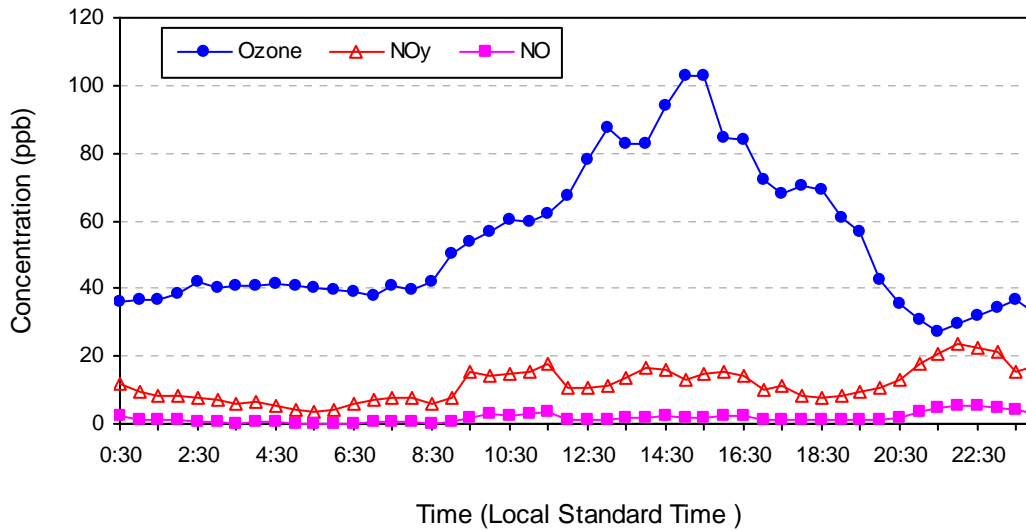
Parameter	Mean (ppb)	S.D. (ppb)	Range (ppb)	N
NO	2.6	1.2	1-7	54
NO <sub>y</sub>	37.3	10.0	21-69	54
NO <sub>x</sub> *	16.3	5.5	10-35	54
NO <sub>z</sub> *	21.0	5.7	9-33	54
O <sub>3</sub>	94.0	26.5	50-172	54

<sup>a</sup> Weekend and weekdays.

The range of  $\text{NO}_y$  concentrations measured at the Santa Ursula site (21-69 ppb) was significantly lower than the range of  $\text{NO}_y$  afternoon levels (10-240 ppb) reported by Sosa et al. (2000) for two sites in Mexico City close to downtown in March 1997. Although care should be taken when comparing the  $\text{NO}_y$  values reported by Sosa et al. (2000), because they based their analysis on a composite of measurements of  $\text{NO}_x$ , PAN,  $\text{HNO}_3$ , and  $\text{NO}_3^-$  that were not taken simultaneously, were not corrected for known interferences ( $\text{NO}_x$  measurements), and, in some cases, were not measured at the same site. However, the range of concurrent calculated  $\text{NO}_z^*$  for the Santa Ursula site (9-33 ppb) was very similar to the range of  $\text{NO}_z$  values (as the sum: PAN +  $\text{HNO}_3$  +  $\text{NO}_3^-$ ) reported by these authors (1-35 ppb). In comparison, the ranges of the measured  $\text{O}_3$  and  $\text{NO}_y$  concentrations and estimated  $\text{NO}_z^*$  values in Santa Ursula were below the ranges of the respective measurements reported for the Los Angeles Basin during August 1987. The Los Angeles measurements were ~ 20-240 ppb for  $\text{O}_3$ ; ~ 25-140 ppb for  $\text{NO}_y$ ; and ~ 2-55 ppb for  $\text{NO}_z$  (Sillman et al, 1997).

#### **4.2.2.2 ININ Site**

Although the monitoring campaign at the ININ site extended from April 1-20, 2004, effective measurements of  $\text{NO}_y$  were possible only from April 16-19. Before April 16, cold weather, cloudy conditions, and continuous rain kept the  $\text{NO}_y$  measurements at a very low level and without variation. There was evidence that the low temperatures (median ~ 9.5°C and lows ~ 2°C) affected the conversion efficiency of the external converter. This situation changed during the week of April 15 when micro-regional weather conditions changed and ambient temperatures rose, coinciding with several smog events in Mexico City. Figure 4.11 shows the hourly average diurnal profiles of  $\text{O}_3$ ,  $\text{NO}_y$ , and NO concentrations from April 16-19 at the ININ site. Table 4.7 presents a statistical summary of the measurements, including meteorological information for this period. According with Figure 4.11, nocturnal and early morning  $\text{O}_3$  levels remained almost at a background level (~ 40 ppb), due to the lack of sufficient NO concentrations to titrate the ozone concentrations.



**Figure 4.11.** Composite diurnal profiles of 30-min average  $O_3$ ,  $NO_y$ , and  $NO$  concentrations measured at the ININ site in southwest Mexico City, April 16 to 19, 2004.

$O_3$  production started after sunrise and was a result primarily of  $NO_x$  emitted in the nearby area. The average rate of  $O_3$  accumulation was  $\sim 7.2$  ppb/hr, once the ozone levels began to increase. The average time of  $O_3$  build-up was around 7 hrs and the average time with ozone concentrations above 80 ppb was 4 hrs. The  $O_3$  peaks occurred in the afternoon hours when the mixing layer reached its maximum development in region. The pass of air parcels coming from the urban area was registered with the occurrence of high ozone levels several hours after the maximum solar radiation.  $O_3$  reached its maximum around 17:00 hr when solar radiation was at a minimum.  $NO$  and  $NO_y$  concentrations were in general low and did not show the typical morning peak observed in urban areas. From  $\sim 9:00$  to  $\sim 17:00$  hr, the  $NO_y$  concentrations rose and fluctuated between 10 and 20 ppb. A peak in  $NO_y$  levels occurred after sunset in concert with a decrease in  $O_3$  concentrations. This effect could be a result of the movement of air parcels containing “fresh”  $NO_x$  from nearby locations. During the four-day smog-event period, one exceedence to the Mexican Ozone Air Quality Standard was observed. The  $NO$  levels did not equal the lowest  $O_3$  concentration along the entire monitoring period. The diurnal  $NO_y$  average was 12.1 ppb, which was almost 4.5 times lower than the same average for the Santa Ursula site.

**Table 4.7.** Statistical summary of daily 30-min average  $\text{NO}_x^{*(a)}$ ,  $\text{NO}_y$  and  $\text{O}_3$  concentrations and 1-hr average meteorological parameters measured at the ININ site in southwest Mexico City (April 16-19, 2004).

Parameter	Unit	Mean	S.D.	Range	No. of Samples
NO	ppb	1.6	1.6	0.01-9.54	151
$\text{NO}_x^{*(a)}$	ppb	10.7	6.69	1.0-41.5	155
$\text{NO}_y$	ppb	12.1	6.6	1.3-35.7	135
$\text{O}_3$	ppb	52.2	23.7	17.6-125.8	161
Temperature	°C	13.1	7.37	3.2-25.6	109
Wind Speed	m/sec	0.97	0.62	0-2.2	109
Wind Direction	degrees	Dominant: 202.5 (28.5%)		0-360	109

<sup>(a)</sup>  $\text{NO}_x$  values as measured by the instrument.

<sup>(b)</sup> 30-min average

Table 4.8 presents a basic statistical analysis of the 30-min average NO,  $\text{NO}_y$ , and  $\text{O}_3$  measured concentrations for the afternoon hours (13:00-17:00 hr) from April 16-19, 2004. This table also shows the estimated  $\text{NO}_x^*$  and  $\text{NO}_z^*$  values at the ININ site.

**Table 4.8.** Statistical summary of 30-min average concentrations of measured NO,  $\text{NO}_y$  and estimated  $\text{NO}_x^*$  and  $\text{NO}_z^*$  between 13:00-17:00 hr (April 16-19, 2004) at the ININ site in southwest Mexico City.

Parameter	Mean (ppb)	S.D. (ppb)	Range (ppb)	N
NO	1.9	0.85	0.56-3.5	25
$\text{NO}_y$	14.6	4.16	7.7-22.3	25
$\text{NO}_x^*$	9.67	2.88	5.0-15.0	25
$\text{NO}_z^*$	2.2	0.41	1.6-3.1	25
$\text{O}_3$	87.5	19.4	56.3-125.8	26

The  $\text{NO}_y$  average afternoon level was 2.5 times lower than the same average in Santa Ursula, and the average afternoon  $\text{NO}_z^*$  value was around 10 times lower than that at Santa Ursula. However, the average afternoon  $\text{O}_3$  concentration was very close to that of the Santa Ursula site for the same dates.

Several samples of gaseous  $\text{HNO}_3$  were collected with an annular denuder system at the ININ site. Table 4.9 presents the results of an analysis of the aqueous extracts from samples taken between April 13-19, 2004.

**Table 4.9.** Statistical summary of the analytical results for aqueous extracts of samples of  $\text{HNO}_3$  collected with an annular denuder system at the ININ site from April 13 to 19, 2004.

Date	Sampling period	Concentration ( $\mu\text{g/mL}$ )	Comment <sup>a</sup>
April 13	10:00-11:07 hr	0.0440	< Detection limit
April 13	11:00-14:13 hr	0.0080	< Detection limit
April 16	12:00-14:00 hr	0.0296	< Detection limit
April 16	9:54-11:54 hr	0.0133	< Detection limit
April 19	11:56-13:53 hr	0.0310	< Detection limit
April 19	Transport blank	0.0032	-

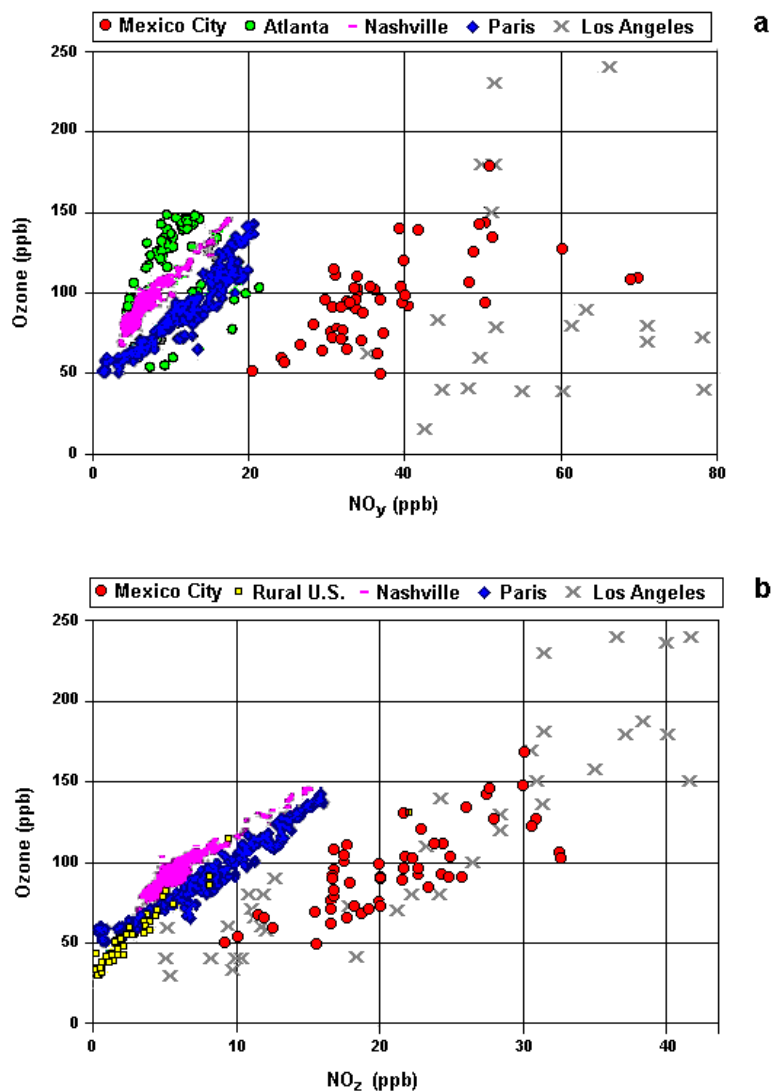
<sup>a</sup> Detection limit = 0.07  $\mu\text{g/mL}$

As shown in Table 4.9, however, the amount of  $\text{HNO}_3$  collected in all the samples was insufficient to reach the detection limit of the analytical method. The very low concentrations of  $\text{HNO}_3$  encountered at the ININ site suggested that mechanisms were removing the  $\text{HNO}_3$  from the air parcels before they reached this area.

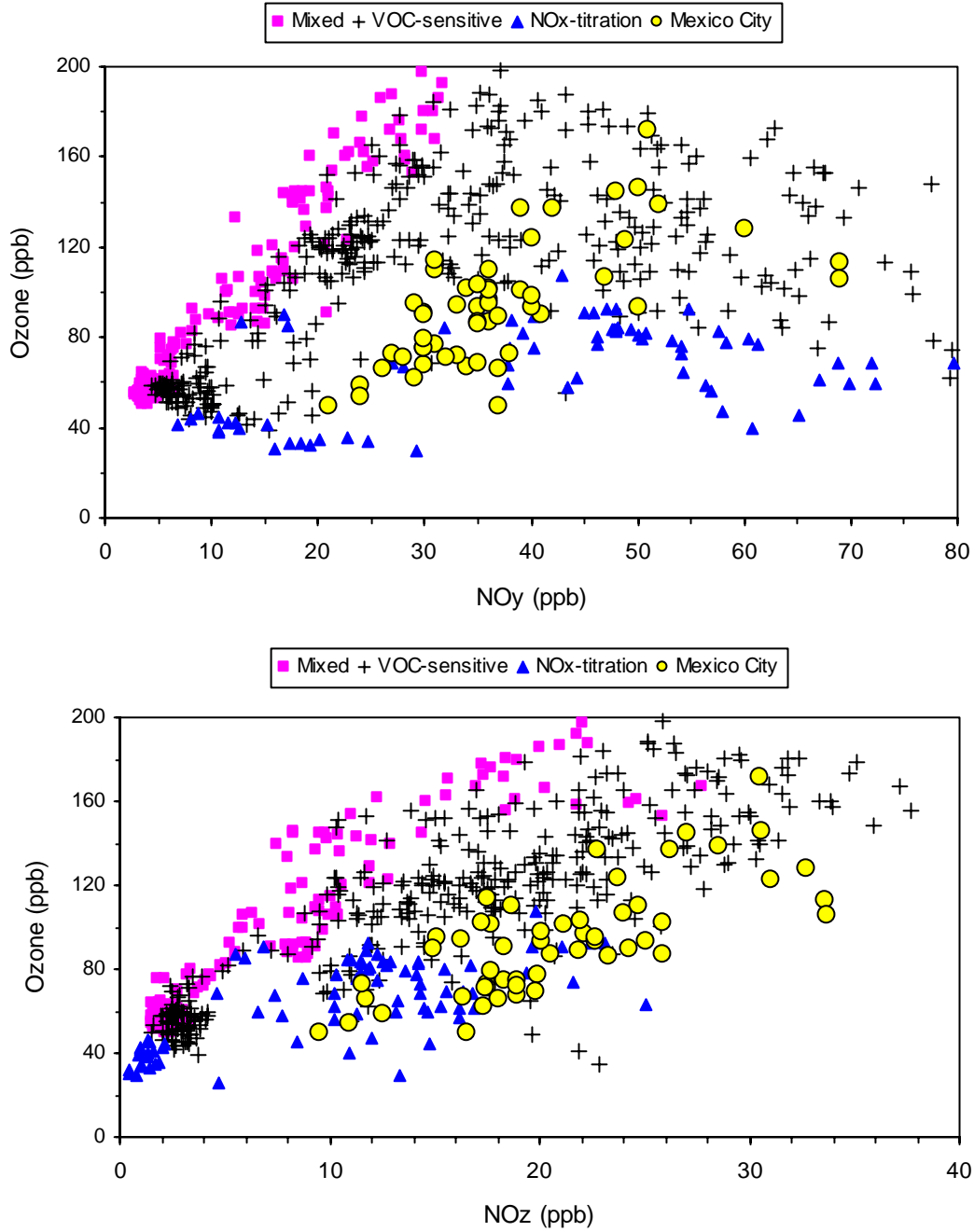
### 4.2.3 Photochemical Indicators Evaluation

#### 4.2.3.1 Santa Ursula Monitoring Site

Figure 4.12 shows the comparison between measured afternoon hourly averages of  $\text{O}_3$  and  $\text{NO}_y$  and measured  $\text{O}_3$  and estimated  $\text{NO}_z^*$  in Santa Ursula and these parameters reported at several sites in the U.S. and in Paris, France (Sillman 2002). Figure 4.12 demonstrates that the measured data points from Santa Ursula matched the pattern of measured data from Los Angeles. Figure 4.13 presents a comparison of the scatterplot patterns of measured afternoon  $\text{O}_3$ , measured  $\text{NO}_y$ , and estimated  $\text{NO}_z^*$  at Santa Ursula with a composite of scatterplot patterns of predicted  $\text{O}_3$  versus  $\text{NO}_y$  and  $\text{NO}_z$  afternoon values for mixed-, VOC-, and  $\text{NO}_x$ -titration conditions for several regions in the U.S. (Sillman and He, 2002).



**Figure 4.12.** Scatterplot patterns of: **(a)** measured  $\text{O}_3$  and  $\text{NO}_y$ , and, **(b)** measured  $\text{O}_3$  and estimated  $\text{NO}_z^*$  at Santa Ursula, Mexico City, compared with measured correlations of  $\text{O}_3$  and  $\text{NO}_y$  and  $\text{NO}_z$  patterns for different locations identified as:  $\text{NO}_x$ -sensitive (Atlanta and Rural U.S.), mixed sensitive (Nashville and Paris), and VOC-sensitive (Los Angeles). Scatterplots adapted from Sillman (2002 and 2003).



**Figure 4.13.** Scatterplot patterns of: **(a)** measured O<sub>3</sub> and NO<sub>y</sub>, and **(b)** measured O<sub>3</sub> and estimated NO<sub>z</sub>\* at Santa Ursula, Mexico City, compared with 3-D model predicted O<sub>3</sub> and NO<sub>y</sub> and NO<sub>z</sub> patterns, respectively, for locations with mixed or with near-zero sensitive conditions (squares), VOC-sensitive conditions (crosses), and dominated by NO<sub>x</sub>-sensitive titration (triangles). Simulation data obtained from data file available from Sillman (2003).

According to Figure 4.13(a), the  $O_3$  versus  $NO_y$  scatterplot correlation for the Santa Ursula data practically superimposed one important segment of the respective simulated correlation for VOC-conditions. However, some data points of the  $O_3$  versus estimated  $NO_z$  data points for Santa Ursula, seen in Figure 4.13(b), overlapped with the pattern for simulated conditions dominated by  $NO_x$ -titration.

Following the procedure used to clarify this same contradiction observed in the data for Cincinnati, Figure 4.14 shows the scatterplot correlation for  $O_3$  versus the ratio  $NO_x^*/NO_y$  for Santa Ursula superimposed on the simulated  $O_3$  versus  $NO_x/NO_y$  ratios for VOC-sensitive and  $NO_x$ -titration conditions. This figure clearly shows that the measured indicator species at Santa Ursula were strongly associated with VOC-sensitive conditions, as predicted.

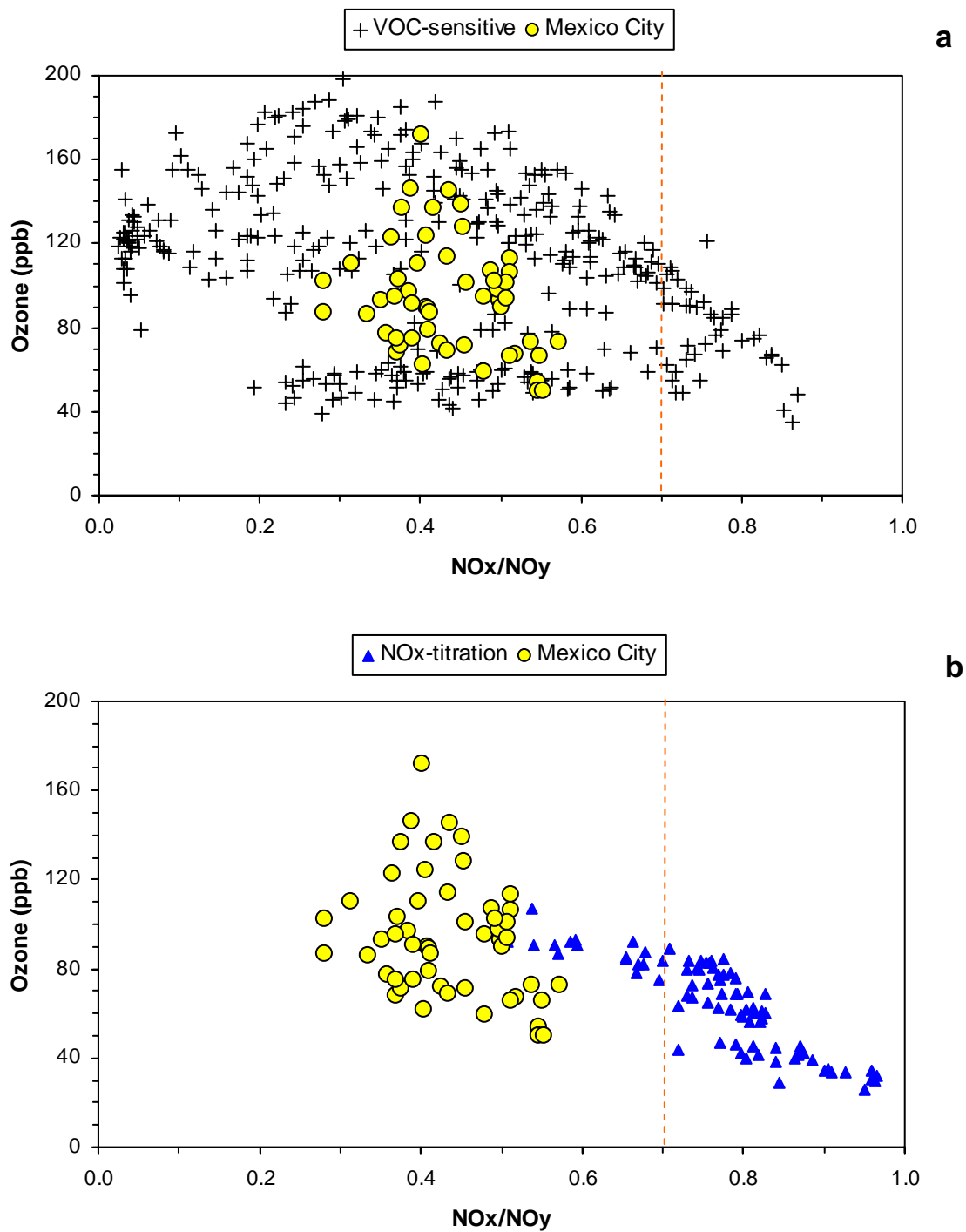
The estimated afternoon average  $NO_x^*/NO_y$  ratio for the Santa Ursula data was  $\sim 0.44$ , suggesting that the air masses that passed over this site during the monitoring campaign were “middle-aged” parcels. Table 4.10 shows the comparison between the calculated average indicator ratios and the average  $NO_y$  concentration from the measurements carried out at Santa Ursula and the transition values suggested by Sillman (1995) and Sillman and He (2002) for moderately  $O_3$  polluted sites. According to this table, the three indicator ratios  $O_3/NO_y$ ,  $O_3/NO_z^*$ , and  $NO_y$  pointed to VOC-sensitive conditions.

Figure 4.15 presents the scatterplot of measured afternoon  $O_3$  versus concurrent  $NO_y$  and of measured  $O_3$  versus concurrent estimated  $NO_z^*$  at Santa Ursula, compared with the mean transition lines for  $NO_x$ - and VOC-sensitive conditions for moderately polluted conditions ( $80 \text{ ppb} < [O_3] < 200 \text{ ppb}$ ).

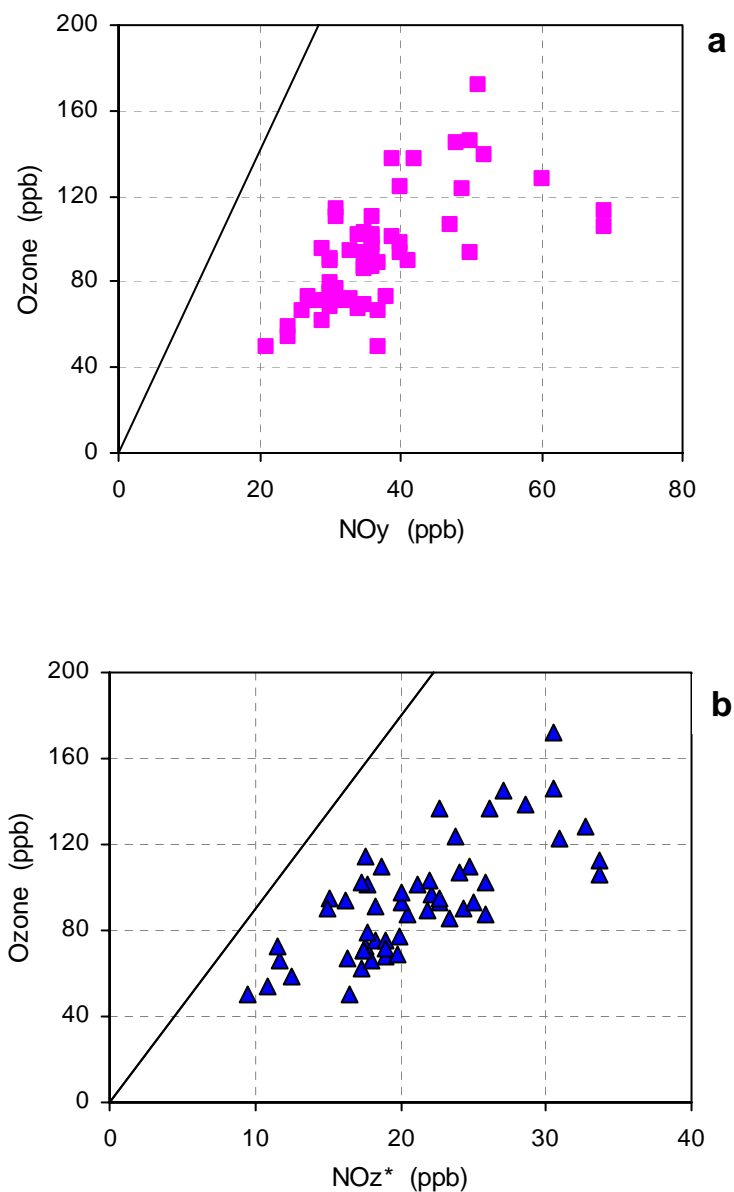
**Table 4.10.** Calculated average indicator ratios observed at the Santa Ursula Monitoring Station (Mexico City) and transition values for moderately ozone (80 to 200 ppb) polluted conditions according to Sillman (1995) and Sillman and He (2002). Values of ratios lower than their transition ranges indicate VOC-sensitive conditions. Values of  $NO_y$  higher than the transition value indicate VOC-sensitive conditions.

Indicator	Santa Ursula average ratios	Transition range
$O_3/NO_y$	2.5	6-8
$O_3/NO_z$	4.47	8-10
$NO_y$	37.3	20





**Figure 4.14.** Scatterplot patterns for afternoon  $\text{O}_3$  versus  $\text{NO}_x^*/\text{NO}_y$  data for Santa Ursula, Mexico City, compared with the calculated  $\text{NO}_x/\text{NO}_y$  ratio from data results of 3-D model simulations for locations with: (a) VOC-sensitive conditions, and (b) dominated by  $\text{NO}_x$ -titration. The vertical dashed line represents the hypothetical limit to differentiate photochemically aged air parcels. Simulation data obtained from data file available from Sillman (2003).



**Figure 4.15.** Scatterplot patterns of: **(a)** measured  $O_3$  and concurrent  $NO_y$ , and **(b)** measured  $O_3$  and concurrent estimated  $NO_z^*$  from 13:00 to 17:00 hr, April 14-25, 2003 at Santa Ursula, Mexico City. The lines represent the transition between  $NO_x$ - and VOC-sensitive chemistry for  $O_3/NO_y$  and  $O_3/NO_z$  indicator ratios for moderately  $O_3$  conditions ( $80 \text{ ppb} \leq [O_3] \leq 200 \text{ ppb}$ ) according to Sillman and He (2002).

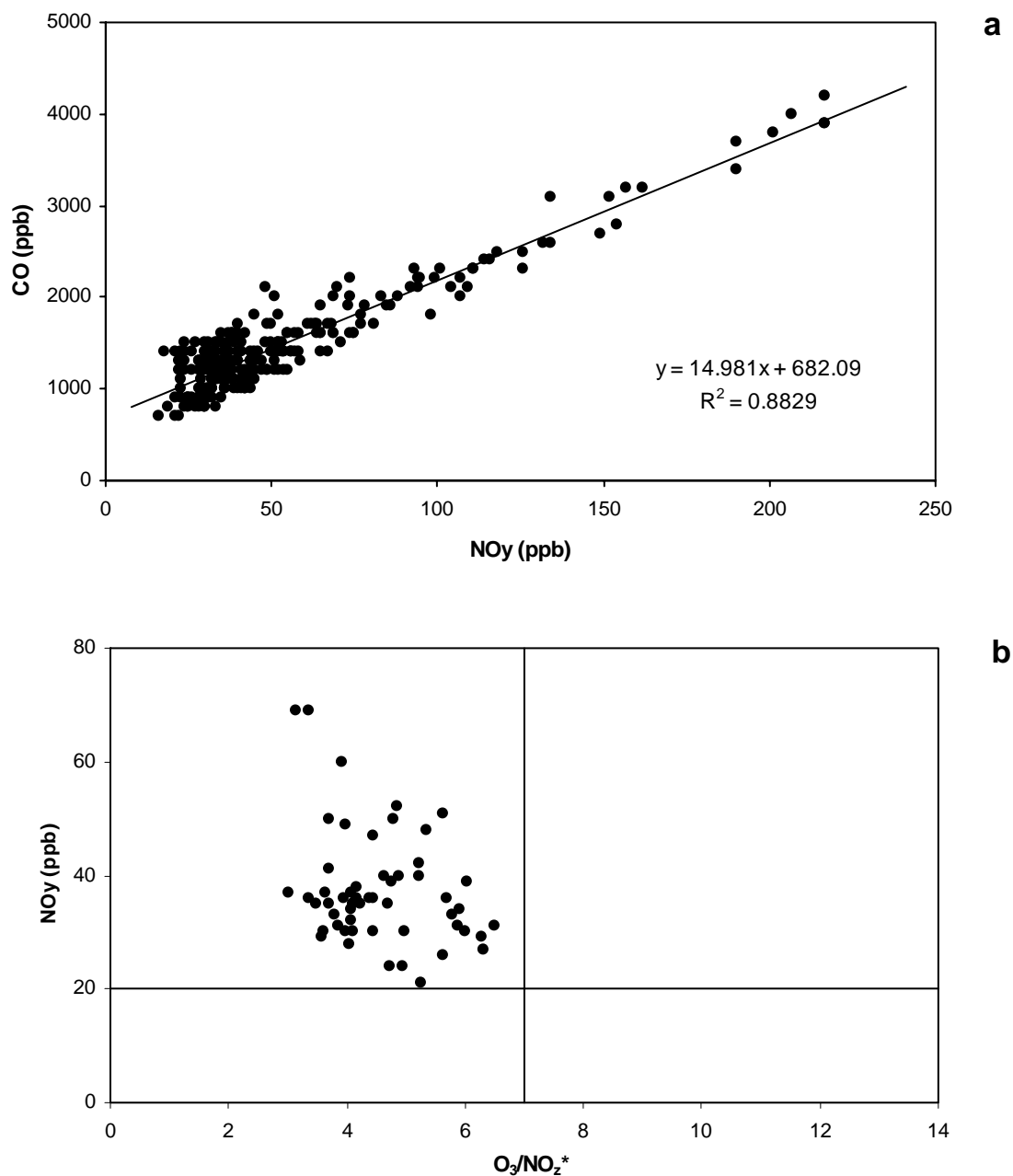
The evaluation of the consistency of the  $\text{NO}_y$  measurements through the correlation between  $\text{O}_3$  and the sum ( $2\text{H}_2\text{O}_2 + \text{NO}_z$ ) could not be performed in this site because  $\text{H}_2\text{O}_2$  measurements were not available. Two indirect alternative approaches were used for this evaluation. The first consisted in the regression analysis of CO concurrent with the  $\text{NO}_y$  monitoring data. As previously indicated, an acceptable correlation coefficient would be indicative that  $\text{NO}_y$  measurements were reliable. The second approach consisted in the empirical scatterplot comparison of the  $\text{NO}_y$  values versus the  $\text{O}_3/\text{NO}_z^*$  ratio. The criterion applied was that, if the suggested sensitivities shown simultaneously for the indicators  $\text{NO}_y$  and  $\text{O}_3/\text{NO}_z^*$  were consistent among them, that would be an indication that the  $\text{NO}_z^*$  values were within reasonable values.

Figure 4.16 shows the correlation between hourly measurements of CO and  $\text{NO}_y$  at SUMS from April 14-25, 2004 and the  $\text{NO}_y$ - $\text{O}_3/\text{NO}_z^*$  correlation for the same data period. As shown in Figure 4.16(a) the correlation between CO and  $\text{NO}_y$  concentrations resulted in an  $R^2 = 0.88$ . The losses in some  $\text{NO}_y$  species by deposition mechanisms could be associated with the withdrawal of the data points from the ideal correlation.

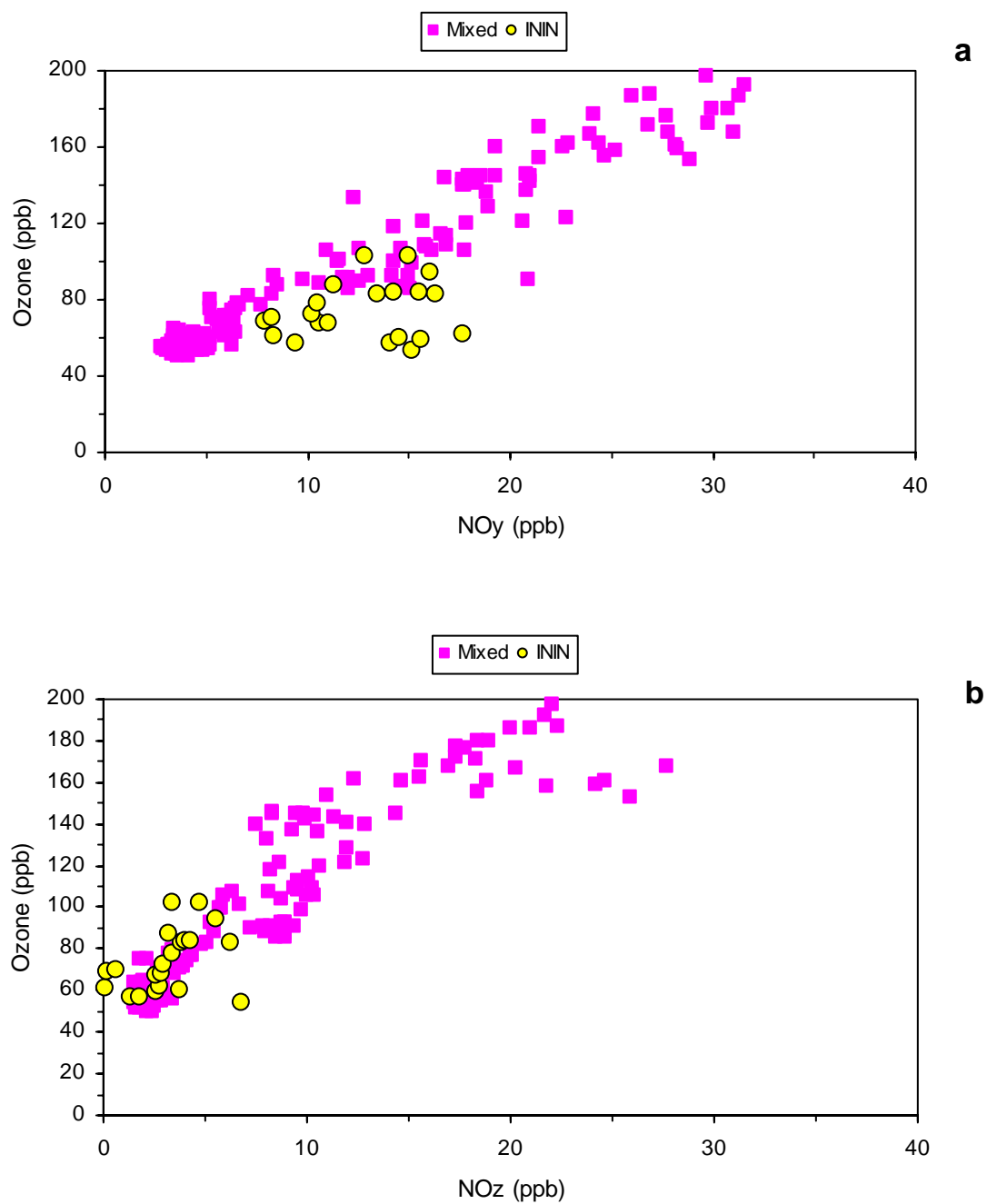
In Figure 4.16(b), all the data points of the  $\text{NO}_y$ - $\text{O}_3/\text{NO}_z^*$  correlation fell in the upper left quadrant ( $\text{NO}_y > 20$  ppb and  $\text{O}_3/\text{NO}_z^* < 7$  for moderate ozone conditions). The measured indicators were consistent with each other simultaneously.

#### 4.2.3.2 ININ Site

Figure 4.17 shows the correlations patterns of measured  $\text{O}_3$  ( $\geq 50$  ppb) versus measured  $\text{NO}_y$  and estimated  $\text{NO}_z^*$  for the ININ site with 3-D model predicted  $\text{O}_3$ ,  $\text{NO}_y$ , and  $\text{NO}_z$  patterns, respectively. The model predictions are for locations with mixed or with near-zero sensitive conditions. Figure 4.17 reveals a contradiction in the PIM theory. While the correlation of  $\text{O}_3$  versus  $\text{NO}_y$  suggests that VOC-sensitive conditions were prevalent at the ININ site, the correlation of  $\text{O}_3$  versus  $\text{NO}_z^*$  indicates that relatively strong  $\text{NO}_x$ -sensitive conditions could dominate this site during smog-events.



**Figure 4.16.** Scatterplot patterns between: **(a)** diurnal CO and concurrent NO<sub>y</sub> hourly concentrations, and, **(b)** correlation between afternoon NO<sub>y</sub> concentrations and O<sub>3</sub>/NO<sub>z</sub>\* ratios for measurements carried out at the Santa Ursula Monitoring Station from April 14 to 25, 2004.



**Figure 4.17.** Scatterplot patterns of: **(a)** measured  $O_3$  ( $\geq 50$  ppb) and  $NO_y$ , and **(b)** measured  $O_3$  ( $\geq 50$  ppb) and estimated  $NO_z^*$  at the ININ site south west Mexico City, compared with 3-D model predicted  $O_3$  and  $NO_y$  and  $NO_z$  patterns, respectively, for locations with mixed or with near-zero sensitive conditions. Simulation data obtained from data file available from Sillman (2003).

This contradiction might be explained by analyzing the pattern of  $O_3$  versus  $NO_z^*$ . The narrow range of  $NO_z^*$  concentrations shown in Figure 4.17 suggests that the air masses that passed over the site most likely suffered the prior removal of species such as  $HNO_3$  and  $NO_3^-$  aerosols. Otherwise the  $NO_z^*$  values should match the mixed-sensitive conditions but not those of the  $NO_x$ -sensitive region.

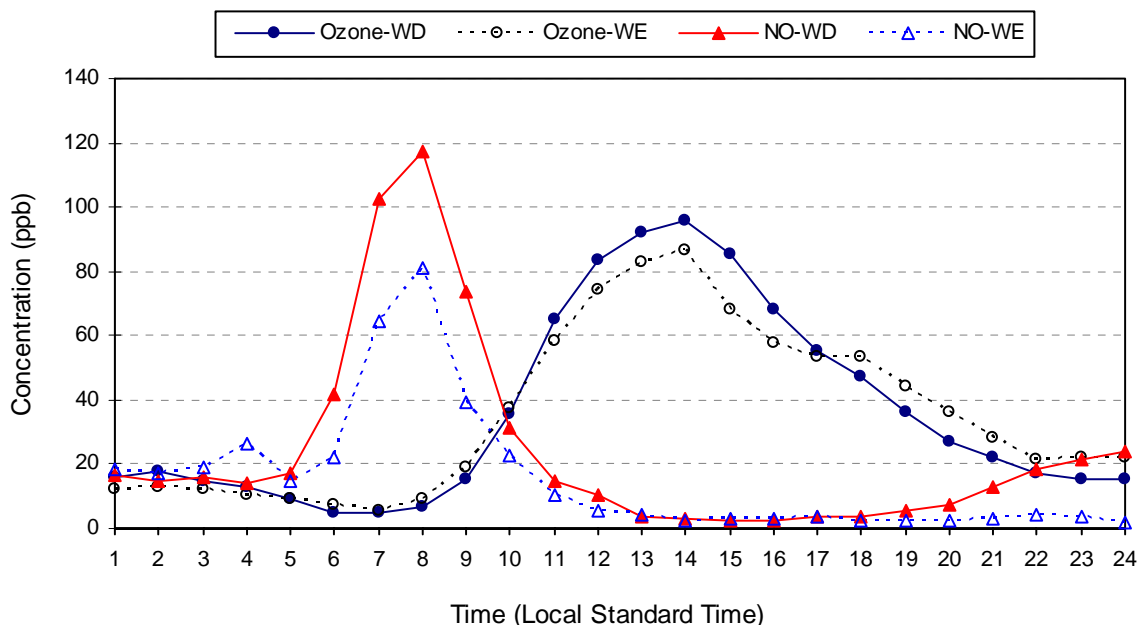
It is also possible that the complete opposite occurred. That is, the heated Mo converter could have a lower conversion efficiency to  $HNO_3$  and PAN species present in the sample, leading to the measurement of relatively smaller  $NO_y$  concentrations and, therefore, to smaller differences in the subtraction of  $NO_y - NO_x^*$ . Singh (1987) suggested that, like ozone, PAN is continually synthesized during the transport of air masses and that PAN acts as both carrier and reservoir of  $NO_2$  due to the thermal instability of its molecule [ $CH_3C(O)-O-O + NO_2 \leftrightarrow CH_3(O)-O-O-NO_2$ ].

As a result, it is possible that the analyzer had measured levels of both PAN and  $NO_2$  in the “ $NO_y$  mode” very similar to those detected in the  $NO_x$  mode. Due to this uncertainty in the  $NO_y$  measurements and, thus, in the estimated  $NO_z^*$  levels, no preliminary conclusions were drawn from the experimental results at this site.

#### **4.2.4 Weekend/Weekday Effect Analysis**

Although the monitoring period for  $NO_y$  at SUMS extended from April 14-25, the weekend/weekday (WE/WD) analysis was performed on the entire data set of hourly averages of  $O_3$  and NO concentrations registered from April 1-30. Figure 4.18 shows the average hourly diurnal WE/WD variations of  $O_3$  and NO for April 2004 at SUMS. Table 4.11 presents the results of the tests for statistical significance in WE and WD for 1-hr and 8-hr average  $O_3$  maximum and 6:00-9:00 hr NO average concentrations for the same sampling period.

According to Figure 4.18, the hourly average  $O_3$  WE and WD daily patterns were very similar. The lowest average ozone concentrations occurred in the morning hours and the peaks around 14:00 hr. The 6:00-9:00 average WD  $O_3$  concentration was slightly lower during morning rush hours than on the WE, probably as a result of  $O_3$  titration by NO emitted from heavy weekday traffic.



**Figure 4.18.** Average weekend and weekday hourly diurnal variations of  $O_3$  and  $NO$  concentrations measured at the Santa Ursula monitoring station, Mexico City from April 1-30, 2004.

However, the average of maximum WD  $O_3$  concentrations was around 10 ppb higher than that on the WE. Immediately after the occurrence of the peak, a relatively constant decrease in  $O_3$  levels was observed in both cases.

In general, the  $O_3$  WE and WD daily patterns in Santa Ursula showed a correspondence with the typical amount of emissions of ozone precursors during these days. Reduced diesel truck and gasoline vehicle traffic in WE could be associated with the reduction in  $NO_x$  emissions leading to a relatively short supply of  $NO$  for morning titration. However, the average  $NO$ - $O_3$  crossover time in WE was only about 30-min earlier than that in WD. The duration of the build-up and accumulation of  $O_3$  was also slightly higher in WE ( $\sim 4.3$  hr) than in WD ( $\sim 4$  hr). The average WE/WD rate of accumulation of  $O_3$  was also different. It was somewhat faster in WD (15.3 ppb/hr) than in WE (13 ppb/hr). This apparent inverse correspondence could be related to a proportional reduction in both anthropogenic VOC emissions and  $NO_x$  in WE.

**Table 4.11.** Results of the tests for statistical significance for WE/WD differences of maxima 1-hr and 8-hr average O<sub>3</sub> concentrations and 6:00-9:00 hr. NO average levels for data registered between April 1 through 30, 2004, at the Santa Ursula monitoring station, Mexico City.

Parameter	Mean (S.D.) ppb	No. of samples	Results of tests of significance	Interpretation
WE max 1-hr O <sub>3</sub>	86.6 (40.5)	8	$p = 0.559$	$H_0^b$ : $\mu_{1\text{-hr O}_3\text{WE}} = \mu_{1\text{-hr O}_3\text{WD}}$ Accepted
WD max 1-hr O <sub>3</sub>	96.1 (37.8)	21		$H_I^c$ : $\mu_{1\text{-hr O}_3\text{WE}} < \mu_{1\text{-hr O}_3\text{WD}}$ Rejected
WE max 8-hr O <sub>3</sub>	74.0 (22.8)	8	$p = 0.505$	$H_0^b$ : $\mu_{8\text{-hr O}_3\text{WE}} = \mu_{8\text{-hr O}_3\text{WD}}$ Accepted
WD max 8-hr O <sub>3</sub>	66.8(29.8)	17		$H_I^c$ : $\mu_{8\text{-hr O}_3\text{WE}} < \mu_{8\text{-hr O}_3\text{WD}}$ Rejected
WE 6-9 a.m. NO	61.5 (59.7)	56	$p = 0.0094$	$H_0^d$ : $\mu_{6\text{-}9\text{am NO WE}} = \mu_{6\text{-}9\text{am NO WD}}$ Rejected
WD 6-9 a.m. NO	97.5 (43.7)	24		$H_I^e$ : $\mu_{6\text{-}9\text{am NO WE}} > \mu_{6\text{-}9\text{am NO WD}}$ Accepted

<sup>a</sup> If  $p \leq 0.05$ , the difference is significant; if  $0.05 \leq p \leq 0.10$  the difference between the means is marginally significant, if  $p > 0.10$ , the difference is insignificant.

<sup>b</sup>  $H_0$ : The change in WE maximum O<sub>3</sub> is insignificant, likely to be caused by a VOC-sensitive condition provided that the reduction in morning WE-NO<sub>x</sub> emissions is significant.

<sup>c</sup>  $H_I$ : The decrease in maximum WE maximum O<sub>3</sub> is significant, likely due to a NO<sub>x</sub>-sensitive condition on condition that morning WE-NO<sub>x</sub> emissions is significant.

<sup>d</sup>  $H_0$ : The reduction in morning WE-NO<sub>x</sub> emissions is insignificant probably due to mere chance.

<sup>e</sup>  $H_I$ : The reduction in morning WE-NO<sub>x</sub> is significant likely due to the reduction in WE activities.

The decreased availability of NO in WE could allow O<sub>3</sub> increased time for accumulation, but the relatively small supply of reactive VOCs in WE could not support a faster rate of O<sub>3</sub> build-up. The combination of both effects could lead to the lower O<sub>3</sub> peak in WE. However, the statistical differences between these peaks (see Table 4.11) show that the difference was not statistically significant at a 95% confidence level. In addition, a stricter two-tail *t-test* with unequal variances for the WE/WD average max 1-hr O<sub>3</sub> as well as the testing of the differences in the average max 1-hr O<sub>3</sub> WE and WD means from a matrix of combinations of WE/WD 2 days (equal variances) were performed. In all cases the results of the



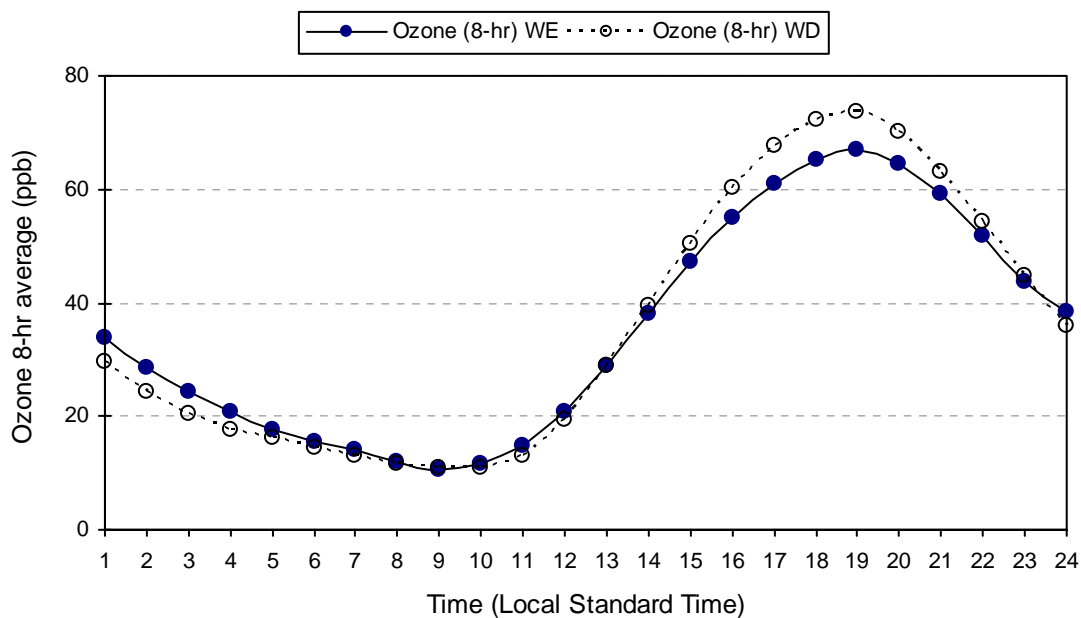
tests were the same as those shown in Table 4.4. Appendix D shows the summary of the results for the statistical analysis for the set of two WE/WD combinations.

On the other hand, Figure 4.18 shows that NO WE/WD hourly average profiles were different in the morning and evening hours. Although morning WE and WD NO peaks occurred at the same hour (~ 8:00 hr), the WD NO peak was almost 1.4 times higher than the morning WE NO peak. The suggested association between the reduction of NO<sub>x</sub> emissions and reduction of traffic during WE seems to be confirmed by this difference. Furthermore, the test of significance of the difference in means of the 6:00-9:00 hr WE NO average (see Table 4.11) showed that this reduction was statistically significant at a 95% confidence level.

After the occurrence of the NO morning peaks, the diurnal WE and WD NO average profiles decreased to reach almost the same level (~ 2-5 ppb) between 13:00 and 18:00 hr. This pattern changed around 19:00 hr in WD and was probably associated with evening traffic, which is the most important source of fresh NO<sub>x</sub> emissions.

WD evening and night rush hours led to a significant increase in the average NO levels compared with the WE NO average concentrations for the same hours, in accordance with decreased WE nighttime traffic in Mexico City. After midnight, the reconstruction of the shallow mixing layer had an effect on the accumulation of NO in both WE and WD. Based on the results of Table 4.11, on the non-significant difference in WE/WD maximum 1-hr O<sub>3</sub> average concentrations, and on the positive significant difference between WE/WD 6:00-9:00 hr NO average levels, it is possible to suggest that a VOC-sensitive conditions dominated.

The O<sub>3</sub> 8-hr average diurnal WE/WD patterns (see Figure 4.19) for the Santa Ursula site followed the same shape as those for the O<sub>3</sub> 1-hr average patterns. In both, somewhat higher levels in WD are seen for almost the middle of the day. The largest difference (~ 6 ppb) occurred at 11:00 hr. The WD O<sub>3</sub> 8-hr average peak occurred at 18:00 hr with a difference of ~ 7 ppb. Test results for the WE/WD O<sub>3</sub> 8-hr average peaks showed that there was not a statistically significant difference between the peaks.

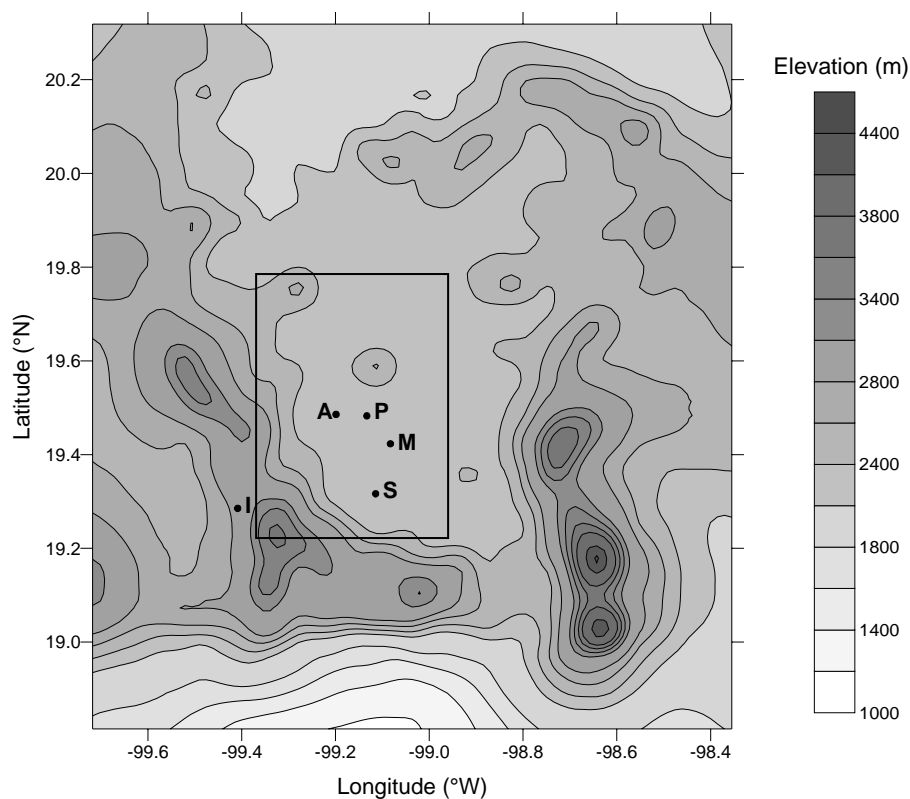


**Figure 4.19.** Average 8-hr O<sub>3</sub> weekend and weekday diurnal profile from April 1-30, 2004, at the Santa Ursula monitoring station, Mexico City.

#### 4.2.5 Modeling Experiment

The Multiscale Climatic and Chemistry Model (MCCM) was applied for three nested domains. The first domain, with a horizontal resolution of 27 km (60 x 60 cell mesh), practically covered the whole of Mexico. The second domain, with a 9 km resolution, (34 x 34 cell mesh) covered the central part of Mexico. The third domain had a resolution of 3 km (40 x 52 cell mesh) and included the Mexico City Metropolitan Area (MCMA), the surrounding mountains, and parts of the neighboring states of Mexico, Morelos, and Puebla (see Figure 4.20).

Meteorological simulations were performed only for Domain 1 and 2, due to lack of emissions data for these areas. For Domain 3, both meteorology and chemistry were simulated. The simulation's vertical domain included several layers, varying from 20 m near the ground to about 15 km above the surface (100 mb). To initialize and to provide boundary values for MCCM, data from the National Center for Environmental Protection's historical archives were used with a time resolution of 6 hr.

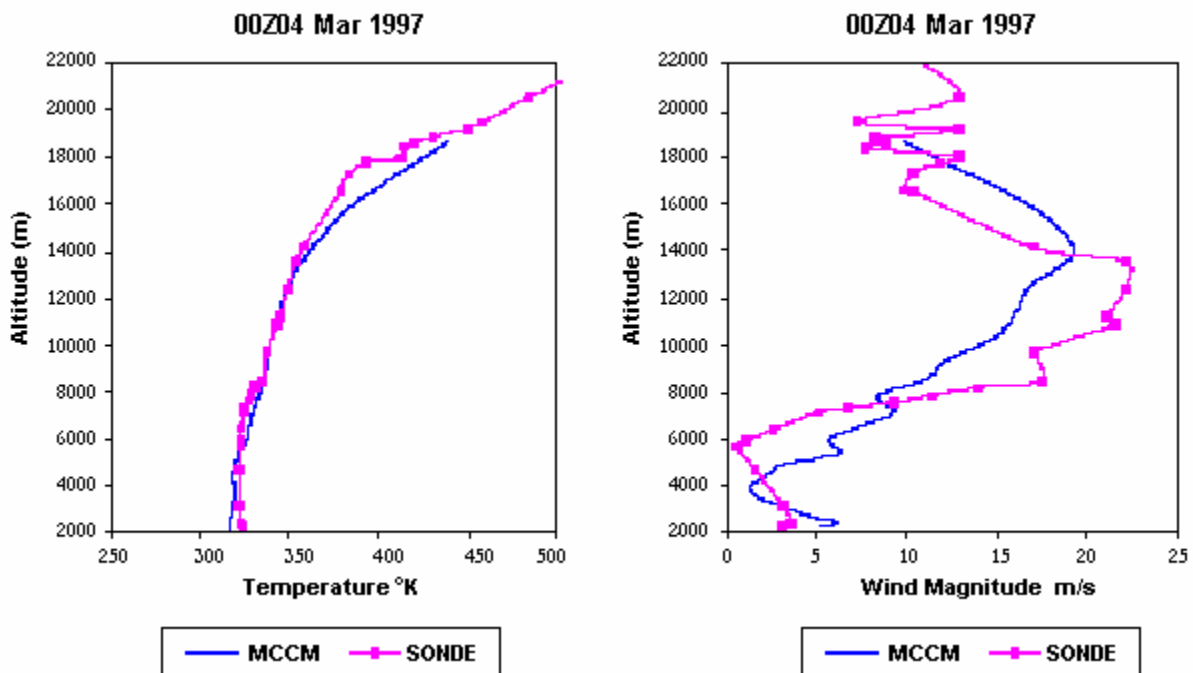


**Figure 4.20.** Topography of the innermost model domain (Domain 3) and position of five reference sites in MCMA. The reference sites are located in the downtown (Merced [**M**]), in the northwest (Azcapotzalco [**A**]), in the north (IMP [**P**]), in the south (Santa Ursula [**S**]), and in the southwest (ININ [**I**]) of Mexico City. The extent of the Domain 3 is around 80 km in East-West direction and around 75 km in the North-South direction. The rectangle in the center shows the approximately extent of the MCMA.

The first 24 hours of modeling were used for initialization only. Boundary concentration values for this domain were obtained from Domain 2, whose boundaries were set to background concentrations. The time period selected for the computational experiments was March 2-5, 1997. This period corresponded to the entrance of a high-pressure anti-cyclonic system over the central region of Mexico, which provided favorable conditions for a typical high pollution scenario (Doran et al., 1998; García, 2002). The simulations were performed at surface level.

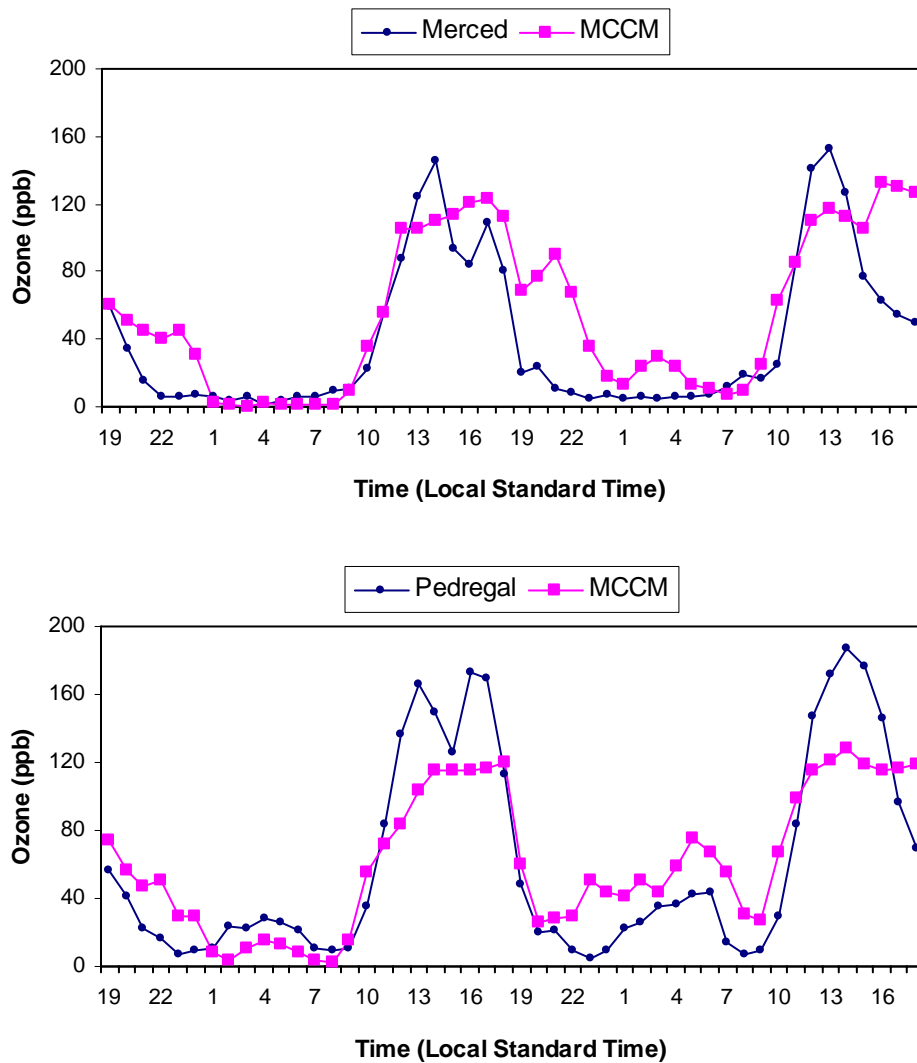
The event chosen for simulation with the MCCM was 14:00-15:00 hr on March 3, 1997. The species considered for the analysis were:  $O_3$ ,  $NO_y$  (as the sum of the simulated  $NO$ ,  $NO_2$ , PAN,  $HNO_3$ , and  $NO_3^-$  species),  $NO_z$  (as the subtraction of  $NO_y - NO - NO_2$ ),  $HNO_3$ , and  $H_2O_2$ . The performance of the

MCCM has been evaluated qualitatively and quantitatively for the Mexico City case (García, 2002; Jazcilevich et al., 2003). Simulated meteorological parameters have resulted in a satisfactory statistical index in agreement with measured data. Figure 4.21 shows the vertical profile of measured and modeled temperature and wind intensity at 18:00 hr on March 4, 1997, as an example of the MCCM's ability to reproduce vertical meteorological conditions.



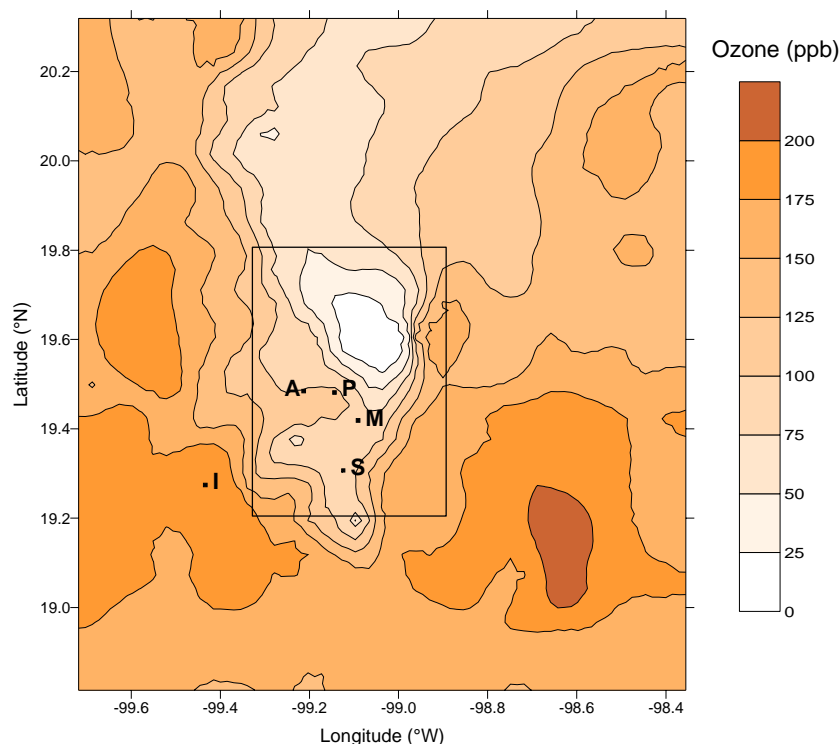
**Figure 4.21.** Vertical profiles comparing measured and modeled temperatures and wind velocity at the International Mexico City Airport at 18 LST March 3, 1997. (From García, 2002).

The accuracy of ozone predictions with the MCCM has been evaluated previously. Jazcilevich et al. (2003) found that the statistical index agreement between modeled and measured ozone values was above 0.7 with an overall average value of 0.78. The index has an optimum value of 1.0 for perfect agreement between model and measurements. Figure 4.22 shows the time series comparing measured and modeled  $O_3$  concentrations for the Merced (downtown) and the Pedregal (southwest) monitoring stations. The Pedregal station is located ~ 5 km to the east of the Santa Ursula site.



**Figure 4.22.** Time series of 1-hr average O<sub>3</sub> concentrations for 60 hours (March 2-4, 1997) using MCCM and two reference monitoring stations in ppb.

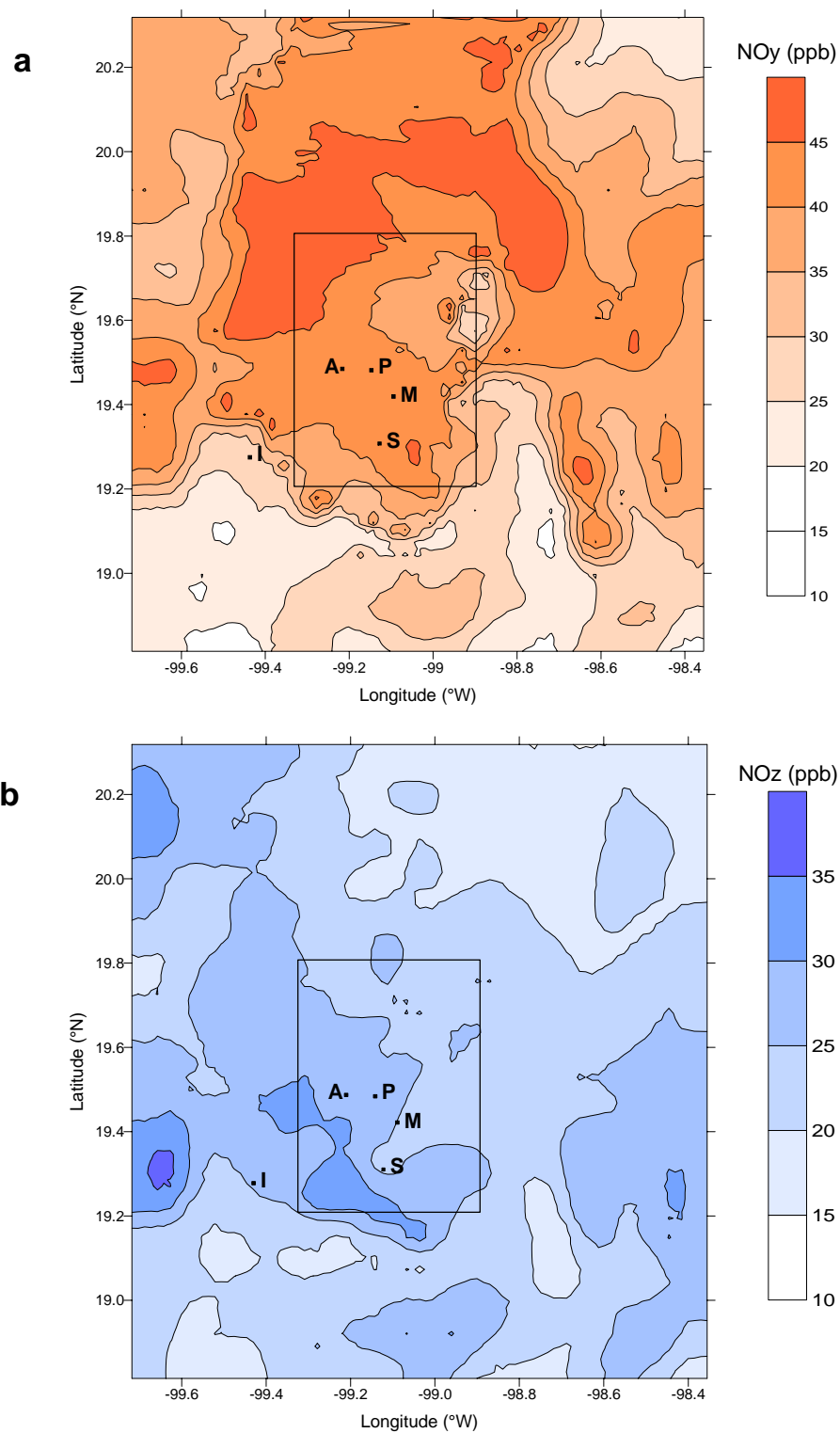
Figure 4.23 shows the spatial distribution of the simulated concentrations of O<sub>3</sub> over Domain 3. According to this figure, simulated O<sub>3</sub> concentrations were distributed across the whole Mexico City air basin and were in good agreement with observed afternoon ozone concentrations registered that day. High O<sub>3</sub> concentrations were found in the southwest and in the mountain range.



**Figure 4.23.** Simulated ozone concentrations distribution in ppb for MCMA air basin at the level surface at 15:00 hr Local Standard Time, March 3, 1997. The rectangle in the center shows the approximately extent of the urban area.

The simulated high  $O_3$  levels over the elevations of the basin could be associated with photochemical reactions between reactive biogenic hydrocarbons from the forests of this region and small amounts of  $NO_x$  transported from the city. The excess of OH radicals available for photochemical reactions could easily accelerate the formation and accumulation of  $O_3$ .

The lowest  $O_3$  concentrations were found northeast of the urban area downwind from a power plant there located. Fresh  $NO_x$  emissions from power plants can create a dominant  $NO_x$ -titration environment along the direction of their plumes. Within these plumes any available  $O_3$  can be easily titrated. However, the nearly stagnant synoptic meteorological conditions of this event-period were able to support the transport and accumulation of relatively high ozone levels over the northeast and west areas of the MCMA. Figure 4.24 shows the spatial distribution of the simulated concentrations of  $NO_y$  and  $NO_z$ .



**Figure 4.24.** Simulated (a)  $\text{NO}_y$ , and (b)  $\text{NO}_z$  concentrations distributions in ppb, for the MCMA air basin at the level surface at 15:00 hr Local Standard Time, March 3, 1997. The rectangle in the center shows the approximately extent of the urban area.

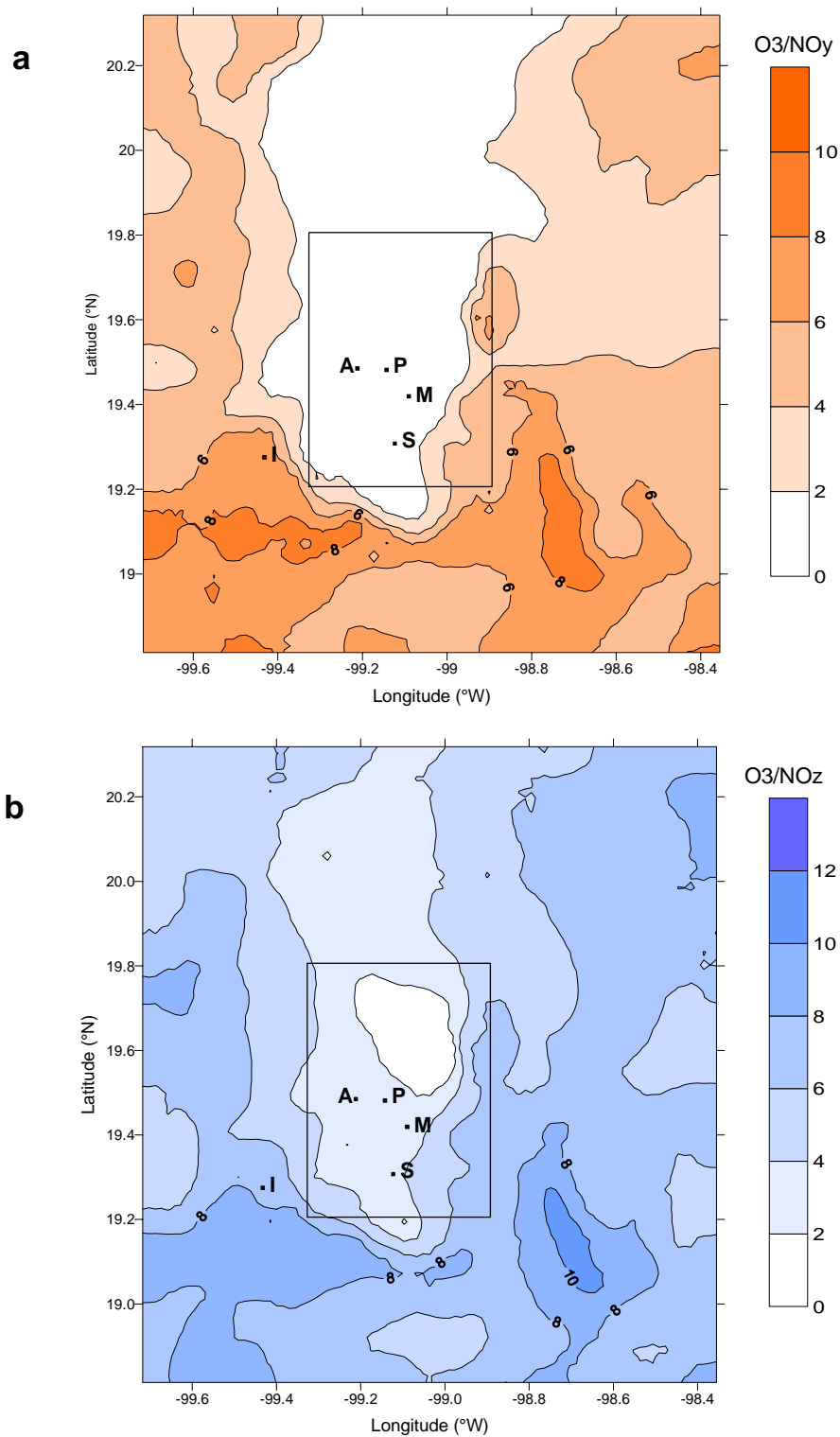
Figure 4.24(a) shows that the plume of  $\text{NO}_y$  species over the city followed a decreasing gradient towards the southwest, following the typical afternoon wind pattern.  $\text{NO}_y$  concentrations higher to 20 ppb were simulated for practically all the MCMA indicating that higher precursors concentrations were associated to this event (Milford et al., 1994; Sillman, 1995). Figure 4.24(b) shows that  $\text{NO}_z$  species increased slightly towards the elevations of the southwest as expected.  $\text{NO}_z$  species are secondary products of the oxidation of  $\text{NO}_x$  emissions. The relatively low  $\text{NO}_z$  concentrations over the northeast of the simulation domain reinforce the above hypothesis that  $\text{NO}_x$  species were more abundant in these areas.

Figures 4.25(a) and (b) show the spatial distribution of the simulated  $\text{O}_3/\text{NO}_y$  and  $\text{O}_3/\text{NO}_z$  ratios, respectively. The contour of the band, in Figure 4.25(a), associated with the transition range between the VOC- and  $\text{NO}_x$ -sensitive conditions (6-8) suggests that most of the  $\text{O}_3$  formed over MCMA did so under VOC-sensitive conditions ( $\text{O}_3/\text{NO}_y \leq 6$ ). The  $\text{NO}_x$ -sensitive areas were located mainly over the southern sector of the simulation domain in agreement with what is typically observed in urban and rural regions of the U.S.A. (Sillman, 1999). The  $\text{O}_3/\text{NO}_y$  ratios within the urban area resulted in an increasing gradient towards the southwest of the city. However, this gradient rapidly changed to  $\text{NO}_x$ -sensitive ratios as the plume left the outskirts of the MCMA. Theoretically, one explanation for this is that, as the polluted air mass ages, its chemistry tends to shift to a  $\text{NO}_x$ -sensitive chemistry.

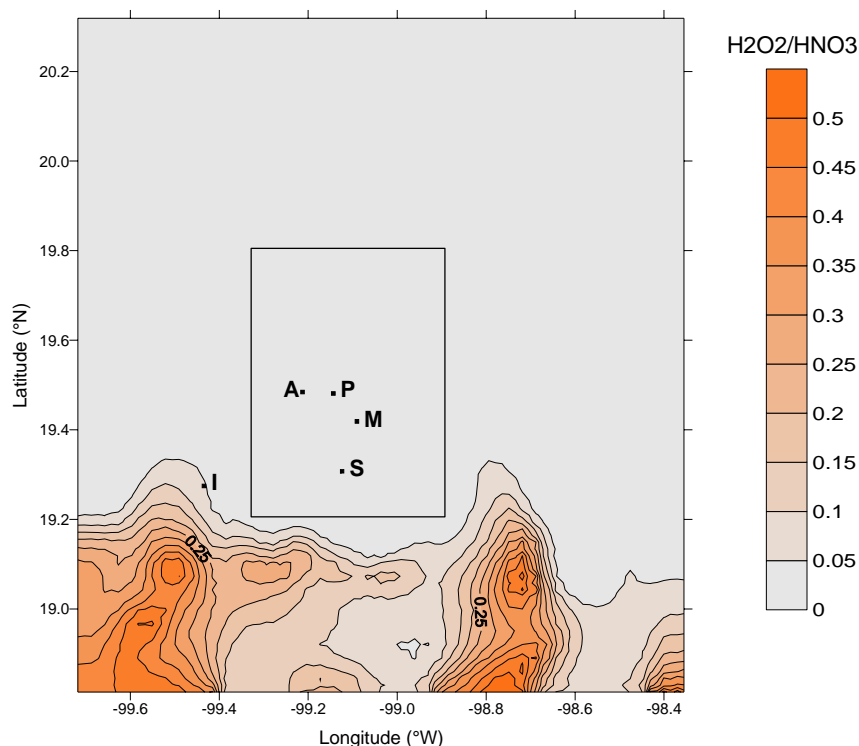
The contour of the band depicting the transition range for the  $\text{O}_3/\text{NO}_y$  ratio (8-10), shown in Figure 4.25(b), suggests also that practically all of the MCMA region was in the VOC-sensitive region ( $\text{O}_3/\text{NO}_z \leq 8$ ). The lowest  $\text{O}_3/\text{NO}_z$  ratios were found towards the northeast of the urban area.

Figure 4.26 shows the simulated distribution of the ratio  $\text{H}_2\text{O}_2/\text{HNO}_3$ . The range of the transition value that separates VOC- and  $\text{NO}_x$ -sensitive conditions (0.25-0.35) confirms the finding of the previous indicator ratios as to the dominance of VOC-sensitive conditions ( $\text{H}_2\text{O}_2/\text{HNO}_3 \leq 0.25$ ). The  $\text{H}_2\text{O}_2/\text{HNO}_3$  ratio, in terms of theory and simulation results, is the most consistent of the indicator ratios for separating VOC- and  $\text{NO}_x$ -sensitive regions (Sillman, 1995; Hammer et al., 2002; Sillman and He, 2002; Thielman et al., 2002).





**Figure 4.25.** Simulated (a)  $O_3/NO_y$  and (b)  $O_3/NO_z$  indicator ratios distribution for the MCMA air basin at the level surface at 15:00 hr Local Standard Time, March 3, 1997. The rectangle in the center shows the approximately extent of the urban area.



**Figure 4.26.** Simulated  $\text{H}_2\text{O}_2/\text{HNO}_3$  indicator ratio distribution for the MCMA air basin at the level surface at 15:00 hr Local Standard Time, March 3, 1997. The rectangle in the center shows the approximately extent of the urban area.

Hammer et al. (2000) reviewed the theoretical basis for the calculation of the transition value of the  $\text{H}_2\text{O}_2/\text{HNO}_3$  indicator ratio in a number of studies from Europe and U.S. They found that even with different theoretical approximations, the transition values ranged between  $\sim 0.18$  (Berlin, Germany) and  $\sim 0.59$  (Po Valley, Italy), although the average transition value was around 0.2.

According to Thielman et al. (2002), the  $\text{H}_2\text{O}_2/\text{HNO}_3$  ratio is close to zero under VOC-sensitive conditions where nitric acid is the dominant sink (i.e., the production of  $\text{HNO}_3 \gg$  the production of  $\text{H}_2\text{O}_2$ ). As the air mass ages,  $\text{NO}_x$  is consumed faster than hydrocarbons (i.e., VOCs in general), resulting in a ratio that approaches zero. The VOC oxidation pathway shifts the ratio to higher values, when the production of  $\text{H}_2\text{O}_2$  becomes important as a sink of radicals. In addition, the ratio  $\text{H}_2\text{O}_2/\text{HNO}_3$  is particularly appropriate because the lifetimes of  $\text{H}_2\text{O}_2$ ,  $\text{HNO}_3$ , and  $\text{O}_3$  (controlled mainly by their deposition rates) are relatively similar.

## 4.2.6 Intercomparison Tests

### 4.2.6.1 Measured Data

As previously mentioned, Sosa et al. (2000) explored the identification of local  $O_3$ - $NO_x$ -VOC sensitivity by applying the photochemical indicators method (PIM) to data collected at different sites of Mexico City (Merced, Azcapotzalco, and IMP) between February and March 1997 (Gaffney et al., 1999; Edgerton, et al., 1999). Table 4.12 presents a comparison between the averages of measured and estimated concentrations of indicator species and photochemical ratios for the Santa Ursula site (this study) and the Merced and Azcapotzalco sites (Sosa et al., 2000).

The results shown in Table 4.12 indicate that, in general, both the  $O_3/NO_y$  and the  $O_3/NO_z$  ratios from Sosa et al. (2000) coincided with the results from SUMS. These ratios indicate VOC-sensitive conditions in Mexico City, which contradicts their finding of  $NO_x$ -sensitive conditions from the  $O_3/(HNO_3 + NO_3^-)$  ratio.

This contradiction can be explained by reviewing the data that Sosa et al. (2000) used for their  $O_3$ - $NO_x$ -VOC sensitivity analysis. In sum, these authors assumed that the concentrations of some species measured at different sites were equal to the expected concentrations of the same species in the other locations. Moreover, they did not take into account the potential interferences and the problems typical of their method during the original measurement campaigns. For example, they did not correct the original  $NO_x$  data for the well-known interferences of PAN and  $HNO_3$  in chemiluminescence  $NO_x$  analyzers. Nor did they review in-depth the potential of  $NH_3$  emissions to influence reported  $HNO_3$  concentrations at the Merced site. They obtained the total  $NO_y$  concentration at the Azcapotzalco and Merced sites as:  $NO_x$  (measured at both sites) + PAN (measured at IMP) +  $HNO_3^-$  (measured only at Merced and as the sum of  $HNO_3$  + volatilized  $NO_3^-$ ).

In addition to these sources of uncertainty in the measurements, the locations selected for the indicators evaluation were not appropriate. The three sites (Merced, Azcapotzalco, and IMP) are located close to the downtown of Mexico City but upwind and far away from the south-southwestern smog

receptor sites of MCMA. Use of these sites and of data inappropriate for the application of the photochemical indicators approach account for Sosa et al.'s inconsistent conclusions.

**Table 4.12.** Comparison of average measured and estimated indicator ratios and photochemical indicator ratios for data collected at Santa Ursula (April, 2004) with results for Merced and Azcapotzalco sites in Mexico City (March, 1997).

Species	Santa Ursula	Merced <sup>a</sup>	Azcapotzalco <sup>a</sup>
	April 14-25, 2004	March 2-19, 1997	March 2-19, 1997
	13:00-17:00 hr	12:00-17:00 hr	12:00-17:00 hr
O <sub>3</sub>	94	92	48
NO <sub>y</sub>	37.3	76.5 <sup>g</sup>	51.2 <sup>g</sup>
NO <sub>x</sub>	16.3 <sup>c</sup> (22.3) <sup>d</sup>	62.7 <sup>d</sup>	37.5 <sup>d</sup>
NO	2.6 <sup>d</sup>	26.6 <sup>d</sup>	5.2 <sup>d</sup>
PAN	-	10.4 <sup>h</sup>	10.4 <sup>h</sup>
PPN	-	-	-
PBN	-	-	-
HNO <sub>3</sub>	-	1.25 <sup>i</sup>	1.25 <sup>i</sup>
NO <sub>3</sub> <sup>-</sup>	-	0.70 <sup>i</sup>	0.70 <sup>i</sup>
NO <sub>z</sub>	21 <sup>e</sup> (15) <sup>f</sup>	12.3 <sup>j</sup>	12.3 <sup>j</sup>
O <sub>3</sub> /NO <sub>y</sub>	2.5	1.2	0.9
O <sub>3</sub> /NO <sub>z</sub>	4.47 <sup>e</sup> (5.8) <sup>f</sup>	7.4	3.8
O <sub>3</sub> /HNO <sub>3</sub>	-	73.6	38.4
O <sub>3</sub> /(HNO <sub>3</sub> + NO <sub>3</sub> <sup>-</sup> )	-	47.1 <sup>k</sup>	24.6 <sup>k</sup>

<sup>a</sup> From Sosa et al. (2000).

<sup>b</sup> From box model applied to measurements performed at the IMP site Gaffney et al., (1999).

<sup>c</sup> NO<sub>x</sub>\* = NO + NO<sub>2</sub><sup>s</sup>, where NO<sub>2</sub><sup>s</sup> was obtained with the BUWAL equation (BUWAL, 1997).

<sup>d</sup> Measured by the NO<sub>x</sub> analyzer at the site.

<sup>e</sup> NO<sub>z</sub>\* = NO<sub>y</sub> - NO<sub>x</sub>\*.

<sup>f</sup> NO<sub>z</sub> = NO<sub>y</sub> - NO<sub>x</sub>, where NO<sub>x</sub> are the measurements reported by the analyzer.

<sup>g</sup> NO<sub>y</sub> = NO<sub>x</sub> + PAN + HNO<sub>3</sub> + volatilized NO<sub>3</sub><sup>-</sup>.

<sup>h</sup> Measured at the IMP site (Gaffney et al., 1999).

<sup>i</sup> Measured at the Merced site from 12:00 to 18:00 hr. NO<sub>3</sub><sup>-</sup> from the volatilization of nitrate aerosols collected on the filter (Chow et al., 2002).

<sup>j</sup> NO<sub>z</sub> = PAN + HNO<sub>3</sub> + NO<sub>3</sub><sup>-</sup>

<sup>k</sup> The sum (HNO<sub>3</sub> + NO<sub>3</sub><sup>-</sup>) is taken as a surrogate of the total HNO<sub>3</sub>.

#### 4.2.6.2 Modeling Data

A direct comparison between the measurements carried out in Santa Ursula and the predictions from the MCCM and other models could not be performed due to differences in their date and time periods. However, several indirect comparisons could be done based on the closeness and relative

contributions of the indicator values to the final indicator ratios and based on the models' overall performance.

The criterion for the selection of the day for the intercomparison between Santa Ursula and the MCCM was that the measured  $O_3$  concentration should be as similar as possible to the predicted concentration. In order to evaluate the consistency of the MCCM results, the predicted values, from this model, were compared with measured values registered at two reference sites and with results from the CIT and box-model simulations. The reference sites were the monitoring stations of Pedregal and Merced. The Pedregal monitoring site was chosen because the Santa Ursula monitoring station was not in operation in 1997 and because it was the closest station to Santa Ursula (~ 5 km west). The CIT modeling results for March 2, 1997 and the general box-model results for March 1997 were used for the overall intercomparison.

Table 4.13 shows the comparison of measured indicator species and indicator ratios for Santa Ursula (March 1997 and April 2004) and Merced (March 1997) with the simulated indicator species and indicator ratios from the modeling predictions of the MCCM, the CIT model (Sosa et al., 2000), and a box model approximation (Gaffney, et al., 1999).

The average 14:00-15:00 hr measured  $O_3$  concentration on April 15, 2004 at Santa Ursula was 93 ppb, while the MCCM simulated concentration for March 3, 1997 at this site was 90 ppb. The measured  $O_3$  concentration on March 3, 1997 at the Pedregal station was 126 ppb, while the predicted concentration at this position was 130 ppb (see Figure 4.23). Similarly, the  $O_3$  predicted concentration at the Merced position was 94 ppb, while the average measured level was 96 ppb. Although the measured and MCCM predicted  $NO_y$  values for Santa Ursula (35 and 42.5 ppb, respectively) resulted in relatively high values, they had a difference of ~ 21%. The predicted MCCM  $NO_y$  concentration for the Merced site was also high, but it was 30% less than the reported 6-hr average measured value. The tendency towards high  $NO_y$  values was also observed with the other measured and predicted results (range of predicted  $NO_y$  concentrations: 42.5-53.4 ppb).

**Table 4.13.** Comparison of measured indicator species and indicator ratios for Santa Ursula (March 3, 1997 and April 15, 2004) and Merced (March 2 and 3, 1997), with the simulated indicator species and indicator ratios predictions from the MCCM (March 3, 1997), the CIT model (March 2, 1997), and a simple box model (March, 1997). All concentrations are in ppb.

Species	Santa Ursula April 15, 2004 measured 14:00-15:00 hr averages	Santa Ursula <sup>a,b</sup> March 3, 1997 simulated (and measured) 14:00- 15:00 hr averages MCCM	Merced <sup>a,b,c</sup> March 3, 1997 simulated 14:00-15:00 hr (and measured 12:00- 18:00 hr) averages MCCM	Merced <sup>c</sup> March 2, 1997 simulated (and measured) 12:00- 18:00 hr averages CIT	MCMA <sup>d</sup> March, 1997 Afternoon hours averages Box-model
O <sub>3</sub>	93	90 (126) <sup>b</sup>	96 (107)	120 <sup>n</sup> (154)	200
NO <sub>y</sub>	35	42.5	42.4 (57) <sup>i</sup>	46 <sup>n,o</sup> (76.8) <sup>i</sup>	53.4 <sup>f,j,r</sup>
NO <sub>x</sub>	12.3 <sup>e</sup>	16.5 (48) <sup>f,h</sup>	17.3 (39) <sup>f</sup>	25 <sup>n</sup> (46) <sup>f</sup>	20 <sup>s</sup>
NO	2 <sup>f</sup>	2.7 (8) <sup>f,h</sup>	2.7	4 <sup>n</sup>	3.4 <sup>f,j</sup>
PAN	-	1.9	2 (15.5) <sup>j</sup>	13 <sup>n,p</sup> (26) <sup>j</sup>	20 <sup>j</sup>
HNO <sub>3</sub>	-	21.9	22 (1.9) <sup>k</sup>	8 <sup>n</sup> (4) <sup>k</sup>	10 <sup>t</sup>
NO <sub>3</sub> <sup>-</sup>	-	1.3	1.1 (0.5) <sup>k</sup>	(0.5) <sup>k</sup>	-
NO <sub>z</sub>	22.7 <sup>g</sup>	25.1	25.1 (17.9) <sup>l</sup>	21 <sup>q</sup> (30.8) <sup>l</sup>	30 <sup>q</sup>
O <sub>3</sub> /NO <sub>y</sub>	2.6	2.1	2.2 (1.88)	2.6 (2.0)	4
O <sub>3</sub> /NO <sub>z</sub>	4.1	3.5	5.5 (5.9)	5.7 (5.0)	6.6
O <sub>3</sub> /HNO <sub>3</sub>	-	4.1	4.3 (44.5) <sup>m</sup>	15 (32.0)	20

<sup>a</sup> Interpolated from the results of the simulation of the spatial distribution.

<sup>b</sup> From historical RAMA data files (<http://sma.df.gob.ma/simat/>).

<sup>c</sup> From Sosa et al. (2000).

<sup>d</sup> From box model applied to measurements performed at the IMP site Gaffney et al., (1999) and additional empirical estimations.

<sup>e</sup> NO<sub>x</sub>\* = NO + NO<sub>2</sub><sup>s</sup>, where NO<sub>2</sub><sup>s</sup> was obtained with the BUWAL equation.

<sup>f</sup> Measured by the NO<sub>x</sub> analyzer at the site.

<sup>g</sup> NO<sub>z</sub>\* = NO<sub>y</sub> - NO<sub>x</sub>\*.

<sup>h</sup> Measured at the Pedregal monitoring station, ~ 5 km west from the Santa Ursula site. No data were available for Santa Ursula by 1997.

<sup>i</sup> NO<sub>y</sub> = NO<sub>x</sub> + PAN + HNO<sub>3</sub> + volatilized NO<sub>3</sub><sup>-</sup>.

<sup>j</sup> Measured at the IMP site (Gaffney et al., 1999).

<sup>k</sup> Measured at the Merced site from 12:00 to 18:00 hr. NO<sub>3</sub><sup>-</sup> from the volatilization of nitrate aerosols collected on the filter (Chow et al., 2002).

<sup>l</sup> NO<sub>z</sub> = PAN + HNO<sub>3</sub> + NO<sub>3</sub><sup>-</sup>

<sup>m</sup> Total HNO<sub>3</sub> as the sum (HNO<sub>3</sub> + NO<sub>3</sub><sup>-</sup>).

<sup>n</sup> From modeling results with the hydrocarbon emissions inventory multiplied by a factor of three.

<sup>o</sup> NO<sub>y</sub> = NO<sub>x</sub> + PAN + HNO<sub>3</sub>.

<sup>p</sup> Modeled at the IMP site position.

<sup>q</sup> NO<sub>z</sub> = PAN + HNO<sub>3</sub>.

<sup>r</sup> Estimated as the sum: NO<sub>y</sub> box-model = "NO<sub>y</sub>" + NO; where reported measured "NO<sub>y</sub>" was NO<sub>y</sub> - NO (Gaffney et al., 1999).

<sup>s</sup> NO<sub>x</sub> = "NO<sub>y</sub>" box model - NO<sub>z</sub>

<sup>t</sup> Assuming that HNO<sub>3</sub> was equivalent to the 50% of the PAN concentration, based on comments of Gaffney et al., (1999).

The relative contributions of  $\text{NO}_x$  and  $\text{NO}_z$  to the total  $\text{NO}_y$  in the measured and the MCCM values were somewhat similar. In the case of the measured concentrations at Santa Ursula, 35% were from  $\text{NO}_x$  and 65% from  $\text{NO}_z^*$ , while the MCCM predicted contributions for this site were 40% from  $\text{NO}_x$  and 60% from  $\text{NO}_z$ . Very similar relative contributions were found in the simulated values for Santa Ursula and those for the Merced site. However, the relative contributions from the MCCM and those measured at Merced were different. While the simulated contributions were 41% from  $\text{NO}_x$  and 59% from  $\text{NO}_z$ , the measurements were 61% from  $\text{NO}_x$  and 39% from  $\text{NO}_z$ .

The intercomparison of the relative  $\text{NO}_x$  and  $\text{NO}_z$  contributions to the predicted  $\text{NO}_y$  at the Merced site from the MCCM for March 3, 1997 and from the CIT model for March 2, 1997 resulted in similar predictions: 41% from  $\text{NO}_x$  and 51% from  $\text{NO}_z$  for the MCCM and 55% from  $\text{NO}_x$  and 45% from  $\text{NO}_z$  for the CIT model. The relative contributions predicted by the box model were also similar: 56% from  $\text{NO}_x$  and 44% from  $\text{NO}_z$ .

The measured  $\text{NO}_z^*$  and MCCM simulated  $\text{NO}_z$  values for the Santa Ursula site were very similar (22.7 ppb and 25.1 ppb, respectively), while the measured and MCCM predicted  $\text{NO}_z$  values at Merced were different (12.3 ppb and 25.1 ppb, respectively). The  $\text{NO}_z$  concentrations from the CIT and the box model, as well as from the measurements for March 2, 1997, had values very close to those of the MCCM.

Despite the differences in the relative contributions of  $\text{NO}_x$  and  $\text{NO}_z$  among the measurements and models, the modeled  $\text{O}_3/\text{NO}_y$  and  $\text{O}_3/\text{NO}_z$  ratios for March 1997 were very similar to the measured conditions at Santa Ursula in April 2004. All of the models indicated VOC-sensitive conditions. However, several inconsistencies emerged from the intercomparison of the predicted and measured values for PAN and  $\text{HNO}_3$ . While the MCCM predicted a PAN contribution to the  $\text{NO}_z$  of 8%, the CIT model predicted a contribution of 61%. The MCCM predicted that  $\text{HNO}_3$  concentrations accounted for 87% to the total  $\text{NO}_z$ , while the CIT model predicted a nitric acid contribution of 38%. Furthermore, the 12:00-18:00 hr average PAN measurements reported for both reference days, March 2 and 3, 1997, at the IMP site were ~ 15.5 and 10.5 ppb, respectively, while the measured  $\text{HNO}_3$  concentrations were ~ 4.2 and 2 ppb respectively.

Sosa et al. (2000) suggested that the PAN and HNO<sub>3</sub> overestimations observed in the CIT simulations could be explained by the model's failure to reproduce both the chemical processes of these compounds and the vertical mixing processes. On the other hand, García (2004) suggested that the overestimation of HNO<sub>3</sub> and underestimation of PAN in the MCCM predictions could be related to deficiencies in the official VOC emission inventory used in the simulation. In the opinion of this researcher, the underestimation of ammonia emissions and the poor balance contributions of aldehydes in the official VOC inventory affected the correct prediction of these compounds.

The lack of NH<sub>3</sub> emissions in a photochemical model may cause errors in both the simulation of HNO<sub>3</sub> because of the particular chemistry of these two compounds in the air. The equilibrium reaction among ammonia and nitric acid ( $\text{HNO}_{3(g)} + \text{NH}_{3(g)} \leftrightarrow \text{NH}_4\text{NO}_{3(s, \text{aq})}$ ) implies that the concentration of these gases in the air depends on the initial proportions of each of them (Finlayson and Finlayson-Pitts, 2000). While the lack of NH<sub>3</sub> emissions in modeling simulations results in an obvious overestimation of HNO<sub>3</sub> (i.e., there is no possibility of conversion of nitric acid to ammonium nitrate), in the sampling and measurement of nitric acid, the presence of high NH<sub>3</sub> levels may interfere strongly with the real ambient HNO<sub>3</sub> concentration. Chow et al. (2002) report that measured NH<sub>3</sub> concentrations at the Merced site were, on average, ~ 3.7 times greater (in ppb) than the HNO<sub>3</sub> concentrations there. Therefore, it is possible that reported measurements of HNO<sub>3</sub> were strongly affected by NH<sub>3</sub> resulting in a “strong” underestimation of the real values. It is not clear if both models overestimated the real HNO<sub>3</sub> concentrations simply because their simulated values were compatible with the range of nitric acid values reported by Limón-Sánchez (2002).

On the other hand, because aldehydes can react with OH radicals to produce peroxyacetyl radicals that rapidly react with NO<sub>2</sub>, thereby producing PAN (Singh, 1987), a poor balance of aldehydes in the VOC emissions inventory used in the models could result an underestimation of PAN concentrations.

Although not stated explicitly, Sosa et al. (2000) artificially increased PAN and HNO<sub>3</sub> concentrations to validate the performance of the CIT model. They used a combination of the measured PAN concentrations from IMP and the CIT predicted HNO<sub>3</sub> concentrations to obtain an adjusted O<sub>3</sub>/NO<sub>y</sub>



ratio. The resulting adjusted ratio was very similar to the  $O_3/NO_y$  ratio estimated from the measurements for the same day. In this way, they indirectly validated the performance of the CIT to study the effect of varying  $NO_x$  and VOC emissions on the spatial distribution of  $O_3$  in the MCMA.

The impact of the overestimation of  $HNO_3$  and underestimation of PAN in the final balance of the  $NO_z$  species predicted by the MCCM cannot be clarified with the available information. However, a readjustment of the emission inventories used by the model probably would not change, to a large extent, the final result because there would be a redistribution of the oxidized nitrogen among these species. In addition, Sillman (1999), Sillman et al. (2003), and Tonessen and Dennis (2000) have suggested that the correlation between  $NO_x$ -VOC sensitivity and  $O_3/NO_y$  and  $O_3/NO_z$  ratios is driven specifically by  $HNO_3$  rather than by PAN and other organic nitrates.

It should be noted, however, that the MCCM predictions were within the range of acceptability given by Sillman (1995). According with this author, in modeling experiments with predicted high  $O_3$  concentrations, the ratio  $O_3/(2H_2O_2 + NO_z)$  should have a near-constant value between 5 and 8, regardless of VOC- or  $NO_x$ -sensitivity. Taking the predicted distribution of the individual concentrations for each of the species in this ratio, the resulting  $O_3/(2H_2O_2 + NO_z)$  ratio was 6.7 (s.d. = 0.71) for 165 data points representative of the MCMA.

The general tendency of the MCCM and the CIT model to predict VOC-sensitive conditions in the Mexico City Metropolitan Area suggests that the results of the indicator measurements in Santa Ursula were consistent with the PIM approach.

### **4.3 Comparative Assessment of the Ozone Formation in Cincinnati and Mexico City**

This section presents a comparative assessment of the main characteristics of the Cincinnati and Mexico City metropolitan areas associated with the formation of ozone and with  $O_3$ - $NO_x$ -VOC sensitivity in the two airsheds. The results of this assessment are combined with several conceptual definitions to evaluate the consistence of the  $NO_x$ -VOC sensitivity dictated by the PIM and the  $VOC/NO_x$  ratio under

severe air pollution conditions. Table 4.14 contains a comparison of selected statistics between the two urban areas for the year 2000.

**Table 4.14.** Comparison of selected statistics between Greater Cincinnati and the Mexico City Metropolitan Area (year 2000).

Statistic parameter	Greater Cincinnati <sup>a</sup>	Mexico City Metropolitan Area <sup>b</sup>
Population	1'817,571 (331,285) <sup>c</sup>	18'171,318 (8'605,239) <sup>d,e</sup>
Total Area (km <sup>2</sup> )	16,583 (206) <sup>c</sup>	5,300 (1,602) <sup>d,e</sup>
Population density (inhabitants/km <sup>2</sup> )	109 (1,608) <sup>c</sup>	12,000 (central area) <sup>e</sup> 2,700 (periphery) <sup>e</sup>
GDP per capita in US dollars	34,646 <sup>f</sup>	7,750 <sup>e</sup>
Vehicle fleet	1.6 million <sup>g</sup>	3.2 million <sup>e</sup>
Average vehicle age	~ 6 years (projected) <sup>h</sup>	~ 10 years <sup>e</sup>
VKT (kilometers per day)	134 million (total) <sup>g</sup>	153 million (total) <sup>e</sup>
Industries (major & small)	~ 1000 <sup>g</sup>	~ 35,000 <sup>e</sup>
NO <sub>x</sub> emissions (tons/yr)	239,277 <sup>g</sup> 63,861 (27% vehicles) 119,115 (49% power plants) 17,832 (7.4% fuel combustion industrial)	193,451 <sup>i</sup> 156,695 (81% vehicles) 11,626 (6% power plants) 6,155 (3.2 % fuel combustion industrial)
VOC emissions (tons/yr)	161,072 <sup>g</sup> 41,510 (25% vehicles) 39,151 (24% biogenic) 308 (0.2% fuel combustion industrial)	429,755 <sup>i</sup> 194,517 (45% vehicles) 15,425 (3.6% biogenic) 7,876 (1.8% fuel combustion industrial)
NO <sub>x</sub> emissions density (tons/day-km <sup>2</sup> )	0.039	0.100
VOC emissions density (tons/day-km <sup>2</sup> )	0.026	0.222
VOC/NO <sub>x</sub> ratio (ton/ton)	0.67	2.2
VOC/NO <sub>x</sub> ratio (ppmC/ppm)	1.9	15.1
O <sub>3</sub> 1-hr maximum 2003 (ppb)	140 <sup>j</sup>	243 <sup>k</sup>
PM <sub>10</sub> 24-hr maximum 2003 (µg/m <sup>3</sup> )	80 <sup>j</sup>	744 <sup>k</sup>
PM <sub>10</sub> arithmetic mean maximum 2003 (µg/m <sup>3</sup> )	23 <sup>j</sup>	59 <sup>k</sup>

<sup>a</sup> Includes 5 counties of Ohio, 6 counties of Kentucky, and 2 counties of Indiana.

<sup>b</sup> Includes 16 delegations from Mexico City and 17 municipalities of the State of Mexico.

<sup>c</sup> City of Cincinnati.

<sup>d</sup> Mexico City.

<sup>e</sup> Molina and Molina (2002).

<sup>f</sup> Demographia/Wendell Cox Consultancy. Belleville, IL 62222, USA, 2003. [www.demographia.com](http://www.demographia.com).

<sup>g</sup> US EPA (2004).

<sup>h</sup> Ohio Department of Transportation. 1998 State of the Transportation System. [www.dot.state.oh.us/sos98/sosPT.htm](http://www.dot.state.oh.us/sos98/sosPT.htm).

<sup>i</sup> SMA (2003).

<sup>j</sup> HCDOES (2004).

<sup>k</sup> SMA (2004).

There are a number of differences between Greater Cincinnati (GC) and the Mexico City Metropolitan Area (MCMA) directly related to the air quality in both regions. By the year 2000, the population in the MCMA was almost 10 times greater than in GC, but its total area was nearly three times smaller than that of GC. This situation is clearly reflected in estimates of population density. The number of inhabitants per square kilometer was almost 12 times greater in the MCMA than that in GC. Strongly related to population density, the use of cars for transportation represents an additional factor in the higher estimations of vehicle emissions in the MCMA. While the vehicle fleet in GC was ~ 1.6 million with an estimated vehicle age of 6 years, the number of vehicles in the MCMA was almost two times that and with an estimated age of 10 years. However, the VKT (vehicle kilometers traveled) per day in both urban areas was very similar. The combination of the vehicle fleet age (around 30% of cars without a catalytic converter) and very severe traffic problems in the MCMA (average speed ~ 11 km/hr) generated much higher vehicle emissions than expected according to the local VKT. The number of industrial facilities in the MCMA was ~ 35 times than in GC. However, the NO<sub>x</sub> and VOC contributions from industrial activity, in both regions, were not significant compared with other sources.

Power plant NO<sub>x</sub> emissions were 10 times greater in GC than in the MCMA. While in GC the contribution of NO<sub>x</sub> emissions from power plants to the total balance was almost 50%, in the MCMA this contribution was only 6%. This apparent inconsistency between population density and power generation is due to the fact that most of the electricity consumed in the MCMA comes from sources located in other parts of the country. Almost 25% of all electricity produced in Mexico is consumed in the MCMA. However, the higher NO<sub>x</sub> emissions from cars in the MCMA, make this region to have its total NO<sub>x</sub> emissions balance very close to that for GC.

On the other hand, the estimated total VOC emissions in MCMA were ~ 2.7 times higher than the same balance for GC. VOC emissions from cars in MCMA were almost 5 times greater than the same emissions contributions in GC, while biogenic VOC emissions in GC were estimated to be nearly 2.6 times greater in this region than in MCMA.

The estimated combined ratio of VOC/NO<sub>x</sub> emissions (as ton VOC/ton NO<sub>x</sub>) seemed to be dominated by NO<sub>x</sub> emissions in GC (VOC/NO<sub>x</sub> = 0.67) and by VOC emissions in MCMA (VOC/NO<sub>x</sub> = 2.2). Sillman (1999) has suggested that in general, NO<sub>x</sub> emissions within an urban area determine the total amount of O<sub>3</sub> that is formed after air moves downwind and chemical processes have run to completion, while VOC emissions control the rate of initial buildup of O<sub>3</sub>. Therefore, it might be expected that O<sub>3</sub> levels were higher in GC. Moreover, the estimated emissions of biogenic VOC precursors in GC were almost twice that of the MCMA. Biogenic VOC species are extremely reactive with most anthropogenic VOC (Chameides et al., 1992). Consequently, the impact of biogenic VOC can be large relative to their ambient concentrations. Even very small concentrations of reactive biogenic VOC may have a very strong impact on the total VOC reactivity. However, the higher reactivity of VOC from car emissions in MCMA and the relatively limited use of catalytic converters give this region a potential higher VOC reactivity than that for GC.

The combined effect of the reactivity-weighted VOC (the sum of each VOC species mass emission weighed by its reaction rate with OH radicals) and the particular physical conditions (total area, emissions density, meteorology, orography, altitude, latitude, etc.) seems to be the difference in the magnitude of the O<sub>3</sub> peak in each region.

#### **4.3.1 Ozone Formation Scenarios**

Topographically, the GC region is a plateau bisected by the Ohio River, composed of a mix of rugged slopes and gently rolling hills. In general, the hills and valleys in the region have only a minimal effect on air movement, although the region's location in the Ohio River Valley makes it prone to morning thermal inversions (OKI, 1979). Large scale movement of air masses in North America, the lack of high elevations across the region, and the relatively fast heating of the ground during the warm months favors the vertical and horizontal dispersion of air pollutants. Because of its latitude (39.8°N) and its relatively high afternoon summer temperatures (average ~ 27.7°C), GC is susceptible to higher ozone concentrations

in the summer and early fall. Ozone is a strongly seasonal secondary pollutant in mid-latitude regions, with the warm months exhibiting higher concentrations than the cool months.

In contrast, the MCMA lies in an elevated basin of the Mexican Plateau. The basin is confined on three sides (east, south, and west) by mountain ridges but has a broad opening to the north and a narrower gap to the south-southwest. The surrounding ridges vary in elevation with several peaks rising to nearly 2000 m above the mean urban floor. Quite the opposite of mid-latitude regions, the subtropical latitude (19°N) and high altitude of the MCMA makes this the area vulnerable to ozone production during winter as well as summer. During the dry winter months, the basin is normally under the influence of high-pressure systems with light winds above the basin and nearly cloudless skies. This favors the formation of strong surface-based inversions at night that persist for several hours following sunrise, trapping morning emissions within a shallow thermal inversion layer that is closed off by the surrounding elevation (Bravo and Torres, 2002). The solar heating of the ground generates sufficient turbulent mixing to erode these inversions by mid-morning, producing deep boundary layers by the afternoon. However, there is sufficient time in the morning hours for photochemical ozone formation.

These differences result in very dissimilar photochemical scenarios for GC and the MCMA. Table 4.15 presents a comparison of ozone formation and related parameters obtained during late summer in Cincinnati (Taft site) and late winter in Mexico City (Santa Ursula site). Although the O<sub>3</sub> concentrations were always higher in Mexico City, there were several similarities between some of the parameters leading to O<sub>3</sub> formation. The average 24-hr NO concentrations were very similar in both cities (~ 14.5 ppb). The afternoon average was similar as well (~ 2.7 ppb).

However, there was a strong difference between the 6:00-9:00 hr NO and NO<sub>x</sub> concentrations in Cincinnati and Mexico City: 25.9 and 73.9 ppb for NO and 47 and 116 ppb for NO<sub>x</sub>, respectively. The apparent resemblance in the 24-hr and afternoon averages suggests that there could be a similarity between NO<sub>x</sub> emissions. The difference in morning NO and NO<sub>x</sub> concentrations seems to be due to the difference in emission density and in the strength of the morning inversion layer. The inversion layer is likely stronger in the MCMA as a result of the barrier effect of the surrounding mountains.

**Table 4.15.** Comparison of ozone formation and related parameters measured at the Taft monitoring station, Cincinnati (September 3-30, 2003) and the Santa Ursula monitoring station, Mexico City (April 14-25, 2004).

Statistic parameter	Taft monitoring station	Santa Ursula monitoring station
O <sub>3</sub> (24-hr average), ppb	24.8	44.7
O <sub>3</sub> max, ppb	73	172
O <sub>3</sub> afternoon average, ppb	59	94
NO (24-hr average), ppb	14.3	14.8
NO (afternoon average), ppb	2.9	2.6
NO (6-9 a.m. average), ppb	25.9	73.9
NO <sub>y</sub> (24-hr average), ppb	41.2	54.7
NO <sub>y</sub> (maximum morning peak), ppb	213	228
NO <sub>y</sub> (afternoon average), ppb	21.6	37.3
NO <sub>x</sub> <sup>a</sup> (6-9 a.m. average), ppb	47.0	116
NO <sub>z</sub> <sup>*</sup> (afternoon average), ppb	7.9	21
O <sub>3</sub> /NO <sub>y</sub>	2.7 (VOC-sensitive)	2.5 (VOC-sensitive)
O <sub>3</sub> /NO <sub>z</sub> <sup>*</sup>	5.9 (VOC-sensitive)	4.4 (VOC-sensitive)
NO <sub>x</sub> <sup>*</sup> /NO <sub>y</sub> (average air masses age)	0.62	0.44
Average rate of O <sub>3</sub> accumulation (ppb/hr)	5.3	21.4
Average time of effective O <sub>3</sub> build up (hr)	5.2	3.3
Average rate of NO depletion (ppb/hr)	3.3	17
WE-O <sub>3</sub> max (ppb)	48.6	86.6
WD-O <sub>3</sub> max (ppb)	46.4	96.1
WE-NO 6-9 a.m. (ppb)	16.1	61.1
WD-NO 6-9 a.m. (ppb)	31.7	97.5
Ambient temperature (average), °C	18.1	19.7
Wind speed (average), m/s	1.8	1.5
Afternoon mixing layer height (average) <sup>a</sup> m	1416	1917

<sup>a</sup> As measured by the NO<sub>x</sub> instrument.

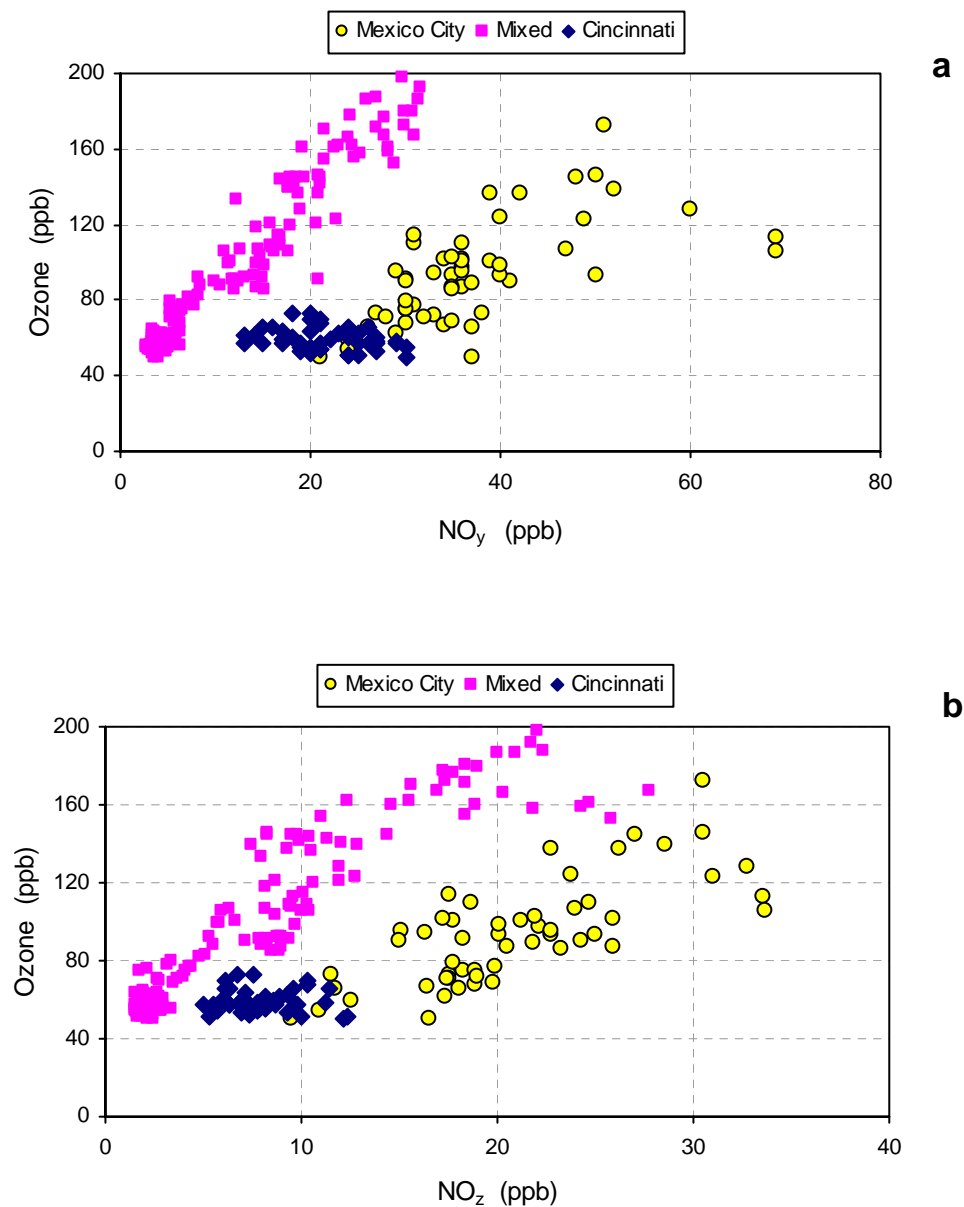
<sup>b</sup> Average of mixing layer estimations from READY model interpolations to FNL datasets. Real-Time Environmental Applications and Display System. National Climatic and Atmospheric Administration. Office of Oceanic and Atmospheric Research. Air Resources Laboratory. [www.arl.noaa.gov/ready.html](http://www.arl.noaa.gov/ready.html).

The average NO<sub>y</sub> morning peak was a little higher in Mexico City (229 ppb) as compared to that in Cincinnati (213 ppb). Nevertheless, the difference increased notably for the average afternoon NO<sub>y</sub> concentrations: 37.3 ppb in Mexico City and 21.6 ppb in Cincinnati. The balance of the NO<sub>y</sub> was likely

associated with the difference in  $\text{NO}_z^*$  afternoon species. While in Mexico City the  $\text{NO}_z^*$  afternoon level had an average of 21 ppm, in Cincinnati the same species had an average of 7.9 ppb.

As previously discussed, measured and simulated PAN and  $\text{HNO}_3$  concentrations strongly suggest that these species are present in the ambient air of the MCMA at relatively high levels compared to other parts of the world. On the other hand, the measured  $\text{HNO}_3$  concentrations in Cincinnati, which are within the range typical of the midwestern U.S., suggest that nitric acid is the most important and abundant oxidized nitrogen species in this region.

The  $\text{O}_3/\text{NO}_y$  ratios were similar for both cities (2.7 in Cincinnati and 2.5 in Mexico City). However, the scatterplot correlations of these two species were different although both indicate VOC-sensitive conditions (see Figure 4.27 (a)). The relatively wide scattering of the data points for Mexico City could imply that the chemistry character of the air parcels in the MCMA airshed were strongly variable, while remaining within the VOC-sensitivity regime. According Sillman (2003), the  $\text{O}_3/\text{NO}_y$  ratio reflects two separate processes, both of which are related to  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity. On the one hand, the ratio reflects the process of photochemical ozone production. On the other, it reflects the process of  $\text{NO}_x$  titration and the immediate removal of  $\text{O}_3$  by reaction with directly emitted NO. The other secondary species are unaffected by  $\text{NO}_x$  titration. The fact that there was relatively little variation in the  $\text{O}_3$  versus  $\text{NO}_y$  correlation in Cincinnati could be interpreted to mean that the  $\text{NO}_x$  titration processes were preventing the accumulation of ozone. Thus, it could be assumed that  $\text{O}_3$  production efficiency in Cincinnati was strongly limited by the relatively high  $\text{NO}_x$  emissions in the region. The large variation observed in the  $\text{O}_3$  versus  $\text{NO}_y$  correlation in Mexico City could indicate that both  $\text{O}_3$  production and  $\text{NO}_x$  were acting simultaneously. However, because  $\text{NO}_y$  by itself may be considered an indicator of the impact of stagnant meteorology on the photochemical processes (Sillman, 1995), the observed variation in the  $\text{O}_3$ - $\text{NO}_y$  correlation in Mexico City could be taken as an indication of the variations of the meteorological conditions during the monitoring campaign. According Sillman, Stagnant meteorology and associated high  $\text{NO}_x$ , VOC, and  $\text{NO}_y$  cause an increase in the photochemical lifetimes of  $\text{NO}_x$  and VOC.



**Figure 4.27.** Scatterplot patterns comparison of (a) measured afternoon O<sub>3</sub> and NO<sub>y</sub> concentrations and, (b) measured afternoon O<sub>3</sub> and estimated NO<sub>z</sub>\* concentrations in Cincinnati (Taft monitoring station) and Mexico City (Santa Ursula monitoring station), and 3-D model predicted scatterplot patterns for O<sub>3</sub> and NO<sub>y</sub>, and O<sub>3</sub> and NO<sub>z</sub> for locations with mixed or with near-zero sensitive conditions. Simulation data obtained from data file available from Sillman (2003).

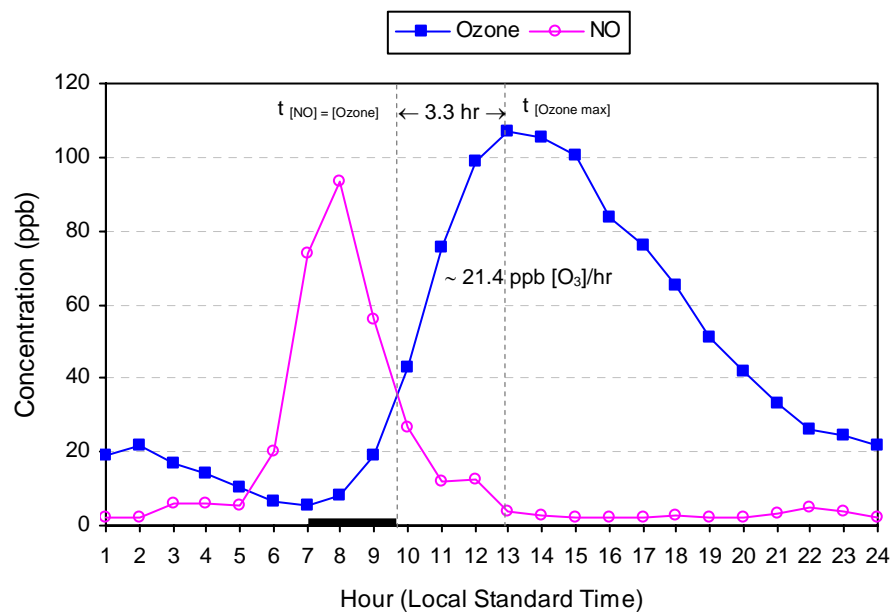
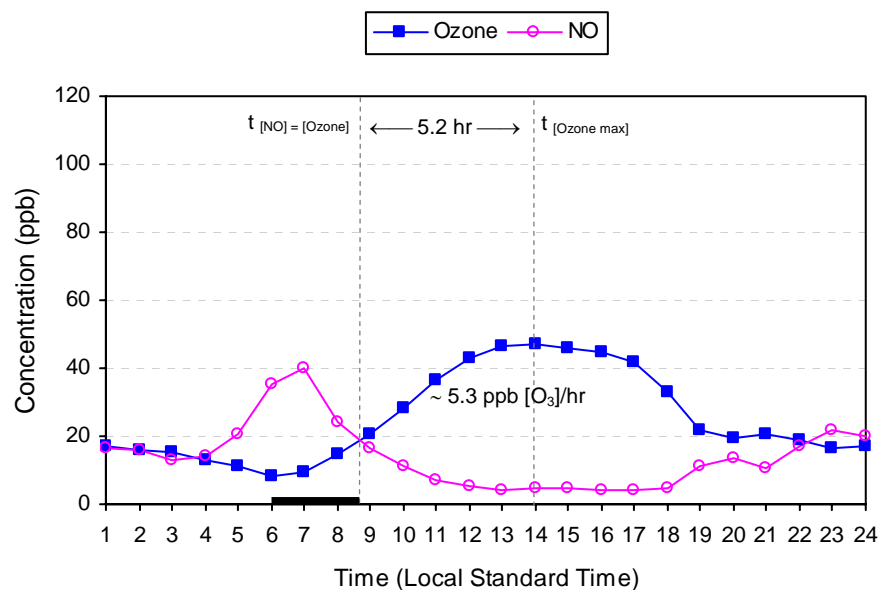


The  $O_3/NO_z^*$  ratios were somewhat different (5.9 in Cincinnati and 4.4 in Mexico City), suggesting that ozone production was more efficient in Cincinnati. However, Figure 4.27 (b) shows that the correlation of  $O_3$  versus  $NO_z^*$  for Cincinnati lacked the covariance seen in Mexico City. This confirms that the apparent excess of  $NO_x$  in Cincinnati was controlling the  $O_3$  accumulation, while in Mexico City the photochemical processes were dominating the NO titration process. In both cases, the PIM method showed that VOC-sensitive conditions were dominant as was confirmed by the statistical analysis of the differences in the WE/WD peak  $O_3$ .

The effect of the VOC emissions on the strength of the  $O_3$  initial buildup in each region can be explained by analyzing conceptual factors associated with ozone formation process. Figure 4.28 presents the estimated averages of: time of  $O_3$  inhibition, time of net  $O_3$  accumulation, and rate of  $O_3$  accumulation for each of the sites and monitoring periods.

Although the predawn NO available to start the morning ozone production was in general lower in Cincinnati compared to those in Mexico City, the morning NO peak in the latter city surpassed by far the respective NO peak observed in Cincinnati, resembling a clear effect of the extent of morning precursors emissions and stagnant morning meteorology in Mexico City. It is generally accepted that in the morning  $O_3$  formation is inhibited by high concentrations of NO, which restrain radical formation by titrating  $O_3$ . During this period, formaldehyde (and to a lesser extent, nitrous acid), are the main source of OH radicals. Because the length of the  $O_3$  inhibition is essentially determined by the concentration of NO, it is expected that this period be shorter in areas of high amounts of fresh  $NO_x$  emissions. However, the average  $O_3$  inhibition period in both cities had approximately the same length (~ 2.7 hr). The possible explanation to this inconsistency is that the production of OH radicals from relatively high morning levels of formaldehyde in Mexico City was of such extent that could rapidly dim the NO titration effect of the photostationary state.

Back in the early 1990s, Bravo et al. (1991) found that the introduction of methyl-t-butyl ether (MTBE) as an additive in Mexican gasoline showed an association with an increase in the morning atmospheric formaldehyde levels and on the  $O_3$  production in MCMA.



**Figure 4.28.** Average hourly diurnal variations of  $O_3$  and NO concentrations and average rate of  $O_3$  accumulation and of time of effective  $O_3$  buildup at the: **(a)** Cincinnati (Taft Monitoring Station, September 3-30, 2003), and **(b)** Mexico City (Santa Ursula Monitoring Station, April 14-25, 2004). The horizontal solid bars represent the approximately length of the  $O_3$  inhibition periods.

Grosjean (2003) has suggested that use of MTBE in unleaded gasoline is an important source of isobutene and formaldehyde emissions from vehicles. Baez et al. (2000) have reported that they measured formaldehyde and acetaldehyde concentrations in the range of 20 ppb from 1993 to 1996 in a smog receptor site of Mexico City. Therefore, it is feasible to assume that VOC emissions from cars in Mexico City could be one additional important source of aldehydes compounds.

As a result of the strong differences in the initial morning emissions and production of OH radicals in both cities, the average NO-O<sub>3</sub> crossover concentration in Mexico City (~ 35 ppb) was almost 1.5 times higher than that for Cincinnati (~ 20 ppb). The average rate of O<sub>3</sub> accumulation and NO depletion was almost four times greater in Mexico City than in Cincinnati. This suggests that, despite the high emission of NO<sub>x</sub> in both regions, the production of NO<sub>2</sub> available for O<sub>3</sub> formation and the removal of NO by organic radicals was strongly enhanced in Mexico City by its higher availability of VOCs.

Fujita et al. (2002) have suggested that timing of maximum O<sub>3</sub> generally coincides with distance of pollutant transport. The ozone peak in Cincinnati and Mexico City occurred in the early afternoon hours suggesting that both airsheds were within the influence of local regional emissions. While the average duration of the net O<sub>3</sub> accumulation was around 1.5 times longer in Cincinnati (~ 5.2 hr) than in Mexico City (~ 3.3 hr) however, the average rate of O<sub>3</sub> build up was almost four times faster in Mexico City (~ 21.4 ppb/hr) than in Cincinnati (~ 5.3 ppb/hr). In addition, the estimated rate of NO depletion was significantly higher in Mexico City (~ 17 ppb/hr) than in Cincinnati (~ 3.3 ppb/hr). The average afternoon mixing layer was relatively higher in the MCMA region than in GC, although the average wind speed was somewhat lower in Mexico City than in Cincinnati.

The above results confirm that despite of GC and MCMA have almost the same load of NO<sub>x</sub> emissions into their air basins, the O<sub>3</sub> formation scenarios are different and strongly influenced by the severity of the local stagnant morning meteorological conditions and by the load and reactivity potential of VOC emissions.

### 4.3.2 Implications of the Results

According with the fundamentals of  $O_3$  photochemistry, under a VOC-sensitive system ( $VOC/NO_x$  ratios  $\ll 8$ ), the short supply of sources of OH radicals and the fact that OH radicals react predominantly with  $NO_2$  lead to enhance the production of  $HNO_3$  and to retard the  $O_3$  formation. On the opposite, under a  $NO_x$ -sensitive system ( $VOC/NO_x$  ratios  $\gg 8$ ) the ample abundance of sources of OH radicals is sufficient to react with the relatively short abundance of  $NO_2$  and with other VOCs to generate more radicals. As a result, the production of  $O_3$  is accelerated and the formation of  $H_2O_2$  is increased.

Following the fundamentals above,  $O_3$  formation in Cincinnati ( $VOC/NO_x \sim 1.9$  ppmC/ppm) should have a VOC-sensitive condition while Mexico City ( $VOC/NO_x \sim 15.1$  ppbC/ppm) a  $NO_x$ -sensitive limitation. Although the above definition for VOC-sensitivity seems to be in agreement with the results of the previous observation driven analyses (PIM and WE/WD effect) for the Cincinnati data, the theoretical  $NO_x$ -sensitivity definition that could be applied to Mexico City looks to be in contradiction with the experimental finding that  $O_3$  formation in Mexico City is likely VOC-sensitive.

The response to this apparent contradiction could be related to the limitation of the  $VOC/NO_x$  ratio to reflect the “real”  $NO_x$ -VOC sensitivity under certain conditions, particularly the true impact of VOC in terms of reactivity-weighted concentrations and severity of the air pollution event (Sillman, 1999). Usually, estimations of VOC to calculate  $VOC/NO_x$  ratios are influenced by uncertainties in the emission inventories calculations and the practical limitation to measure all the VOC species present in ambient air samples. As a result, the final estimated VOC are generally dominated by relatively less reactive but very abundant alkanes. This situation seems to be an important factor in the misleading of the true load of VOC with direct impact on  $O_3$  formation in Mexico (Molina and Molina, 2002). However, after adjustments in the MCMA VOC reactivity, the  $VOC/NO_x$  ratio still indicates a  $NO_x$ -sensitive condition ( $VOC/NO_x > 8$ ).

Milford et al. (1994) and Sillman (1999) have used a combination of theoretical analysis and results from photochemical models to demonstrate the existence of stronger correlations between events of

higher overall concentrations of  $O_3$  precursors and VOC-sensitive chemistry systems, and between events with lower concentrations of ozone precursors and  $NO_x$ -sensitive conditions, even with identical initial VOC/ $NO_x$  emissions ratios. They have suggested that environmental conditions prone to have higher ozone precursors can be associated with either higher emission densities (i.e., larger urban areas) or with more restricted meteorological conditions (i.e., light winds and low daytime vertical mixing) or with a combination of both. Sillman has proposed an explanation for the severity of event effect by doing an analogy with the theoretical concept of Kleinman (1991, 1994) that states that, the split between the  $NO_x$ -sensitive and  *$NO_x$ -saturated* (or VOC-sensitive) regimes is related to the relative supply of  $NO_x$  (from emissions) in comparison with supply of radicals generated by sunlight. Freshly emitted plumes of polluted air have an initial  $NO_x$  supply that greatly exceeds the supply of radicals. As the plume ages, the total amount of radicals created during the process of photochemical evolution becomes equal with and eventually surpasses the initial  $NO_x$ -source, causing a switch from  *$NO_x$ -saturated* to  $NO_x$ -sensitive conditions. Therefore, if the source of  $NO_x$  source per unit volume in an airbasin is high (after considering the effect of dilution through daytime vertical mixing), it will take more time for the accumulated source of radicals (from the VOCs) to become equal with and surpass the  $NO_x$  source. As a result, a  *$NO_x$ -saturated* (VOC-sensitive) condition will persist for a longer time. But, if the  $NO_x$  source is low, the accumulated radical source will exceed the  $NO_x$  source after a short period of photochemical aging and the system will shift to  $NO_x$ -sensitive conditions.

Because MCMA has both a high precursors emission density and prevalent stagnant meteorology conditions, it is possible that the combination of these two factors is favoring the persistence of a  *$NO_x$ -saturated* (VOC-sensitive) condition, situation that is not seen by the initial VOC/ $NO_x$  ratio. Furthermore, the very high PAN concentrations observed in Mexico City may allows one to hypothesize that this is an additional factor influencing the limitation of the radical sources to surpass the  $NO_x$  source in the MCMA airbasin. PAN is produced by reaction of acetaldehyde,  $CH_3CHO$ , with OH and  $O_2$  to form  $CH_3CO_3$ , which combines with  $NO_2$  to form PAN (net reaction R12). Using steady-state approximations on the chemistry of PAN formation, Sillman et al. (1990) have shown that PAN concentration increases as

O<sub>3</sub> increases sequestering temporarily OH radicals and carbonyl molecules ( $[PAN] \propto k^*[OH][CH_3CHO][O_3]$ ; where  $k^*$  represents a composite of all of the reaction constants involved in the reactions sequence). Gaffney et al. (1999) have reported the occurrence of a very strong diurnal variation of PANs (PAN, peroxyacetyl nitrate; PPN, peroxypropionyl nitrate; and PPB, peroxybutiryl nitrate) with PAN levels as high as 30 ppb. Therefore, it is reasonable to assume that PAN formation is acting as an important temporal reservoir of radicals from the accumulated burden of radicals necessary to switch the VOC-sensitive system to NO<sub>x</sub>-system.

All of the mentioned above factors seem to be well represented by the evaluation of the NO<sub>y</sub> and the O<sub>3</sub>/NO<sub>y</sub> indicators for MCMA and justify the experimental finding of the prevalence of VOC-sensitive conditions in the Mexico City air basin. The observed high afternoon O<sub>3</sub> concentrations suggests that there were large radical sources, while the high NO<sub>y</sub> values suggests that there were a higher number of radical sinks, such as the OH + NO<sub>2</sub> and CH<sub>3</sub>CHO + OH reactions.

The resulting VOC-sensitivity in Mexico City would seem to be in contradiction with prior studies. There are two reasons that could explain this contradiction. The first one that has been thoroughly discussed by Molina and Molina (2002), points out that the MCMA emission inventories, particularly the VOC inventory, are highly uncertain. As a consequence, photochemical simulations and other studies that used these inventories have given confusing O<sub>3</sub>-sensitivity results. The second is that the use of morning VOC/NO<sub>x</sub> ratios is not appropriate as a direct justification for the MCMA NO<sub>x</sub>-VOC sensitivity because this ratio is not able to detect the effect of the severity of the event and the “true” reactivity of the VOCs.

Another implication of the PIM experimental results is that present model-based O<sub>3</sub> control strategies for Cincinnati and Mexico City should be carefully reviewed. This research has shown that the current strategy of reducing NO<sub>x</sub> emissions in both airshed regions might, in fact, result in an increase in O<sub>3</sub> concentrations.

## 4.4 References

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## Chapter 5

### Conclusions and Recommendations

Understanding the sensitivity of ozone concentrations to changes in emissions of  $\text{NO}_x$  and VOC has been an important task for formulating effective policies to reduce  $\text{O}_3$  formation in urban areas. The use of 3-D photochemical models has been the traditional scientific tool used to predict and design  $\text{O}_3$  air quality strategies. However, a number of uncertainties associated with the required input data, such as the  $\text{NO}_x$  and VOC emission inventories, have called into question the resulting predictions and, therefore, the proposed control strategies. Observation-based methods (OBMs), along with advances in the instrumentation for measuring  $\text{O}_3$  precursors, have been suggested as supplementary means for deriving the  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity and for testing the reliability of model predictions. However, OBMs have not been universally accepted within the regulatory and scientific communities because of a lack of research on the effectiveness of their methods.

In this study, the combination of an OBM with additional conceptual and data analysis was used to investigate the  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity of two dissimilar urban areas. This approach was selected with the hope that it would provide an improved diagnosis for the  $\text{O}_3$ -air-quality management process.

Measurements of  $\text{O}_3$ ,  $\text{NO}_y$ , and  $\text{NO}$  in one representative site of Cincinnati, Ohio (U.S.A.) and Mexico City, Federal District (Mexico) were used as the basis to evaluate these cities'  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity through the photochemical indicator method (PIM). Photochemical indicators are species ratios of the  $\text{O}_3$  concentration, which has built-up in the time since the precursor emissions occurred, to the total reactive nitrogen species or another secondary species. Individual species such as  $\text{NO}_y$  are also reliable indicators of the history of photochemically aged air parcels. The evaluation of the indicator species and indicator ratios was complemented with additional measurements of  $\text{HNO}_3$  and  $\text{H}_2\text{O}_2$  in Cincinnati and with 3-D simulation results of  $\text{NO}_y$ ,  $\text{NO}_z$ , PAN,  $\text{HNO}_3$ , and  $\text{H}_2\text{O}_2$  in Mexico City. In addition, a semi-empirical estimate of the  $\text{NO}_z^*$  concentration was applied in both cases. The estimated  $\text{NO}_2^*$  afternoon concentrations were within the operational uncertainty range suggested for  $\text{NO}_x$  chemiluminescent

analyzers. The consistency of the measured and the estimated indicator ratios was evaluated through the comparison of their correlation patterns with 3-D predicted correlation scatterplots for reference locations.

Assuming as “universal” the Sillman transition values for the different indicators between  $\text{NO}_x$  and VOC sensitivity conditions it was found that in both cities the chemistry of  $\text{O}_3$  formation was likely VOC-sensitive.

In both cases, the above finding was corroborated with a statistical analysis of the differences in the weekend/weekday (WE/WD) effect of morning NO emissions (with NO as a surrogate of the total  $\text{NO}_x$  emissions) on the respective maximum  $\text{O}_3$  concentrations. The statistical test was formulated based on the chemistry of  $\text{NO}_x$ -VOC sensitivity. The analysis showed that, although there was a statistically significant difference in the reduction of NO emissions on WE with respect to the WD emissions, the difference on the maximum ozone concentrations was not statistically significant in either city. In a  $\text{NO}_x$ -sensitive environment, a reduction of  $\text{NO}_x$  emissions would bring down the maximum  $\text{O}_3$  concentration. The statistical analysis showed that this was not the case. Therefore, the results of these “natural” experiments indirectly pointed to the existence of VOC-sensitive conditions in both regions.

The consistency of the PIM diagnosis in Mexico City was exercised in an intercomparison of the indicator results with previous preliminary measurements and modeling experiments done in this urban area. The results of the intercomparison showed that, independent of the methods and approaches used to arrive at the indicator ratios, all of them tended to diagnose VOC-sensitive conditions. Furthermore, an indirect comparison of the photochemical indicators results with the predictions from the 3-D photochemical model (i.e., the MCCM) led to the conclusion that the performance of the model could be improved if the actual contribution of aldehydes and ammonia emissions was included in the emissions inventory input data files.

The conceptual analysis of the descriptive parameters related to  $\text{O}_3$  formation helped to assess the behavioral differences of the photochemical process for both cities. The rate of accumulation and the peak levels of  $\text{O}_3$ , as well as the rate of depletion of NO, were higher in Mexico City than in Cincinnati. This phenomenon was seen as the result of the combination of stronger morning stagnant meteorological

conditions and larger contribution of reactive VOCs to the photochemical process in Mexico City, even though  $\text{NO}_x$  emissions for both areas seemed to be similar.

The experimental finding of VOC-sensitive conditions in both regions was interpreted as the result of the predominance of the reaction for formation of  $\text{HNO}_3$  in an environment with low free radical sources (“low” VOC emissions) but higher radical sinks (“high”  $\text{NO}_x$  emissions) in Cincinnati, and of the predominance of  *$\text{NO}_x$ -saturated conditions* (continuous “high”  $\text{NO}_x$  supply exceeding the supply of radical sources) combined with formation reactions of PAN and  $\text{HNO}_3$  in an environment with large free radical sources (“high” reactive VOC emissions) and high radical sinks (“high”  $\text{NO}_x$  emissions) in Mexico City. The combination of these results led to confirm the inconsistency of the initial emissions VOC/ $\text{NO}_x$  ratio as a tool for diagnosis of the  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity in regions with high precursors emissions densities and restrictive natural ventilation conditions.

The findings of this applied research suggest that present model-based  $\text{O}_3$  control strategies for Cincinnati and Mexico City should be carefully reviewed. This research has shown that the current strategy of reducing  $\text{NO}_x$  emissions in both airshed regions could, in fact, result in an increase in  $\text{O}_3$  concentrations.

As demonstrated by this study, the combination of the PIM with the WE/WD statistical analysis and the conceptual analysis of empirical  $\text{O}_3$ -formation related parameters has the potential to be a more comprehensive complementary tool for the identification of the  $\text{O}_3$ - $\text{NO}_x$ -VOC sensitivity of even dissimilar geographic and urban regions.

Although the results of this applied research have suggested the dominance of VOC-sensitive conditions in both regions, more research is needed to further investigate the ozone sensitivity in the entire Greater Cincinnati and the Mexico City Metropolitan Area regions. This implies the simultaneous collection of  $\text{O}_3$  and  $\text{NO}_y$  data in at least three sites of both airshed regions (one upwind and two downwind of the urban area), including at least one site with  $\text{HNO}_3$  and  $\text{H}_2\text{O}_2$  measurements. If possible, these measurements should include the determination of PAN and  $\text{NO}_3^-$  concentrations.

One effort to measure photochemical indicators simultaneously in Cincinnati and Mexico City during one period in which the conditions were relatively similar is recommended. Although summer is the season more conducive to ozone formation in southwest Ohio, this is not the case in Mexico City because this period corresponds to the rain season. The most favorable month could be May; the end of the smog season in Mexico City and the beginning of the smog season in Cincinnati.

In addition, the implementation of a 3-D photochemical modeling effort for the southwestern Ohio region (including parts of Ohio, Kentucky, and Indiana), calibrated against the results of the indicators measurements, is also recommended. The sampling and analysis of 6:00-9:00 hr VOC species concentrations in Cincinnati is mandatory for the appropriate setting of the photochemical model. In the case of Mexico City, it is recommended to calibrate the MCCM photochemical model considering scenarios with different aldehydes and ammonia contributions and also to calibrate the results against the indicator measurements. Once the model is calibrated, it would be desirable to investigate if the transition values suggested by Sillman should be adjusted for the particular geographic situation of Mexico City.

It is hoped that the present applied research will contribute to the development of better O<sub>3</sub> air quality management policies as recommended by the NARSTO for North America.

## APPENDIX A

### CLASSIFICATIONS OF O<sub>3</sub>- SENSITIVITY CHEMISTRY IN 3-D MODELING SIMULATIONS

The classifications for predicted NO<sub>x</sub>-VOC sensitivity with 3-D chemistry models are based on the relative impact of reduced NO<sub>x</sub> versus reduced VOC. These simulations predict distinct patterns of the correlation between simulated afternoon O<sub>3</sub> and NO<sub>y</sub> and between O<sub>3</sub> and NO<sub>z</sub> among other combinations. The patterns have been classified according to the differences found between simulated O<sub>3</sub> in a model base case and equivalent scenarios with 25% or 35% reductions in NO<sub>x</sub> and in anthropogenic VOC in several locations of the U.S. Locations have been classified according to the following definitions:

NO<sub>x</sub>-sensitive: A given reduction in NO<sub>x</sub> emissions causes a significant (>5 ppb) reduction in simulated O<sub>3</sub>, in both the base case and in the scenario with the same equivalent reduction in anthropogenic VOC.

VOC-sensitive: A given reduction in VOC results in an O<sub>3</sub> concentration at least 5 ppb lower than in both the base case and in the scenario with the same equivalent reduction in NO<sub>x</sub>.

Mixed or near-zero sensitive: Reduced NO<sub>x</sub> and reduced VOC emissions result in a simulated O<sub>3</sub> within 5 ppb of each other, and both scenarios result in a simulated O<sub>3</sub> lower than in the base case by at least 5 ppb.

dominated by NO<sub>x</sub>-titration: A given reduction in NO<sub>x</sub> causes a significant increase in the simulated O<sub>3</sub> in the base case by at least 5 ppb. At the same time, a given reduction in VOC results in a simulated O<sub>3</sub> not lower than 5 ppb or more relative to the base case.

#### Reference

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## **APPENDIX B**

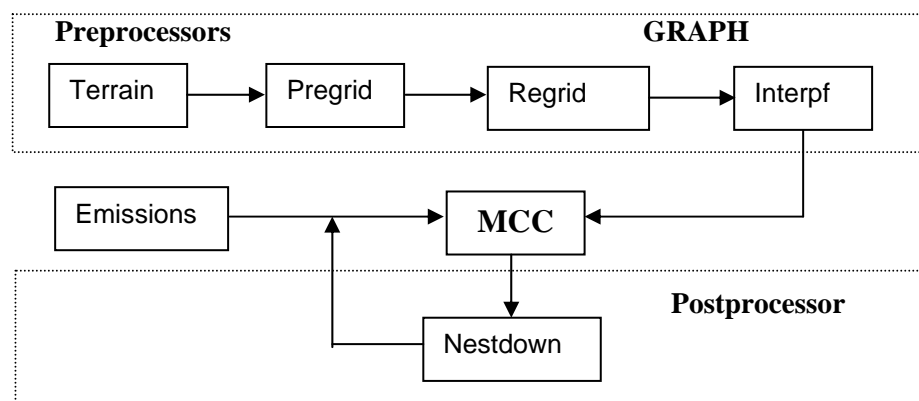
### **SUMMARY OF THE MAIN CHARACTERISTICS OF THE MCCM**

The air quality model: Multiscale Climate Chemistry Model (MCCM) is a coupled complex meteorology/chemistry model, which is based on the fifth generation Penn State/NCAR nonhydrostatic mesoscale model (MM5) and the Regional Air Deposition Model version 2 (RADM2). MCCM does include modules for meteorology, photolysis, biogenic emissions, radiation, and deposition among others. MM5 is non hydrostatic with terrain following coordinates, has a multiscale option, is capable of four-dimensional data-assimilation interface with actual weather forecast models (general circulation model and observations), contains explicit cloud schemes and multilevel soil/vegetation parameterization.

An advantage of MCCM is that MM5 is directly coupled with a chemistry-transport-model and photolysis module. Several parameterizations schemes are used for the calculation of turbulent transport variables. The parameterization scheme includes a soil/vegetation scheme that calculates the soil temperature and moisture stratification, and determines the heat and moisture fluxes at the interface between atmosphere and soil. The photolysis module uses a radiative transfer model that calculates photolysis frequencies for the 21 photochemical reactions of the gas-chemistry model at each grid point. This module considers changes in the radiation with height and changes in air composition such as ozone, aerosols and water vapor. The biogenic module calculates organic emissions of isoprene, monoterpenes, other biogenic VOCs, and nitrogen emissions by the soil. This module takes into account the effects of temperature, radiation and type of vegetation because they affect directly the emissions of biogenic VOCs. The dry deposition module calculates the elimination of trace compounds from the atmosphere depending on deposition velocity, which is calculated using aerodynamic, sub-layer and surface resistance. Because of the simultaneous on-line coupling between the meteorology and the chemistry, the MCCM gives consistent results without the necessity of interpolation as with the models that use non-coupled schemes.

The gas-phase chemistry is based on the RADM2 mechanism. This mechanism considers for inorganic species, 14 stable species, 4 reactive intermediates and 3 abundant (oxygen, nitrogen and water). The organic part is represented by 26 stable species and 16 peroxy radicals. The RADM2 mechanism represents organic chemistry through a reactivity aggregated molecular approach. Similar organic compounds are grouped together into a limited number of model groups through the use of reactivity weighting with respect to the OH radical. For the numerical solution, MCCM uses a Quasi-Steady-State Approximation (QSSA) method with 22 diagnosed, 3 constant and 38 predicted species. MM5 contains a four-dimensional data assimilation technique. This technique insures that the model solution does not diverge strongly from observations.

The input data to the MCCM consists of meteorological information, anthropogenic emissions, use of land, and topography files, which must be preprocessed in order to prepare the data for running the model. A scheme of the structure of the main preprocessors used to run the MCCM is shown in Figure B.1.



**Figure B.1.** Schematic flow diagram of the different processors used to run the MCCM.

In the *Terrain* preprocessor, the physical domains and geographical locations, as well as the nested strategy are specified. This module prepares the topography and the use of land information through an interpolation process. The purpose of the *Regrid* preprocessor is to read synoptic scale meteorological files or predicted output data and to interpolate the information to the topographical mesh previously specified in *Terrain*. The *Pregrid* preprocessor reads the files and prepare the output data in the format required by *Regrid*. Within the *Regrid* preprocessor it is included a preprocessor called *Rawins* which improves the quality of the predicted results from the first approximation through an intercomparison with observed rawinsondes. The *Interpf* program generates the initial meteorological field, and the upper and lower boundary conditions through an interpolation scheme. *GRAPH* is part of MM5 and allows one to visualize the preliminary output information. *Nestdown* is a post-processor that generates the input data for a “new” nested domain from the results of a mother-domain generated from a previous run. The *Emissions* preprocessor prepares the anthropogenic emissions from area and lineal sources in the format required by MCCM. The input VOC data must be in terms of the several reactivity family groups defined by RADM2.

The general resolution of the different meteorological domains is shown in Table B.1. The temporal resolution of the output data is generally of 1-hr.

**Table B.1. Details of the extent of domain and scale of the modeling time of MCCM.**

Domain	Cell Size Resolution (km)	Number of Cells		Modeling time (hr)
		x	y	
1	27	60	60	96
2	9	34	34	72
3	3	52	40	72

The limitations of the MCCM are all related with the quality of information, particularly, the meteorological information for the larger domain. As well as other photochemical models, the quality of the emission inventory is reflected in the modeling results. Other important operational limitation of the

MCCM is the preparation of the input data. This process requires the conversion of a number of different input data formats into specific formats required by the model.

## References

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Grell, G.A., Emeio, S., Stockwell, W.R., Schenemeyer, T., Forkel, R., Michalakes, J., Knoche, R., Seidl, W. (2000) Application of a multiscale, coupled MM5/chemistry model to the complex terrain of the VOTALP valley campaign. *Atmos. Environ.* **34**, 1435-1453.

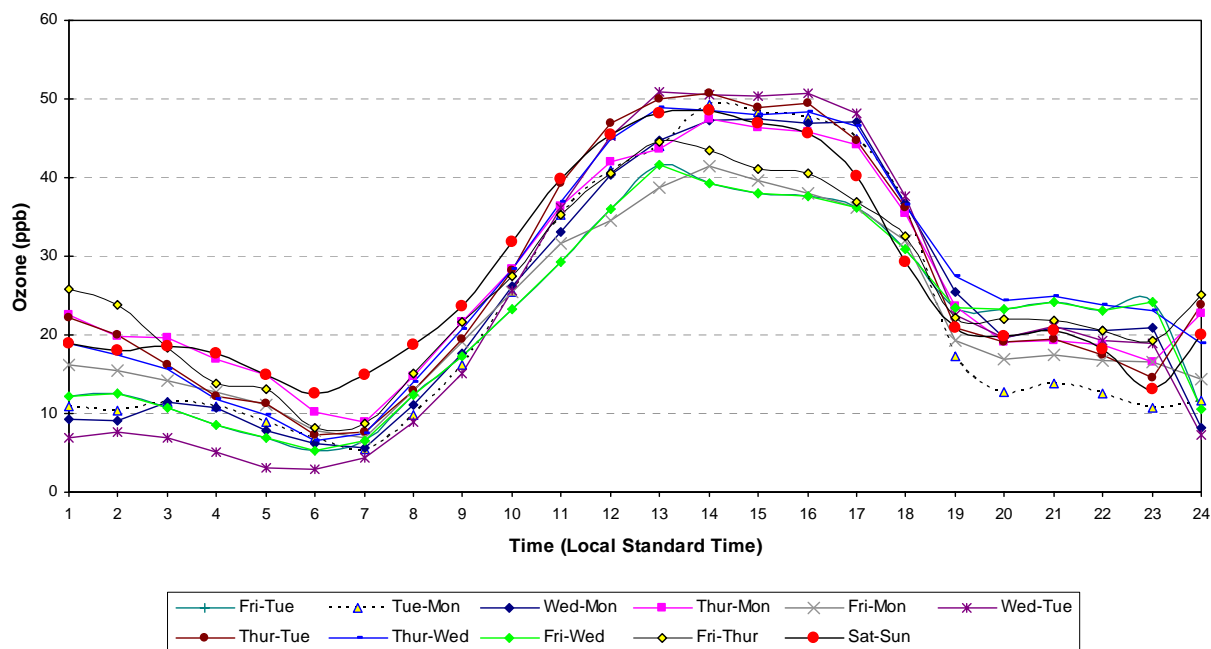
## APPENDIX C

### SUMMARY OF THE STATISTICAL ANALYSIS OF DIFFERENCES OF 1-hr MAX O<sub>3</sub> MEANS FROM A MATRIX COMBINATION OF WEEKDAYS AND WEEKEND DAYS IN CINCINNATI

**Table C.1.** Results of the tests of statistical analysis on differences (two-tail *t*-test; equal variances; *p* at a 95% confidence level), in average 1-hr max O<sub>3</sub> concentrations from a matrix combination of two WD and two WE days for data observed at the Taft Monitoring Station (Cincinnati), September 3-30, 2003.

Pair of days	Mean (S.D.) ppb	No. of Samples	Result of test of significance <sup>a</sup>	Interpretation <sup>b</sup>
Mon-Tue	49.2 (15.3)	5	<b>p = 0.947</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Mon-Wed	47.3 (20.6)	7	<b>p = 0.885</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Mon-Thur	47.4 (12.5)	7	<b>p = 0.867</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Mon-Fri	41.4 (13.2)	7	<b>p = 0.335</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Tue-Wed	50.5 (19.8)	6	<b>p = 0.840</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Tue-Thur	50.7 (8.4)	6	<b>p = 0.763</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Tue-Fri	43.7 (11.7)	6	<b>p = 0.505</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Wed-Thur	48.6 (16.)	8	<b>p = 0.999</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Wed-Fri	39.3 (21.4)	8	<b>p = 0.322</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>
Thur-Fri	43.5 (9.6)	8	<b>p = 0.416</b>	<b>H<sub>0</sub>: Accepted</b>
Sat-Sun	48.6 (14.4)	8		<b>H<sub>1</sub>: Rejected</b>

- a. If  $p \leq 0.05$ , the difference in means is significant; if  $0.05 \leq p \leq 0.10$ , the difference in means is marginally significant; if  $p \geq 0.10$ , the difference in means is insignificant.
- b.  $H_0$ :  $\mu_{O_3WE} = \mu_{O_3WD}$ ; the change in WE average maximum O<sub>3</sub> is insignificant; likely to be caused by a VOC-sensitive condition provided that the reduction in morning WE-NO<sub>x</sub> emissions is insignificant.  
 $H_1$ :  $\mu_{O_3WE} < \mu_{O_3WD}$ ; the change in WE average maximum O<sub>3</sub> is significant; likely to a NO<sub>x</sub>-sensitive condition provided that the reduction in morning WE-NO<sub>x</sub> emissions is insignificant.



**Figure C.1.** Diurnal 1-hr average  $O_3$  concentrations from a matrix combination of two weekdays and two weekend days registered at the Taft Monitoring Station (Cincinnati), September 3-30, 2003.

## APPENDIX D

### SUMMARY OF THE STATISTICAL ANALYSIS OF DIFFERENCES OF 1-hr MAX O<sub>3</sub> MEANS FROM A MATRIX COMBINATION OF WEEKDAYS AND WEEKEND DAYS IN MEXICO CITY

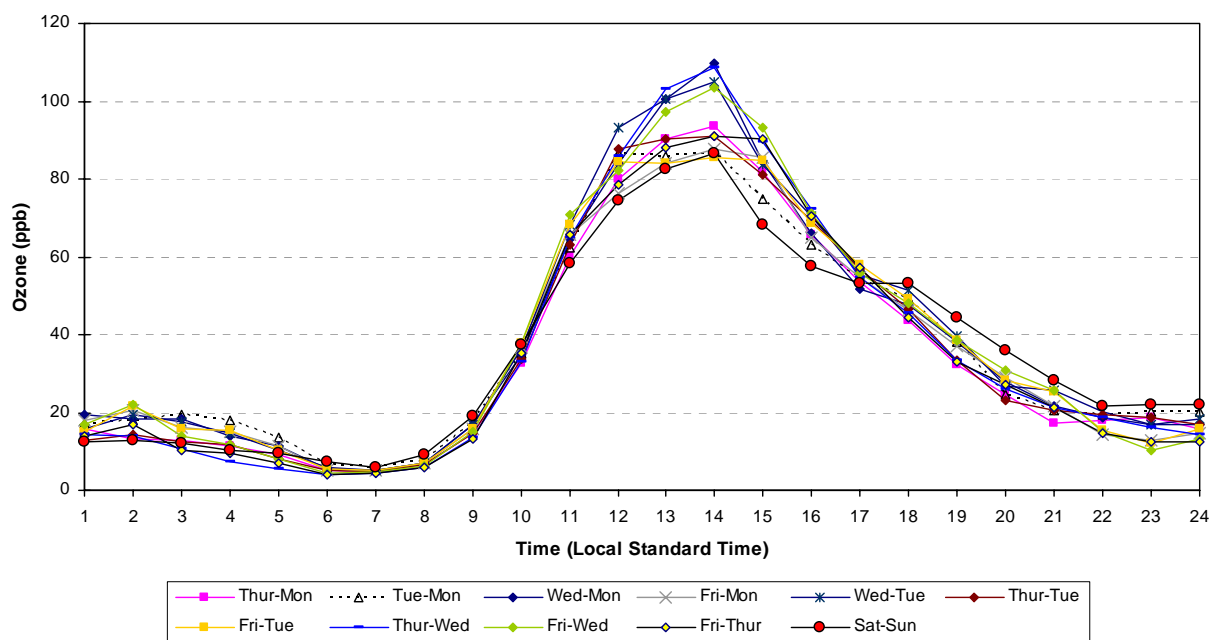
**Table D.1.** Results of the tests of statistical analysis on differences (two-tail *t*-test; equal variances; *p* at a 95% confidence level), in average 1-hr max O<sub>3</sub> concentrations from a matrix combination of two WD and two WE days for data observed at the Santa Ursula Monitoring Station (Mexico City), April 1-30, 2004.

Pair of days	Mean (S.D.) ppb	No. of Samples	Result of test of significance <sup>a</sup>	Interpretation <sup>b</sup>
<b>Mon-Tue</b>	87.1 (39.2)	7	<b>p = 0.98</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Mon-Wed</b>	109.8 (47.6)	7	<b>p = 0.33</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Mon-Thur</b>	93.5 (33.7)	8	<b>p = 0.72</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Mon-Fri</b>	87.6 (38.3)	8	<b>p = 0.96</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Tue-Wed</b>	104.8 (44.9)	8	<b>p = 0.41</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Tue-Thur</b>	90.9 (30.8)	9	<b>p = 0.81</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Tue-Fri</b>	85.7 (34.8)	8	<b>p = 0.43</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Wed-Thur</b>	108.56 (37.9)	9	<b>p = 0.27</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Wed-Fri</b>	103.3 (43.7)	9	<b>p = 0.43</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>
<b>Thur-Fri</b>	90.9 (30.9)	10	<b>p = 0.80</b>	<b>H<sub>0</sub>: Accepted</b>
<b>Sat-Sun</b>	86.6 (40.5)	7		<b>H<sub>1</sub>: Rejected</b>

<sup>a</sup>. If  $p \leq 0.05$ , the difference in means is significant; if  $0.05 \leq p \leq 0.10$ , the difference in means is marginally significant; if  $p \geq 0.10$ , the difference in means is insignificant.

<sup>b</sup>.  $H_0$ :  $\mu_{O_3WE} = \mu_{O_3WD}$ ; the change in WE average maximum O<sub>3</sub> is insignificant; likely to be caused by a VOC-sensitive condition provided that the reduction in morning WE-NO<sub>x</sub> emissions is insignificant.

$H_1$ :  $\mu_{O_3WE} < \mu_{O_3WD}$ ; the change in WE average maximum O<sub>3</sub> is significant; likely to a NO<sub>x</sub>-sensitive condition provided that the reduction in morning WE-NO<sub>x</sub> emissions is insignificant.



**Figure D.1.** Diurnal 1-hr average O<sub>3</sub> concentrations from a matrix combination of two weekdays and two weekend days registered at the Santa Ursula Monitoring Station (Mexico City), April 1-30, 2004.